

**Accurate
measurements of CO
in humid air using the
CRDS technique**

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**Accurate measurements of carbon
monoxide in humid air using the cavity
ring-down spectroscopy (CRDS)
technique**

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Abstract

Accurate measurements of carbon monoxide (CO) in humid air have been made using the cavity ring-down spectroscopy (CRDS) technique. The measurements of CO mole fractions are determined from the strength of its spectral absorption in the near infrared region ($\sim 1.57 \mu\text{m}$) after removing interferences from adjacent carbon dioxide (CO_2) and water vapor (H_2O) absorption lines. Water correction functions that account for the dilution and pressure-broadening effects as well as absorption line interferences from adjacent CO_2 and H_2O lines have been derived for CO_2 mole fractions between 360–390 ppm. The line interference corrections are independent of CO mole fractions. The dependence of the line interference correction on CO_2 abundance is estimated to be approximately $-0.3 \text{ ppb}/100 \text{ ppm CO}_2$ for dry mole fractions of CO. Comparisons of water correction functions from different analyzers of the same type show significant differences, making it necessary to perform instrument-specific water tests for each individual analyzer. The CRDS analyzer was flown on an aircraft in Alaska from April to November in 2011, and the accuracy of the CO measurements by the CRDS analyzer has been validated against discrete NOAA/ESRL flask sample measurements made on board the same aircraft, with a mean difference between integrated in situ and flask measurements of -0.6 ppb and a standard deviation of 2.8 ppb . Preliminary testing of CRDS instrumentation that employs new spectroscopic analysis (available since the beginning of 2012) indicates a smaller water vapor dependence than the models discussed here, but more work is necessary to fully validate the performance. The CRDS technique provides an accurate and low-maintenance method of monitoring the atmospheric dry mole fractions of CO in humid air streams.

1 Introduction

Atmospheric carbon monoxide (CO) plays an important role in tropospheric chemistry (Logan et al., 1981), and acts as a useful tracer for emissions of CO_2 , CH_4 , and H_2 from

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biomass and fossil fuel burning (Andreae and Merlet, 2001; Levin and Karstens, 2007; Ehhalt and Rohrer, 2009; Turnbull et al., 2009). A considerable number of techniques have been employed to perform atmospheric measurements of CO, such as nondispersive infrared spectroscopy (NDIR) (Dickerson and Delany, 1988), vacuum ultra-violet resonance fluorescence (VURF) (Gerbig et al., 1999), tunable diode laser absorption spectroscopy (TDLAS) (Sachse et al., 1987), closed path Fourier Transform Infrared (FTIR) absorption (Griffith et al., 2012), gas chromatography combined with a mercuric oxide detector or a flame ionization detector (GC/HgO or GC/FID) (Novelli, 1999), and more recently, quantum cascade laser (QCL) (McManus et al., 2008), integrated cavity output spectroscopy (ICOS) (O’Keefe, 1998), and cavity ring-down spectroscopy (CRDS) (Crosson, 2008).

Comprehensive comparisons of measurements made using these techniques are given by Zellweger et al. (2009, 2012). Because variations of water vapor in the atmosphere modify the mole fractions of CO (known as the dilution effect), sample air has always been dried in previous techniques to obtain dry air mole fractions. A major advantage of some recently available techniques (QCL/ICOS/CRDS) is that water vapor mole fractions are simultaneously measured. This makes measurements of CO in humid air feasible as the H₂O values can be used to correct for dilution and potential spectroscopic effects. Zellweger et al. (2012) have derived water corrections for the QCL and ICOS analyzers, but not for the CRDS analyzer due to the fact that the more complicated nature of the spectral interference requires more complicated interference testing and analysis procedures. Carbon monoxide is detected by QCL or ICOS in the mid-infrared region where the CO absorption is well separated from other absorption features; in contrast, the CRDS technique measures CO in the near infrared region (~ 1.57 μm) where there are significant interferences from adjacent absorption lines of CO₂ and H₂O at ambient mole fractions of CO, CO₂, and H₂O. An example of the spectral absorptions for typical ambient mole fractions of CO (238 ppb), CO₂ (386.2 ppm), and H₂O (0.578 %) is given in Fig. 1a. In the scanned spectral region of the CO absorption line, there are four CO₂ absorption lines and three H₂O absorption

