Systematic Physics Constrained Parameter Estimation of Stochastic Differential Equations

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Abstract

A systematic Bayesian framework is developed for physics constrained parameter inference of stochastic differential equations (SDE) from partial observations. The physical constraints are derived for stochastic climate models but are applicable for many fluid systems. A condition is derived for global stability of stochastic climate models based on energy conservation. Stochastic climate models are globally stable when a quadratic form, which is related to the cubic nonlinear operator, is negative definite. A new algorithm for the efficient sampling of such negative definite matrices is developed and also for imputing unobserved data which improve the accuracy of the parameter estimates. The performance of this framework is evaluated on two conceptual climate models.

Keywords: Global Stability, Parameter Inference, Imputing Data, Stochastic Differential Equations, Physical Constraints, Stochastic Climate Models

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1. Introduction

In many areas of science the inference of reduced order hybrid dynamic-stochastic models, which take the form of stochastic differential equations (SDE), from data is very important. For many applications running full resolution dynamical models is computationally prohibitive and in many situations one is mainly interested in large-scale features and not the exact evolution of the fast, small scale features, which typically determine the time step size. Thus, reduced order stochastic models are an attractive alternative. Examples are molecular dynamics (20), engineering turbulence (19), and climate science (14, 15, 21).

The inference of such models has been done using non-parametric methods (39, 16, 8) from partial observations. These non-parametric methods need very long time series for reliable parameter estimates and can be used only for very low-dimensional models because of the curse
of dimension. More importantly, they do not necessarily obey conservation laws or stability properties of the full dimensional dynamical system. In many areas of science one can derive reduced order models from first principles (25, 30) such that certain fundamental properties of the full dynamics are still valid. These methods provide us with parametric forms for the model fitting. Physical constraints then not only constrain the parameters one has to estimate but they can also ensure global stability. Thus, there is a need for systematic physics constrained model and parameter estimation procedures (31, 27).

For instance, the climate system is governed by conservation laws like energy conservation. Based on this energy conservation property the normal form of stochastic climate models has been derived by (25) using the stochastic mode reduction procedure (22, 23, 24, 14, 15). This procedure allows the systematic derivation of reduced order models from first principles. This normal form provides a parametric form for parameter estimation from partial observations which we will use in this study.

The fundamental form of climate models is given by

$$\frac{dz}{dt} = F + Lz + B(z, z),$$

(1)

where $z \in \mathbb{R}^N$ denotes the N-dimensional state vector, $F$ the external forcing, $L$ a linear and $B$ a quadratic nonlinear operator. The nonlinear operator $B$ is conserving energy $z \cdot B(z, z)$. For current climate models $N$ is of the order of $10^6 - 10^8$. This shows that running complex climate models is computationally expensive. But for many applications like extended-range (periods of more than 2 weeks), seasonal and decadal climate predictions one is only interested in the large-scale circulation of the climate system and not whether there will be a cyclone over the North Atlantic on a particular day next year. The large-scale circulation can successfully be predicted using reduced order models (36, 1, 14, 15, 21).

The stochastic mode reduction procedure (22, 23, 24) provides a systematic framework for deriving reduced order climate models with a closure which takes account of the impact of the unresolved modes on the resolved modes. In order to derive reduced order models one splits the state vector $z = \begin{pmatrix} x \\ y \end{pmatrix}$ into resolved $x$ and unresolved $y$ modes. The stochastic mode reduction procedure now enables us to systematically derive a reduced order climate model which only depends on $x$

$$dx = \left( \tilde{F} + \tilde{L}x + \tilde{B}(x, x) + M(x, x, x) \right) dt + a(x, \sigma) dW,$$

(2)

where $M$ denotes a cubic nonlinear term, $W$ is the Wiener process and $\sigma$ the diffusion parameters.

In this study we will develop a systematic Bayesian framework for the efficient estimation of the model parameters using Markov Chain Monte Carlo (MCMC) methods from partial observations. We are dealing with partial observations because we now only have knowledge of the few resolved modes $x$ and are ignorant about the many unresolved modes $y$.

Stochastic climate modeling is a complex problem and empirically estimating the parameters poses several problems. First, the nonlinearity of climate models requires an approximation of the likelihood function. While it can be shown that this approximation converges to the true likelihood, this is not necessarily the case for real world applications. Here we develop a MCMC algorithm for the first time for stochastic climate models and demonstrate that this algorithm performs well. Second, the nonlinearity of the problem causes the space of parameters leading to stable and physical meaningful solutions to become small as the dimension of the problem increases. We show that a lot of the posterior mass is on parameter values which lead to solutions
exploding to infinity in finite time. To solve this problem we derive global stability conditions. These conditions take the form of a negative definite matrix. Hence, we devise a novel sampling strategy based on sampling non-negative matrices. We show that this sampling strategy is computationally efficient and leads to stable solutions.

In section 2 we introduce stochastic climate models and derive conditions for global stability. Previous studies have shown that reduced climate models with quadratic nonlinearity shown unphysical finite time blow up and long time instabilities (18, 26, 27, 40). Here we use the normal form of stochastic climate models (25) and derive sufficient conditions for global stability for the normal form of stochastic climate models. These normal form stochastic climate models have cubic nonlinearities. The here derived stability condition is more general than the one in Majda et al. (25). In section 3 we develop a Bayesian framework for the systematic estimation of model parameters using physical constraints. Here we develop an efficient way of sampling negative-definite matrices. Without this constrain the MCMC algorithm would produce about 40% unphysical solutions which is clearly very inefficient. Here we also demonstrate that for these kinds of SDEs imputing data improves the parameter estimates considerably. In section 4 we demonstrate the accuracy of our framework on conceptual climate models. We summarize our results in section 5.