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lines. The ratios of the largest line for each species in the scanned spectral region are 1 : 23 : 22 for CO : CO₂ : H₂O. For each raw spectrum (collected every few seconds), the absorption lines of CO, CO₂, and H₂O are fit using a Levenberg-Marquardt optimization algorithm with a Galatry profile for CO (Varghese and Hanson, 1984) and a composite, empirically-derived spectral profile for CO₂ and H₂O. The mole fraction of CO is determined from the peak height of the CO spectral component (see Fig. 1b). Although it is challenging to get precise measurements of CO in the near infrared region using the CRDS technique, a big advantage is that this technique employs reliable, cheap, and stable infrared lasers rather than expensive quantum cascade lasers, and the same apparatus can also be employed to measure both CO₂ and CH₄ precisely. In fact, CRDS analyzers that measure CO₂/CH₄/CO/H₂O or CO₂/CO/H₂O (Picarro, Inc., Santa Clara, CA, USA, the G2401 or G2302 series) in one unit are commercially available. Extensive efforts have been carried out to make accurate measurements of CO₂ and CH₄ possible without drying the sample air (Chen et al., 2010; Winderlich et al., 2010; Nara et al., 2012; Rella et al., 2012). It is therefore beneficial to be able to accurately measure CO in humid air with the same analyzer as well.

In this paper, we derive water correction functions for the CO measurements by the CRDS technique to account for the dilution and spectroscopic effects. In Sect. 2, we present the results from a series of laboratory experiments that are then used to derive the water correction functions. The comparison of water correction functions for CO from different CRDS analyzers is given in Sect. 3. The field comparison between in situ and flask measurements on board an aircraft is shown in Sect. 4. In the final section, we present the discussion and conclusion.

2 Water correction functions for CO measurements

Mole fractions of CO and other trace gases are influenced significantly by variations of water vapor in the atmosphere. For this reason, atmospheric CO measurements need to be reported as mole fractions in dry air. In principle, there is a simple linear

relationship between the CO mole fraction in wet air (CO_{wet}) and in dry air (CO_{dry}):

$$\text{CO}_{\text{wet}}/\text{CO}_{\text{dry}} = 1 - \text{H}_2\text{O} \quad (1)$$

where H_2O is the water vapor mole fraction.

In the case of CO measurements by the CRDS technique, there are a number of factors that complicate the correlation between CO_{wet} and CO_{dry} : (1) the slope is not equal to -1 if water vapor is not accurately determined; (2) a nonlinear term exists as a result of pressure broadening (the mole fraction of CO is determined from the peak height of the absorption profile instead of the integrated spectral absorption area which would not be sensitive to pressure broadening); (3) residual line interference remains due to imperfect quantifications of adjacent absorption lines of CO_2 and H_2O . These effects can be compensated for by the following equation:

$$(\text{CO}_{\text{wet}} - \Delta\text{CO})/\text{CO}_{\text{dry}} = \left(1 + a' \times \text{H}_2\text{O}_{\text{pct}} + b' \times \text{H}_2\text{O}_{\text{pct}}^2\right) \quad (2)$$

Where CO_{wet} is the mole fraction of CO in humid air. For comparability between different analyzers of the same type, we consistently calculate CO_{wet} in ppb from the reported raw analyzer signal “peak84_raw” or “P84_RAW” as “peak84_raw” $\times 0.427 \times 1000$ or “p84_RAW” $\times 0.427 \times 1000$ (0.427 is the default factory calibration factor). CO_{dry} is the mole fraction of CO in dry air; ΔCO is the line interference adjustment to account for the residuals in the line fits of adjacent absorption lines of CO_2 and H_2O , and is a function of CO_2 and H_2O mole fractions, not of CO mole fraction; a' and b' are coefficients in the quadratic function to correct for the dilution and pressure-broadening effects and have units of $\%^{-1}$; $\text{H}_2\text{O}_{\text{pct}}$ is determined from the detection of the absorption lines of H_2O that are adjacent to the absorption line of CO, and has a unit of $\%$. Since $\text{H}_2\text{O}_{\text{pct}}$ directly affects the CO absorption line, it is appropriate to use measurements of these water vapor lines to correct all water vapor effects for CO measurements.

The CRDS analyzers mentioned above use the same spectroscopic fitting algorithms to derive the mole fractions of CO and CO_2 , and therefore the dilution and pressure-broadening effects for CO are expected to be similar to those for CO_2 . The corrections

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for these effects are a function of water vapor mole fractions, and are proportional to the mole fractions of CO. On the other hand, the line interferences caused by the residuals in the quantifications of adjacent CO₂ and H₂O absorption lines are independent of the mole fractions of CO as they only affect the baseline of the CO absorption. Since the parameters associated with dilution and pressure broadening (a' , and b') and line interference (ΔCO) in Eq. (2) cannot be directly fit using measurements of CO_{wet} and known CO_{dry}, we solve for these two effects separately based on water test data from laboratory experiments.

The water test data presented in this paper were obtained based on two different methods described in Rella et al. (2012). Method#1 employs a dew point generator (LI-COR model 610) or a micromodule (<http://www.liqui-cel.com/product-information/micromodule.cfm>) to humidify a gas stream from a tank. The advantage of this method is that mole fractions of CO in humidified gas are determined when the system has reached some equilibrium state at each humidity level. Method#2 can be described as a water droplet test. In this method, a small amount (0.2–0.8 ml) of distilled (or deionized) water (can be slightly acidified) is injected into the inlet of the CRDS analyzer to humidify air from a tank. The water droplet added to the inlet line is held at the hydrophobic particulate filter of the analyzer, through which water can only pass in the form of water vapor. This method is easy to perform, and provides water vapor ranging from 0 to 6%, but the vapor amount changes rapidly. To obtain slowly changing water vapor mole fraction, a small amount of silica gel soaked with deionized or acidified water and housed inside a small flow through chamber has been used at NOAA/ESRL to humidify the gas stream. Water test data from a list of experiments are summarized in Table 1.

2.1 Corrections for dilution and pressure-broadening effects

At typical CO mole fractions found in the atmosphere (50–500 ppb), the effects of dilution and pressure broadening are of the same order of magnitude as the line interferences. Therefore, these two types of effects cannot be differentiated in one experiment

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with ambient levels of CO. To minimize the effect of line interferences from CO₂ and H₂O, a water test was performed at Picarro Inc. in August 2010 on a CO₂/CO/H₂O analyzer (S/N: CKADS017) using tank air with an extremely high mole fraction of CO (~ 10⁶ ppb) and no CO₂. The experiment was carried out under a fume hood, using the simple water droplet test (Method#2, with a heated filter) and water vapor mole fractions ranging from 0 to 6%. In this case, the line interferences are negligible when compared to the influences by dilution and pressure broadening. The experimental data and analysis results are shown in Fig. 2. Both dilution and pressure-broadening effects can be accounted for using the quadratic function in Eq. 2 ($\Delta CO \approx 0$), and the derived coefficients $a' = -1.24 \times 10^{-2} \%^{-1}$, and $b' = -6.0 \times 10^{-4} \%^{-2}$. These two coefficients are very close to those for CO₂ $a = -1.20 \times 10^{-2} \%^{-1}$, $b = -2.67 \times 10^{-4} \%^{-2}$ (Chen et al., 2010). The residuals of the quadratic fit applied to 200 ppb dry mole fraction of CO are much smaller than the WMO recommended accuracy of ± 2 ppb for atmospheric CO measurements, with a standard deviation of 0.06 ppb (Fig. 2d). Transferring the water correction coefficients for CO₂ from one CRDS analyzer to another causes uncertainty on the order of 0.1 ppm out of 400 ppm for CO₂ (Chen et al., 2010; Rella et al., 2012), which is at least one order of magnitude more accurate than needed for the transferability of CO water corrections for dilution and pressure broadening among different analyzers (with a goal of 2 ppb out of 200 ppb ambient mole fraction). Therefore, these coefficients can be applied to all CO measurements made by the same type of CRDS analyzer to account for dilution and pressure-broadening effects.