2. Stochastic Climate Models and Global Stability

Here we study the following D dimensional normal form of stochastic climate models (which has the same structural form as Eq. (2), see (25)): 

\[ dx_i = \left( \alpha_i + \sum_{j=1}^{D} \beta_{i,j} x_j + \sum_{j=1}^{D} \sum_{k=1}^{j} \gamma_{i,j,k} x_j x_k + \sum_{j=1}^{D} \sum_{k=1}^{j} \sum_{l=1}^{k} \lambda_{i,j,k,l} x_j x_k x_l \right) dt \]  

\[ + \sum_{j=1}^{D} a_{i,j} dW_j + \sum_{j=1}^{D} \sum_{k=1}^{j} b_{i,j,k} x_j dW_k. \]

which we write for convenience in a more compact form 

\[ dx = \mu(x, A) dt + a(x, \sigma) dW. \]  

The parameters \( \alpha, \beta, \gamma, \) and \( \lambda \) are written as one matrix \( A \in \mathbb{R}^{D \times P}. \) We allow for the inclusion of all possible linear, quadratic and cubic terms with forcing terms entering first, followed by linear, quadratic and then the cubic terms. We include them into a matrix \( A \) as \( A_{i,1} = \alpha_i, A_{i,j+1} = \beta_{i,j}, A_{i,(j+1)2+k+D+1} = \gamma_{i,j,k} \) and \( A_{i,(j+1)2+k+D+1} = \lambda_{i,j,k,l}. \) The index function for the cubic term is given by 

\[ f(j, k, l) = 1 + D + \frac{D(D+1)}{2} + \frac{j(j-1)(j+1)}{6} + \frac{k(k-1)}{2} + l. \]  

Global stability implies that the cubic terms acts as a nonlinear damping. For climate models energy conservation implies global stability. However, in general global stability does not necessarily imply energy conservation. The energy equation can be written as 

\[ \frac{1}{2} \frac{dE}{dt} = \sum_{i=1}^{D} x_i \frac{dx_i}{dt} = \sum_{i=1}^{D} \sum_{j=1}^{D} \sum_{k=1}^{j} \sum_{l=1}^{k} A_{i,j,k,l} x_j x_k x_l. \]
Here we only consider the cubic term because this term will ultimately determine global stability. Majda et al. (25) have shown that the normal form of stochastic climate models allows for linearly unstable modes. These linearly unstable modes are associated with important weather systems and waves and are an intrinsic and important part of climate models. Once these linearly unstable modes have reached a certain amplitude the nonlinear cubic terms will govern their evolution and, thus, ensure global stability.

We consider now the vector \( \mathbf{v} \) with \( \frac{1}{2}(D+1) \) components of the form \( v_i = x_ix_j \) with \( 1 \leq j \leq i \leq D \). Now we find a negative definite matrix \( \mathbf{M} \) such that (25)

\[
v^T \mathbf{M} \mathbf{v} = \frac{1}{2} \frac{dE}{dt} \leq 0. \tag{7}\]

A sufficient solution is as follows: Let matrix \( \mathbf{M} \in \mathbb{R}^{(D+1)D/2 \times (D+1)D/2} \) be

\[
\mathbf{M} = \begin{pmatrix}
A_{i,f(i,j)} & \ldots & A_{i,f(i,k)} & \ldots & A_{i,f(i,D)} \\
0 & \ldots & 0 & \ldots & 0 \\
A_{i,f(i,j)+A_{i,f(i,k)}} & \ldots & A_{i,f(i,j)+A_{i,f(i,k)}} & \ldots & A_{i,f(i,j)+A_{i,f(i,D)}} \\
0 & \ldots & 0 & \ldots & 0 \\
\end{pmatrix} \tag{8}
\]

where \( 1 \leq j \leq i \leq D \) and \( 1 \leq l \leq k \leq D \). While this solution is not necessarily unique the imposition of this constraint is still necessary in order to reduce the amount of parameter values leading to unstable models and, thus, to reduce the computational expense of the parameter inference. As we will show below, not imposing this constraint will lead to unstable and unphysical solutions in about 40% of parameter values in our MCMC scheme.

In summary, the stochastic climate model in Eq. (4) is globally stable if the tensor \( \mathbf{M} \) is negative definite and \( \mathbf{M} \) determines the components of the cubic operator \( \lambda \). This is an important result for the constrained parameter estimation of stochastic climate models (25, 31).

3. Physics Constraint Parameter Sampling

We use a Markov Chain Monte Carlo algorithm for the parameter inference which was proposed by (7) and (17). We use this approach because of its flexibility. For instance, the exact algorithms by (5) and (6) cannot be easily applied to multidimensional. Furthermore, the exact algorithms require that the drift function must be the gradient of a potential; stochastic climate models cannot be written in such a form. The novel aspect of our MCMC algorithm is the physics constraint sampling. This algorithm overcomes the dependency between the diffusion parameters and the missing data by changing variables to the underlying Brownian motion \( \mathbf{W} \in \mathbb{R}^d \) and conditioning on this when performing the parameter update.

Assuming we have the SDE in Eq. (4) and \( \mathbf{a}(\mathbf{x}_t, \sigma) \in \mathbb{R}^{d \times d} \) is invertible in its first argument then

\[
d\mathbf{W}_t = \mathbf{a}^{-1}(\mathbf{x}_t, \sigma)(d\mathbf{x}_t - \mu(\mathbf{x}_t, \mathbf{A})dt) \tag{9}\]

is a \( d \)-dimensional Brownian motion. We define a process \( \mathbf{Z} \), which conditions on the endpoint \( x_T \), by

\[
d\mathbf{x}_t = \mathbf{a}(\mathbf{x}_t, \sigma)d\mathbf{Z}_t + \frac{x_T - x_i}{T-t}dt, \quad x_0 = x_0, \tag{10}\]

where \( T \) is the next observation time. In discrete time the transformation is

\[
X_{i+1} = X_i + \mathbf{a}(X_i, \sigma)(Z_{i+1} - Z_i) + \frac{x_T - X_i}{m - i} \tag{11}\]
where $m - 1$ is the number of imputed points between two observations.

Now we write down an algorithm that does not degenerate as $m$ increases. It is based on the reparametrisation of the conditional density of the data

$$p(X|A, a) = p(g(Z, a)|A, a) \left| \frac{\partial g}{\partial Z} \right|$$

(12)

where the Jacobian determinant is equal to the product

$$\left| \frac{\partial g}{\partial Z} \right| = \prod_{i=1}^{N-1} |a(X_i, \sigma)|,$$

(13)

and $X_i = g(Z_i, a)$ denotes the transformation of all the data.

The density on the right hand side can be written as a Radon-Nikodym derivative of the law of $Z$ with respect to Brownian motion and so the dominating measure is parameter free $[9]$. This means that the Metropolis-Hastings (MH) acceptance probability will have non zero numerator and denominator.

We sample $\sigma$ according to Algorithm $[1]$. We use zero-based numbering and consider there to be $N$ observations and therefore, $N - 1$ observation intervals indexed $0 \ldots N - 2$. We assume that the inter-observation times $\Delta t$ are all equal and that there are $m - 1$ imputed points per interval, giving a time interval of $\delta = \Delta t/m$. We use the notation $X_\delta = X_{\delta t}$ and $\mu_j = \mu(X_{\delta t}, A)$. The extension to variable inter-observation times is straightforward. For simplicity, we write the algorithm for perfect observation of the system so that $X_{\text{im}}, i = 0, \ldots, N - 1$ are fixed. In Algorithm $[1]$, $\phi$ denotes a Gaussian distribution and $q$ the Gaussian proposal density.

**Algorithm 1** Sample parameters entering the diffusion matrix.

- Draw $\sigma^* \sim q(\sigma^*|\sigma)$
- Initialize $\alpha = \log(q(\sigma^*|\sigma)) - \log(q(\sigma^*|\sigma)) + \log(p(\sigma^*)) - \log(p(\sigma))$
- for $i = 0$ to $N - 2$
  - for $j = 0$ to $m - 2$
    - $Z_{\text{im}+j+1} = Z_{\text{im}+j} + a^{-1}(X_{\text{im}+j}, \sigma)(X_{\text{im}+j+1} - X_{\text{im}+j} - \frac{X_{\text{im}+1} - X_{\text{im}}}{m-1})$
    - $X_{\text{im}+j+1} = X_{\text{im}+j} + \frac{X_{\text{im}+1} - X_{\text{im}}}{m-1} + a(X_{\text{im}+j}, \sigma^*)(Z_{\text{im}+j+1} - Z_{\text{im}+j})$
    - $\alpha = \alpha + \log(\phi(X_{\text{im}+j+1}; X_{\text{im}+j}, \mu, \Delta t, \delta \Sigma_{\text{im}+j}) + \log |a(X_{\text{im}+j}, \sigma^*)|$
    - $- \log(\phi(X_{\text{im}+j+1}; X_{\text{im}+j}, \mu, \Delta t, \delta \Sigma_{\text{im}+j}) - \log |a(X_{\text{im}+j}, \sigma)|$
  - end for
- end for
- Set $[\sigma, X] = [\sigma^*, X^*]$ with probability $\min(1, \exp(\alpha))$ else retain $[\sigma, X]$.

To update missing data between observations we use an independence sampler as in $[35]$ using the proposal process

$$dX^* = \xi(X^*, X_{\delta t})dt + a(X^*, \sigma)dW^*,$$

(14)
where \( X_T \) is the next observation, \( X^* \) the proposed data, and where \( \alpha(X^*, \sigma) \) is the same diffusion function as that in Eq. 3. \( \xi \) denotes the modified linear bridge (see below). This process will have measure that is absolutely continuous with respect to the target process in Eq. (4) because of their common diffusion function. To update all of the missing data we propose a block at a time from Eq. (14) and then accept this according to the MH ratio. If the inter-observation interval is large then the acceptance rate may become very low and so one may sub-sample smaller blocks.

For some interval \( i \) we set \( X^*_0 = X_{\text{im}} \) and \( X^*_m = X_{\text{im}+1 \text{m}} \), then we propose \( X^*_j : X^*_{m-1} \) and accept or reject using the MH acceptance probability

\[
\alpha = \frac{\rho(X^*_m | X^*_{m-1}, A) \prod_{j=0}^{m-2} \rho(X^*_{j+1} | X^*_j, A) q_0(X_{\text{im}+j+1 | X_{\text{im}+j}, \xi, \sigma)} \rho(X_{\text{im}+j+1 | X_{\text{im}+j}, A) q_0(X^*_j | X^*_j, \xi, \sigma)} \right)
\]

where \( \rho_0 \) is the transition density of the target

\[
dX_j = \mu(X_i, A)dt + a(X_i, \sigma)dW_j, \quad X_0 = x_0, \quad t \in [0, T]
\]

over time interval \( \delta \) and \( q_0 \) is the transition density of the proposal. We choose proposal processes so that given \( X^*_j, X^*_{j+1} \) is approximately Gaussian distributed. Eq. (14) is not a true Gaussian process because of the state dependent noise term. Details for updating the missing data are given in Algorithm 2.