2.2 Corrections for line interferences

An additional correction to the derivation of CO_{dry} involves the line interferences caused by the residuals in the quantification of adjacent absorption lines of CO₂ and H₂O; these are functions of CO₂ and H₂O mole fractions. To simplify the problem, we develop a function to correct for the line interferences caused by the residuals in the quantification of adjacent absorption lines of H₂O only, and only estimate the influences from variations of CO₂ mole fractions. For operation over a wider range of CO₂

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mole fractions, a more complete analysis is required. Five water tests have been performed on one CO₂/CH₄/CO/H₂O analyzer (S/N CFKBDS2007) at the laboratory of NOAA ERSL (see Table 1): one was made in February 2011 using a micromodule (Method#1), and four were completed in March 2012 using water droplets with different mole fractions of CO₂ and CO (silica gel soaked with deionized water, Method#2).

We fit a fourth-order polynomial function (Eq. 3) to the data from three droplet tests with 360–390 ppm CO₂ in air using a weighted least squares method, and the weights in the curve fit were calculated based on the density of the measurements in the water vapor domain so that equal water vapor range contributed equally to the curve fit:

$$\Delta\text{CO} = A \times \text{H}_2\text{O}_{\text{pct}} + B \times \text{H}_2\text{O}_{\text{pct}}^2 + C \times \text{H}_2\text{O}_{\text{pct}}^3 + D \times \text{H}_2\text{O}_{\text{pct}}^4 \quad (3)$$

A , B , C , D are coefficients of the polynomial fit, and ΔCO is calculated as $\text{CO}_{\text{wet}} - \text{CO}_{\text{dry}} \times (1 + a' \times \text{H}_2\text{O} + b' \times \text{H}_2\text{O}^2)$, with $a' = -1.24 \times 10^{-2} \%^{-1}$, and $b' = -6.0 \times 10^{-4} \%^{-2}$ from Sect. 2.1.

The fourth-order function was empirically chosen. We tried to fit these data with a series of polynomial functions from the second order to the fifth order, and found the residuals of those fits below the fourth-order exhibited a clear dependence on water vapor (now shown). Besides this, the standard deviations of the residuals of the fits were 3.3 ppb, 2.3 ppb, 1.1 ppb, and 1.1 ppb, respectively. It is important to note that the coefficients A – D represent the water vapor bias in the CO measurement after the basic spectral model for water vapor has been removed from the raw spectral data. In other words, if the spectral model describing this spectral region was perfect, then these coefficients would be zero. The experimentally derived ΔCO and the fourth-order polynomial fit are given in Fig. 3. The February 2011 test was not fit due to the lack of data below 1 % H₂O.

The derived coefficients of the weighted fourth-order polynomial fit are $A = 16.628 \%^{-1}$, $B = 12.460 \%^{-2}$, $C = -10.692 \%^{-3}$, $D = 1.275 \%^{-4}$. The ΔCO values from the experiment made in February 2011 are plotted on the same figure. The 0.5 % bin averaged residuals are mostly within ± 2 ppb, except for two points that are slightly

greater than 2 ppb for the four tests with CO₂ ranging from 360–390 ppm. There is no evidence of drift in the water correction functions from February 2011 to March 2012.

The tests cover CO mole fractions ranging from 131 ppb to 303 ppb, CO₂ mole fractions ranging from 360 ppm to 390 ppm, and H₂O mole fractions ranging from 0 to 4%.

The fact that all these data can be fit to one single function with small residuals demonstrates that this correction applies to different CO mole fractions for the range of CO₂ from 360 ppm to 390 ppm. An extensive characterization of the dependence of CO water corrections on CO₂ abundance requires determination of a liner term in CO₂ and a term that is linear in the product of CO₂ and H₂O. We do not have enough test data to quantify these two terms from the water tests. However, we made a water test using 550 ppm CO₂ for a different CRDS analyzer and did not observe significant differences (see Sect. 2.3, Fig. 4). The dependence of CO measurements in dry air on CO₂ is presented in Sect. 2.4. Note that the size of the third and fourth order terms are large at high water concentrations, and the correction factors can only be used for the tested water vapor range of 0–4%. Extrapolating the polynomial curve to higher water vapor values will cause a large uncertainty.