**Algorithm 2** Sample missing data between observations.

```
for i = 0 to N - 2 do
    Set \( X^*_0 = X_{\text{im}} \)
    Set \( \alpha = 0 \)
    for j = 0 to m - 2 do
        \( X^*_{j+1} \sim q_0(X^*_j | \xi(X^*_j, X^*_{m-1}), \sigma) \)
        \( \alpha = \alpha + \log(\rho(X^*_j+1 | X^*_j + \delta \mu^*, \delta \Sigma^*) + \log(q_0(X_{\text{im}+j+1 | X_{\text{im}+j}, X_{\text{im}+1}, \sigma) )
        - \log(\rho(X_{\text{im}+j+1 | X_{\text{im}+j} + \delta \mu_{im+j}, \delta \Sigma_{im+j}) - \log(q_0(X^*_j+1 | X^*_j, \xi(X^*_j, X^*_{im+1}), \sigma))
        end for
    end for
    if \( \exp(\alpha) > \mathcal{U}(0, 1) \) then
        for j = 0 to m - 2 do
            \( X^*_{im+j+1} = X^*_{j+1} \)
        end for
    end if
```

Algorithms 1 and 2 are combined with standard MH updates for the parameters \( A \) entering into the drift function. One could use Random-Walk proposals but in our case of polynomial models it is more efficient to implement another Gibbs sampling step. Repeatedly alternating between these three steps will produce MCMC samples that can be used to estimate the parameters. In practice we increase the amount of missing data \( m \) until we see convergence in the marginal distributions of the parameters.

6
3.1. Sampling of Diffusion Paths

Now we discuss how we simulate diffusion paths from Eq. \( \text{[3]} \) that are conditioned upon start \( X_0 = x_0 \) and end \( X_m = x_m \) points. The target density, with respect to Lebesgue measure, is the product

\[
\pi(X_1 \cdots X_{m-1}|X_0 = x_0, X_m = x_m; \mathcal{A}, \sigma) = \prod_{i=0}^{m-1} p_\delta(\tau_{k+1}, x_{k+1} | \tau_k, x_k; \mathcal{A}, \sigma),
\]

where now \( p_\delta \) is the transition density for the process in Eq. \( \text{[3]} \) and is, in practice, approximated by the Euler transition density. We consider the total time interval \( \tau_m - \tau_0 = \Delta \) divided into \( m \) equidistant sub-intervals so that \( \tau_{k+1} - \tau_k = \Delta/m = \delta \).

Having \( N \) observations, for each of \( i = 0, 1, 2, \ldots N-1, X_m \) is an observation. Between every pair of observations the diffusion bridge will need to be simulated. We use an independence sampler with proposal density of the form \( q(X'|X) = q(X^*) \). Here, we consider proposal processes of the form of Eq. \( \text{[14]} \).

We use a Modified Linear Bridge proposal for sampling of parameters of the drift of equations of the form

\[
dZ_t = (Q(X)Z_t + r(X, t))dt + \Sigma(X)dB_t,
\]

This is a local linearization of the nonlinear diffusion over a small time window \( \text{[20], [38]} \). First we construct bridge distributions for general multivariate linear diffusions \( \text{[4]} \). If at time \( s \) we have \( X_s = d \) and at time \( T, X_T = e \) then the distribution of \( X_t \) for \( 0 \leq s < t \leq T \) can be shown to be Gaussian with mean

\[
v_{d,e}(s, t) = \Gamma(t, T)\Gamma(s, T)^{-1} m^*_d(s, t) + \Gamma(s, T)^T \Gamma(s, T)^{-1} m^*_e(t, T),
\]

where

\[
\Gamma(s, t) = \int_s^t e^{(t-u)Q} \Sigma_s^T e^{(u-t)Q^T} du,
\]

\[
m^*_d(s, t) = x + \int_s^t e^{(t-u)Q} r(u)du \quad \text{and} \quad m^*_e(s, t) = x - \int_s^t e^{(t-u)Q} r(u)du.
\]

The covariance matrix is given by

\[
\Sigma(s, t) = \Gamma(t, T)\Gamma(s, T)^{-1} \Gamma(s, t).
\]

In general this matrix can be computed as follows: if we diagonalize \( Q \) so that \( Q = U\Lambda U^{-1} \) then compute the matrix \( A \) with components

\[
A_{ij} = \left( U^{-1} \Sigma_i^T U^{-1} \right)_{ij} \left( e^{(t-u)\Lambda_i} - e^{(t-u)\Lambda_j} \right),
\]

then

\[
\Gamma(s, t) = U\Lambda U^T.
\]

The proposal distribution is given by

\[
q(X_k+1|X_k, X_m, A) = \phi(X_k+1; \nu_{X_kX_m}(k\delta, (k+1)\delta), \Sigma_{\nu}(k\delta, (k+1)\delta)),
\]

where \( \nu_{X_kX_m}(j\delta, (j+1)\delta) \) and \( \Sigma_{\nu}(j\delta, (j+1)\delta) \) are given in Eqs. \( \text{[15]} \) and \( \text{[20]} \) respectively.

In contrast to a linear bridge sampler here we update at each imputed point. This means recomputing the matrices \( \Gamma(s, t) \) at each point although \( Q \) and \( U \) are only calculated once.
3.2. Inference for Drift Parameters

Now we give details of the computational implementation of the sampling of parameters in the drift function. Since the drift parameters enter linearly we can construct a Gibbs sampler where their conditional posterior is Gaussian. This greatly improves the mixing of the Markov Chain.

3.2.1. Gibbs Sampler

Consider $N$ observations with time interval $\delta$. We set $Y_t = X_{t+1} - X_t$ and let $U \in \mathbb{R}^{N-1 \times P}$ be the design matrix of the data, scaled by $\delta$. The columns of $U$ are indexed in the same way as the columns of the parameter matrix $A$.

The log likelihood can be written

$$L(A; \mathbf{X}) = -\frac{1}{2} \sum_{i=1}^{N-1} [\Sigma_i] - \frac{1}{2} \sum_{i=1}^{N-1} \sum_{j=1}^{D} \left( Y_i - \sum_{k=1}^{P} U_{ik} A_{jk} \right) \Sigma_{ij}^{-1} \left( Y_i - \sum_{k=1}^{P} U_{ik} A_{jk} \right),$$

where the instantaneous covariance matrix $\Sigma_i$ is computed from $\Sigma_i^{1/2} = (d_{ij} + \sum_{l=1}^{D} e_{ij} X_{il}) \Lambda^{1/2}$.

We have $DP$ parameters to infer in the matrix $A$. We use a zero mean Gaussian prior with covariance matrix $\Gamma \in \mathbb{R}^{DP \times DP}$. Let $\Lambda \in \mathbb{R}^{DP \times DP}$, be a matrix with components

$$\Lambda_{i(i-1)P+j, (k-1)P+l} = \sum_{i=1}^{N-1} U_{il} \Sigma_{ik}^{-1} U_{jk} + \Gamma_{i(i-1)P+j, (k-1)P+l}^{-1}$$

where $i, k = 1 \ldots D$ and $j, l = 1 \ldots P$. Let $e \in \mathbb{R}^{DP}$ with components

$$e_{i(i-1)P+j} = \sum_{i,k} U_{i,j} \Sigma_{ik}^{-1} Y_{ik}.$$

The posterior mean $\mu_{i(i-1)P+j}$ of $A_{ij}$ is given by the solution of $\Lambda \mu = b$ and the posterior covariance is $\text{Cov}(A_{ij}, A_{kl}) = \Lambda_{i(i-1)P+j, (k-1)P+l}^{-1}$.

We applied the above Gibbs sampler to a large data set from a two dimensional model of the form of Eq. (23) with random values for the diffusion parameters. We chose a fine observation interval of $\delta = 10^{-2}$ and long observation period $T = 10,000$. Fig. 3 displays the trace plots for all 20 parameters (note that the indices are from 0 rather than 1 as in the text). Using this large data set the algorithm is able to reproduce the true values shown in red.

We performed a further simulation study to test the dependence of the posterior estimates upon the data set used. We inferred all of the drift parameters for a simple two dimensional model of the form of Eq. (23) using data sets of length $T = \{10, 100, 1000\}$ and with observation interval $\Delta = \{0.1, 0.01, 0.001\}$. Note that the diffusion function is arbitrary in this model. The results are shown in Table 1. For each parameter we estimated the posterior mean and the posterior 10th-90th percentiles. A cell is colored blue where the true value falls within this range. Note that if we were using the true likelihood, rather than an approximation, then we would expect there to be around 80% blue boxes. The error of the estimates for each data set can be quantified using the quadratic Posterior Expected Loss (PEL) function

$$f(\hat{\theta}, X_{obs}) = \int_{\Theta} (\theta - \hat{\theta})^2 \pi(\theta | X_{obs}) d\theta,$$
Figure 1: Output of a Gibbs sampler for 20 drift parameters of a two dimensional model from Eq. (3). The observation interval is $\delta = 10^{-3}$ and $T = 10,000$. The true values are shown in red.
<table>
<thead>
<tr>
<th>$T = 10$</th>
<th>$T = 100$</th>
<th>$T = 1000$</th>
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</thead>
<tbody>
<tr>
<td>0.1</td>
<td>0.01</td>
<td>0.001</td>
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<tr>
<td>0.01</td>
<td>0.001</td>
<td>0.0001</td>
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</tbody>
</table>

| $A_{00} = 0$ | $\{0.786, 0.786\}$ | $\{0.999, 0.999\}$ |
| $A_{01} = 5$ | $\{1.0, 1.0\}$ | $\{1.0, 1.0\}$ |
| $A_{02} = 0$ | $\{0.999, 0.999\}$ | $\{1.0, 1.0\}$ |
| $A_{10} = 0$ | $\{0.999, 0.999\}$ | $\{1.0, 1.0\}$ |
| $A_{11} = 0$ | $\{0.999, 0.999\}$ | $\{1.0, 1.0\}$ |
| $A_{20} = 0$ | $\{0.999, 0.999\}$ | $\{1.0, 1.0\}$ |
| $A_{21} = 0$ | $\{0.999, 0.999\}$ | $\{1.0, 1.0\}$ |

Table 1: Drift parameter estimates for a two dimensional cubic model with arbitrary diffusion function. On the left is the true value of the parameter. The length of the data set used for the inference is labeled as $T$ and the observation interval is $\Delta = (0.1, 0.01, 0.001)$. In each cell the parameter is estimated from the posterior mean and in brackets is shown the 10th-90th percentiles of the posterior. The blue coloring is where the true value falls in each case. The bottom of the table shows the Posterior Expected Loss in each case.
where $\theta$ represents all of the parameters, $\hat{\pi}$ is the estimated posterior distribution and $\theta^*$ is the true value of the parameter.