2.3 Corrections for line interferences for analyzers with a new fitting algorithm

Since the beginning of 2012, a new line fitting algorithm has been applied by Picarro Inc. to the CO analyzers, which reduces the line interferences from water vapor on the CO measurements. We performed two water droplet tests (Method #2, 380 ppm and 550 ppm CO₂) at NOAA/ESRL on a CO₂/CH₄/CO/H₂O analyzer (CFKBDS2059) that uses the new fitting algorithm. The same fourth-order polynomial function was used to fit the data from the water test (see Fig. 4), and the derived coefficients are: $A = 2.032\%^{-1}$, $B = 1.821\%^{-2}$, $C = -0.671\%^{-3}$, $D = 0.048\%^{-4}$. Compared to the corrections shown in Fig. 3, the magnitude of the correction has been significantly reduced, especially in the higher order terms. Note that no significant differences have been observed for the test data using 550 ppm CO₂.

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2.4 Dependence of CO measurements on CO₂ mole fractions

A dependence of CO measurements on CO₂ mole fractions is expected due to line interferences. To characterize the dependence, one experiment was performed on one CO₂/CH₄/CO/H₂O analyzer (S/N CFKBDS2007) at Picarro Inc. in January 2011, and the apparatus is shown in Fig. 5.

To start with, air from a low CO₂ tank was provided to the CRDS analyzer after passing through a CO scrubber (Sofnocat 514), and then switched to air from a high CO₂ tank. The Sofnocat responds very slowly to changes in CO₂ (it was adsorbed on the sofno-cat substrate) so that a slow change in CO₂ was obtained (Fig. 6a). The water level was about 0.4 %, and was stable throughout the measurement (Fig. 6b). The reading of raw CO was not close to zero due to the presence of water vapor (Fig. 6c). Water corrections were then applied based on Eq. (5) and the coefficients in Sects. 2.1 and 2.2 to derive dry CO, which was indeed around zero (Fig. 6d). From a linear fit of dry CO to CO₂, we derive a dependence of CO measurements of -0.3 ppb/100 ppm CO₂, and the 100 ppm bin averaged residuals of the linear fit are smaller than ± 1 ppb (Fig. 6e, f).

Alternatively, the CO₂ dependence can be derived from measurements of a series of CO standards with different CO₂ mole fractions by fitting the following equation:

$$\text{CO}_{\text{tank}} = A' + B' \times \text{CO}_{\text{raw}} + C' \times \text{CO}_2 \quad (4)$$

Here CO<sub>tan-
k</sub> is the known CO mole fraction, CO_{raw} is the measured mole fraction by a CRDS analyzer (“peak84_raw” $\times 0.427 \times 1000$), and CO₂ is the CO₂ mole fraction. To be able to determine the parameters A' , B' , C' accurately, a large span of CO₂ is required (our calibration data are not shown due to a narrow span of CO₂).

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2.5 Application of corrections for CO measurements

Combining the corrections for dilution and pressure broadening and for line interferences, Eqs. (2)–(3), the corrected CO values can be expressed as

$$\text{CO}_{\text{corrected}} = \frac{\text{CO}_{\text{wet}} - (A \times \text{H}_2\text{O}_{\text{pct}} + B \times \text{H}_2\text{O}_{\text{pct}}^2 + C \times \text{H}_2\text{O}_{\text{pct}}^3 + D \times \text{H}_2\text{O}_{\text{pct}}^4)}{1 + a' \times \text{H}_2\text{O}_{\text{pct}} + b' \times \text{H}_2\text{O}_{\text{pct}}^2} \quad (5)$$

- 5 The coefficients $a' = -1.24 \times 10^{-2} \%^{-1}$, and $b' = -6.0 \times 10^{-4} \%^{-2}$ can be used for different analyzers, whereas coefficients A, B, C, D need to be determined from instrument-specific water tests (see Sect. 3). Note that we do not include any CO_2 dependence in the correction function because it is determined to be small for CO_2 mole fractions ranging from 360–550 ppm.