We performed a test with both the Gibbs sampler and data imputation. In Tab. 2 the data is observed at interval $\Delta = 0.1$. The smaller intervals $\Delta = \{0.01, 0.001\}$ are obtained by imputing data with $m = \{10, 100\}$ respectively. The table shows that imputing data approximately doubles the Posterior Expected Loss. As expected the confidence intervals are broader but with more imputed data the algorithm can recover the true values. This shows that our data imputing strategy successfully improves the parameter estimates.

Our aim is to infer models that can be used for prediction. This can be problematic when dealing with non-linear models as some (generally unknown) regions of the parameter space will give solutions that explode to infinity with probability 1. This is a particular problem when, as exemplified by Table 1, large amounts of data are needed to regain the true values.

To demonstrate this problem we performed an inference on a two dimensional cubic model using $N = 1,000$ observations at $\Delta = 0.1$. For each inferred parameter value we then simulated the solution for $T = 100$. After this time we recorded whether the solution retained finite values or had exploded. The marginal posterior distributions of the cubic parameters are plotted in Fig. 2. Each plot shows two histograms: one in blue records the distribution of stable parameter values and in red are those that exploded. Notice that, when looking at the marginal distributions, the stable and unstable regions largely overlap; it is difficult to separate the two regions. In this case 40% of values were unstable. Tests (not shown) indicate that this is an even bigger problem in higher dimensions. Therefore, it is essential to use constraints on the parameter space to enable only physically meaningful solutions. The necessary conditions have been derived in section 2.

Thus, as shown above, when updating the drift parameters we ensure that $M$ is negative definite. In practice it is sufficient to check only whether the symmetric part $(M + M^T)/2$ is negative-definite. In the next section we will develop a systematic way of sampling negative definite matrices.

### 3.3. Sampling Negative Definite Matrices

To sample negative definite matrices we use the Component Wise algorithm (31). Here we sample the density of a $n \times n$ matrix $M$, with normally distributed components, subject to the constraint that it is negative definite. This algorithm updates $M$ component wise and is based on the following property: an $n \times n$ matrix is negative definite if and only if all $k \leq n$ leading principal minors obey $|M^{(k)}|(-1)^k > 0$. The $k$th principal minor is the determinant of the upper left $k \times k$ sub-matrix. Consider the parameters along the main diagonal. As they only enter $M$ once each will have an associated upper bound. The Algorithm works by calculating the upper bound associated with the constraints from each principal minor. It does this to find the least upper bound and thereby the truncation point of the normal distribution.

Here, $\mathcal{N}(\mu, u, \sigma^2)$ is the right truncated normal distribution with mean $\mu$, standard deviation $\sigma$ and upper bound $u$. The off-diagonal parameters enter twice so there will be a quadratic function determining their limits for each leading principal minor. For parameters in element $M_{ij}^{(k)}$ there will be an associated quadratic relation $a_{ij}^{(k)}M_{ij}^{2} + b_{ij}^{(k)}M_{ij} + c_{ij}^{(k)} = 0$ where the coefficients
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Table 2: Drift parameter estimates for a two dimensional cubic model with arbitrary diffusion function. On the left is the true value of the parameter. The data used is the same as that of Table 1 sampled at the \( \Delta = 0.1 \) interval. In this case data is imputed to obtain the intervals \( \Delta = 0.01, 0.001 \). The bottom of the table shows the Posterior Expected Loss in each case.
Figure 2: Marginal distributions of the cubic parameters inferred from a data set with $N = 1,000$ observations at interval $\Delta = 0.1$. The blue histogram shows the parameters that gave stable solutions to the SDE, while the red is for those that gave unstable solutions. The true values are given by the red lines.
Algorithm 3 Sample parameters along diagonal

for $i = 1$ to $n$

for $j = i$ to $n$

$U_i = 0$

$x = \left( \sum_{k=l}^{j} (\frac{1}{\sqrt{\det(M_{l:j,k:j})}}) \right) f_j(M_{l:j,k:j})$

end for

if $x < U_i$ then

$U_i = x$

end if

$M_{ij} \sim \mathcal{N}(\mu_i, U_i, \sigma_i^2)$

end for

are functions of the other parameters. These coefficients are found to be

$$a_{ij}^{(k)} = -|M_{l:j,k:j}|$$  \hspace{1cm} (28a)

$$b_{ij}^{(k)} = (-1)^{i+j} \sum_{l=1}^{j-1} M_{l:k}(-1)^{i+j+k}|M_{l:j,k:j}|$$  \hspace{1cm} (28b)

$$+ (-1)^{i+j} \sum_{k=1}^{N} M_{k:j}(-1)^{i+j+k}|M_{l:j,k:j}|$$  \hspace{1cm} (28c)

$$+ (-1)^{i+j} \sum_{k=1}^{N} M_{k:j}(-1)^{i+j+k}|M_{l:j,k:j}|$$  \hspace{1cm} (28d)

$$c_{ij}^{(k)} = \sum_{k=1}^{N} M_{l:k}(-1)^{i+j+k} \left( \frac{1}{\sqrt{\det(M_{l:j,k:j})}} \right)^2$$  \hspace{1cm} (28e)

$$\sum_{l=1}^{N} M_{l:k}(-1)^{i+j+k}|M_{l:j,k:j}|$$  \hspace{1cm} (28f)

where $|M_{l:j,k:j}|$ represents the $k$th principal minor with rows $i$ and $j$ and columns $l$ and $k$ removed. For each component $M_{ij}$ this quadratic form can be solved to give upper and lower bounds on the parameter. The matrix $M$ can be cycled through updating each parameter in turn. Algorithm 4 describes the sampling of off-diagonal elements using the coefficients in Eq. (28g).

Here, the notation, $\mathcal{N}(\mu, \sigma^2)$ refers to the doubly truncated normal distribution with mean $\mu$, left truncation $u^-$, right truncation $u^+$ and standard deviation $\sigma$.

To simulate from truncated normal distributions we are using the inverse Cumulative Density Function (CDF) method. One simply calculates the corresponding CDF of the lower and upper
Algorithm 4 Sample parameters off diagonal

for $i = 1$ to $n$ do
  for $j = i + 1$ to $n$ do
    $u^* = \infty$
    $u^- = -\infty$
    for $k = j$ to $n$ do
      Calculate $a_{ij}^{(k)}$, $b_{ij}^{(k)}$ and $c_{ij}^{(k)}$ and solve $a_{ij}^{(k)}x^2 + b_{ij}^{(k)}x + c_{ij}^{(k)} = 0$.
      Set $mn = \min(x_1, x_2)$ and $mx = \max(x_1, x_2)$
    end for
    if $mx < u^*$ then
      $u^* = mx$
    end if
    if $mn > u^-$ then
      $u^- = mn$
    end if
    $M_{ij} \sim \mathcal{N}(\mu_{ij}, u^-, u^+, \sigma_{ij}^2)$
  end for
end for

boundaries and then draws a uniform random variable between these numbers. Inverting the CDF gives a random variable from the Normal distribution restricted to this region.

For our problem we use the rejection sampler method proposed by \textsuperscript{33}. This method draws uncorrelated samples directly from the target density. Rejection sampling from a distribution $h(x)$ is based on a proposal distribution $g(x)$ such that $h(x) \leq Mg(x)$ holds for some constant $M$ and all of the support of $h(x)$. For a one sided truncated Normal the exponential distribution is a good proposal. First it is translated to coincide with the truncation point then the rate parameter is optimized in order to closely match the tail of the Normal distribution.

$$g(z; \alpha, \mu^-) = \alpha \exp(-\alpha(z - \mu^-))I_{z \geq \mu^-}$$  \hfill (29)

The optimal value of $\alpha$ is calculated by maximizing the expected acceptance probability and is shown to be

$$\alpha^*(\mu^-) = \frac{\mu^- + \sqrt{(\mu^-)^2 + 4}}{2}$$  \hfill (30)

More details are given in \textsuperscript{33}.

We performed a numerical study to compare the standard Normal and Exponential proposals. The efficiency of proposing $x$ from the standard normal then accepting if $x > \mu^-$ falls to approximately 0.023 while for the optimized exponential proposal is approximately 0.5.

For the doubly truncated Normal one uses either an exponential or uniform distribution, as a proposal, depending upon the size of the truncated region. If the following holds

$$u^+ > u^- + \frac{2 \sqrt{\sigma}}{u^- + \sqrt{(u^-)^2 + 4}} \exp(\frac{(u^-)^2 - u^- \sqrt{(u^-)^2 + 4}}{4})$$

then it can be shown that the exponential is more efficient, otherwise the uniform is better \textsuperscript{33}. Fig. 3 shows both the uniform and exponential approximations for both cases.
Figure 3: Doubly truncated normal distribution. The top figure has \( u^- = 2 \) and \( u^+ = 3 \) and is better approximated with the exponential distribution. The bottom figure has \( u^- = 2 \) and \( u^+ = 2.5 \) and the uniform is more efficient.
We use Algorithms 3 and 4, along with the methods of sampling truncated Normal variables, to sample the stability matrix $M$. We tested this algorithm on a three dimensional model with $T = 100$. In this case the dimension of $M$ is $n = D(D+1)/2 = 6$. Note that this MCMC algorithm still mixes well. Fig. 4 compares the posterior distributions estimated from the Component Wise Algorithm and the standard Gibbs sampler. Notice the large differences between the distributions in each case. Similar results (not shown) are obtained for the off diagonal parameters using Algorithm 4.

4. Results

4.1. Deterministic Double Well Potential Model

The first conceptual climate model we consider is a cubic model coupled to the chaotic Lorenz system (28). It is fully deterministic and consists of a slow variable representing a climate process and three fast variables representing chaotic weather fluctuations. The slow variable moves inside a double well potential and is perturbed by the chaotic Lorenz system, which acts effectively as noise when $\epsilon \to 0$. The equations are as follows

$$\frac{dx}{dt} = x - x^3 + \frac{4}{90\epsilon}y_2$$  
$$\frac{dy_1}{dt} = \frac{10}{\epsilon^2}(y_2 - y_1)$$  
$$\frac{dy_2}{dt} = \frac{1}{\epsilon^2}(28y_1 - y_2 - y_1y_3)$$  
$$\frac{dy_3}{dt} = \frac{1}{\epsilon^2}(y_1y_2 - \frac{8}{3}y_3)$$

Sample paths are displayed in Fig. 5. We now fit a one dimensional cubic SDE to the data. We just consider the general cubic form (25)

$$dX_t = (a_1 + a_2X_t + a_3X_t^2 + a_4X_t^3)dt + \sigma dW_t$$

and estimate all of the parameters $\{a_1, a_2, a_3, a_4, \sigma\}$ from sparse observations of the system: again using $\Delta = 10.0$ and $N = 1000$. To update the drift parameters we use the Gibbs sampler of Section 3.2.1 The estimated posterior distributions are shown in Fig. 6. A lot of imputed data is needed before the estimates start to converge towards the values predicted by homogenization but the inference demonstrates that there is enough information in the sparse data set if the likelihood is well approximated.