3 Transferability of water corrections across different analyzers

When using the water corrections given in Eq. (4), it is important to understand how consistent they are from one instrument to the other. In total, water tests have been performed on five analyzers with the old fitting algorithm in groups from NOAA/ESRL (CFKBDS2008 and CKADS2027), MPI-Jena (CFKBDS2003 and CKADS2004), and MPI-Mainz (CKADS018) using Method#1 and/or Method#2 (see Table 1). The differences between the corrected CO values based on the coefficients derived in Sects. 2.1 and 2.2 and the measured dry mole fraction of CO for each test and analyzer are shown in Fig. 7. Note that several test results may have been averaged for each individual analyzer. The results clearly show that each analyzer requires instrument-specific water corrections (i.e. coefficients A, B, C, D) to meet the WMO recommended accuracy of ± 2 ppb. The variations in the CO correction function between analyzers captures the small differences in the spectrograms delivered by each instrument, due to tiny differences between the reported wavelengths, absorbances, temperature, or pressure.

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4 Field validation

The CRDS analyzer tested in Sect. 2 (CFKBDS2007) was flown without a drying system on board an aircraft in Alaska from April to November 2011, and from April 2012 until the present. A detailed description of the field deployment is given by Karion et al. (2012). A total of 12 flights have been made in 2011, and each flight took ~ 8 h, during which 24 flasks were sampled. In-flight calibrations were made by providing air from three tanks to the CRDS analyzer every 30 min. The flasks were analyzed at NOAA/ESRL, with an accuracy of 2 ppb (Novelli et al., 2003). The in situ CO measurements were integrated using weighting functions that accounted for observed atmospheric variability during the flask sampling period (Chen et al., 2012). To accurately represent the weighting functions, knowing the times when the pressurizing process started and ended is critical. Since only the times when the pressurizing process ended were recorded, the times when the pressurizing process started were estimated based on the ambient pressure and the time it takes to fill a flask on the ground. The comparison of integrated in situ measurements with flask measurements of CO during all flights shows a mean bias of -0.6 ppb and a standard deviation of the differences of 2.8 ppb, which is better than the WMO recommended Inter-Laboratory comparability or accuracy of ± 2 ppb. Further, no significant dependence of the differences between in situ and flask measurements on any of the variables (day, $\text{H}_2\text{O}_{\text{pct}}$, CO_2 , CO, CH_4 , and ambient pressure) has been found, partially due to the relative large scatter in the differences between in situ and flask measurements (Fig. 8).

5 Discussion and conclusion

Accurate measurements of carbon monoxide (CO) in humid air have been made using the cavity ring-down spectroscopy (CRDS) technique. The measurements of CO mole fractions are determined from the strength of its spectral absorption in the near infrared region ($\sim 1.57 \mu\text{m}$) after removing interferences from adjacent carbon dioxide

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(CO₂) and water vapor (H₂O) absorption lines. To measure humid ambient air without drying the air, laboratory experiments have been carried out to derive the water correction functions for CO. These water correction functions account for the dilution and pressure-broadening effects as well as line interferences from adjacent CO₂ and H₂O absorption lines for CO₂ mole fractions ranging from 360–390 ppm, and are independent of CO mole fractions. We determined a dependence of CO measurements on CO₂ mole fractions of –0.3 ppb/100 ppm CO₂. This result quantifies the potential errors of CO measurements under varying CO₂ mole fractions. For measurements of CO on aircraft or at a background observational site, the dependence seems to be rather small, since variations of CO₂ larger than 100 ppm CO₂ are rarely seen (corresponding to a bias of CO ±0.3 ppb). For the cases when significant CO₂ variations are possible, e.g. exhaust measurements, soil chamber experiments, the dependence of CO measurements on mole fractions of CO₂ needs to be better characterized.