Figure 7 shows the predictive skill of the one dimensional reduced model for $\sigma$ estimated for various $m$. Figures 7a and 7b show that the reduced model can reproduce the double well distribution of the full model although the separation of each well is underestimated for $m = 2$ and $m = 4$ due to the larger noise. For $m \geq 8$ the model reproduces well the full models marginal distribution for $x$. It is not clear whether there is much difference between $m = 8$ and $m = 64$. However, observing Figures 7c and 7d we see that the autocorrelation function for the full model is much better approximated when $m = 64$. This shows that the ability of our framework to impute data is a powerful way of deriving accurate reduced order models.
Figure 4: Estimate posterior distributions for parameters from a two dimensional model of the form Eq. 3 with \( N = 100 \) and \( \sigma = 0.1 \). The parameters, which are randomly generated, are written in the matrix notation introduced in Section 3.2.1. The histograms are the posterior distributions with uninformative prior, in red are the posterior distributions for parameters with stable SDEs and in black are the posterior distributions which include the stability matrix prior information derived in this chapter.
4.2. Model Reduction for Triad Systems

Now we apply our model fitting procedure to a triad model with a high dimensional deterministic system with two slow, climate variables coupled to fast chaotic dynamics. The reduction strategy has two challenges: to successfully approximate the deterministic variables by a stochastic process and to be insensitive to a lack of time scale separation.

The full system is given by

\[
\begin{align*}
\frac{dx_1}{dt} &= \frac{b_1}{\epsilon} x_2 y_1 \\
\frac{dx_2}{dt} &= \frac{b_2}{\epsilon} x_1 y_1 \\
\frac{dy_k}{dt} &= \frac{b_3}{\epsilon} x_1 x_2 \delta_{1,k} - \text{Re} \frac{ik}{2\epsilon^2} \sum_{p+q+k=0} \hat{u}_p^* \hat{u}_q^* \\
\frac{dz_k}{dt} &= -\text{Im} \frac{ik}{2\epsilon^2} \sum_{p+q+k=0} \hat{u}_p^* \hat{u}_q^*
\end{align*}
\]

where unknown parameters \( \sigma \) and \( \gamma \) have been introduced. Here we estimate them using the Algorithms 1 and 2 from observations of the climate variables alone. For convenience we con-
Figure 6: Posterior distribution estimates from MCMC output applied to a sparse data set (Δ = 10). Different distributions correspond to increasing amounts of missing data. The distribution in brown, for m = 64, agrees with the theoretical values predicted by the homogenization procedure.

Consider the inference problem where the model Eq. is driven by two independent Brownian motions.
Figure 7: Predictive statistics for the reduced double well model coupled to chaotic Lorenz system: Eq. (31d) for two values of $\epsilon$. In each plot the lines correspond to the inferred one dimensional model for different $m$.

We apply the inference to a data set with total time $T = 500$ and observation interval $\Delta = 0.1$. We simulate the system for $\epsilon = \{0.01, 0.1, 0.25, 0.5, 0.8, 1.0\}$. Posterior estimates for the case $\epsilon = 0.8$ are shown in Fig. 9. This value corresponds to a moderately small, though realistic (14), amount of time scale separation. The vertical lines are the mean values of the posterior distributions for the case $\epsilon = 0.01$. Note that, as expected, there is a small discrepancy between the estimates for $\epsilon = 0.8$ and $\epsilon = 0.1$.

For the parameter estimation we assume the most general model (Eq. 3) and, thus, we have to estimate 28 unknown parameters. We estimated these using Algorithms 1 and 2 as well as the
Gibbs sampling algorithm. We implemented the restricted parameter space using the constraints by using the element by element method. Fig. 10 shows the results for the drift parameters estimated for the case $\epsilon = 0.8$. As can be seen, for increasing the number of imputed data the posterior distributions tend to converge for most parameters.

We use the estimates from $m = 8$ to form a predictive model. Fig. 11 shows the autocorrelation function of mode $x_1$ of the full model with the one computed from the predictive model. As it can be seen, the decay of the autocorrelation function is very accurately captured. This is a particular hard test case because the time scale separation parameter $\epsilon = 0.8$ and we have only a very moderate, though realistic for the climate system (14), amount of time scale separation.
5. Summary

Here we developed a systematic Bayesian framework for the inference of the parameter of SDEs constrained by the physics of the underlying system. The physical constraints not only constrain the parameter space but also enforce global stability of the reduced order models. For conceptual climate models we use a constraint based on energy conservation which ensures global stability of the effective SDE. This constraint takes the form of a negative definite matrix. We then develop a new algorithm for the sampling of negative definite matrices. We
Figure 11: Autocorrelation plots of the full (vertical bars) and reduced (red) models for $\epsilon = 0.8$.

also develop a new algorithm for imputing data and show that imputing data improves the accuracy considerably of the parameter estimation. We demonstrated its power successfully on two conceptual climate models.

While we focused on climate models in this study our method is general enough that it can also be applied to other areas of fluid dynamics. Furthermore, also many other physical systems observe conservation laws and, thus, stability conditions can be derived which will be useful for parameter estimation procedures.

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References