Unlike the good transferability of the water correction coefficients for CO₂ and CH₄ among different CRDS analyzers (Rella et al., 2012), comparisons of water correction functions for CO from five different analyzers of the same type show significant differences, making it necessary to perform instrument-specific water tests for each individual analyzer. The magnitude of the water corrections has been significantly reduced by employing a new fitting algorithm (in Picarro CRDS CO analyzers newer than early 2012). Further tests need to be done to determine if the transferability of the water corrections for analyzers using the new fitting algorithm have been improved. From our limited experience, significant drift has not been observed in the water corrections for one CRDS analyzer (CFKBDS2007) over a period of one year. Based on these laboratory studies showing the sensitivity of the water correction to CO₂ mole fraction and the analyzer to analyzer variability it is recommended that fits be made to each analyzer periodically (3–6 months) with standard gas concentrations that closely resemble both the span of CO₂ and CO that will be observed.

The CRDS analyzer was flown on an aircraft in Alaska from April to October in 2011, and the accuracy of the CO measurements by the CRDS analyzer has been validated

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against discrete NOAA/ESRL flask sample measurements made on board the same aircraft, with a mean difference between integrated in situ and flask measurements of -0.6 ppb and a standard deviation of 2.8 ppb. No significant dependence of the differences between in situ and flask measurements could be detected, partly because the noise in the differences is relative large. Direct comparisons of in situ measurements of CO by two or more side-by-side analyzers, as already investigated by Zellweger et al. (2012), could potentially discover or exclude issues that influence the accuracy of the CO measurements by the CRDS technique.

Although the precision of the CO measurements by the CRDS technique on a short-time period is not as high as other available techniques, such as QCL and ICOS, it is possible to make accurate measurements of CO using this technique without drying the air if measurements over a sufficiently long period can be integrated and calibrated against standards. The CRDS analyzer does not need frequent calibration (Winderlich et al., 2010; Karion et al., 2012; Richardson et al., 2012) or a drying system, and could be an accurate and low-maintenance instrument useful for monitoring the atmospheric mole fractions of CO.

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Table 1. Summary of water tests for CO on a series of CRDS analyzers.

Group	Analyzer S/N	CO (ppb)	CO ₂ (ppm)	Method
NOAA/ESRL	CFKBDS2007	131–303	360–390	#1 & #2
	CFKBDS2008	301	385	#2
	CKADS2027	177	450	#1
	CFKBDS2059	129	380–550	#2
Picarro, Inc.	CKADS017	~ 10 ⁶	0	#2
MPI-Mainz	CKADS018	81–492	365–470	#2
MPI-Jena	CFKBDS2003	213	395	#1 & #2
	CFKBDS2004	188–442	390–440	#1

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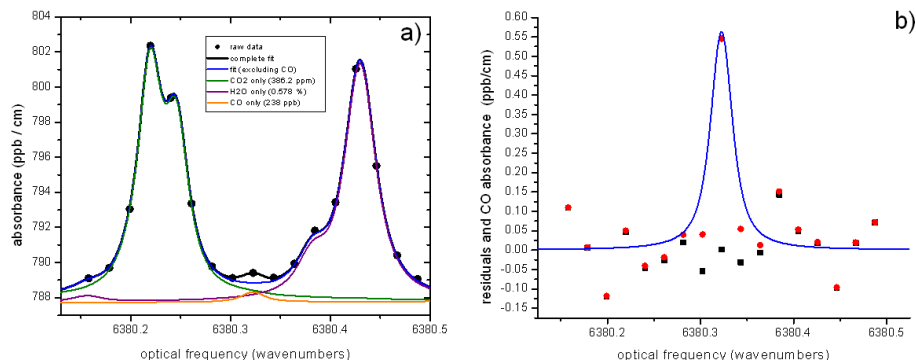


Fig. 1. The spectroscopic data and fits of CO and adjacent absorptions of CO₂ and H₂O: **(a)** the spectral fits of absorption data (black solid dots) for CO₂ (four overlapping lines) only, H₂O (three lines) only, CO only, excluding CO, and all data; **(b)** the residuals of a multi-component fit to the data in the left panel, with CO included in the fit (black solid squares), and with CO not included (red solid dots), along with the CO spectral fit (blue curve).

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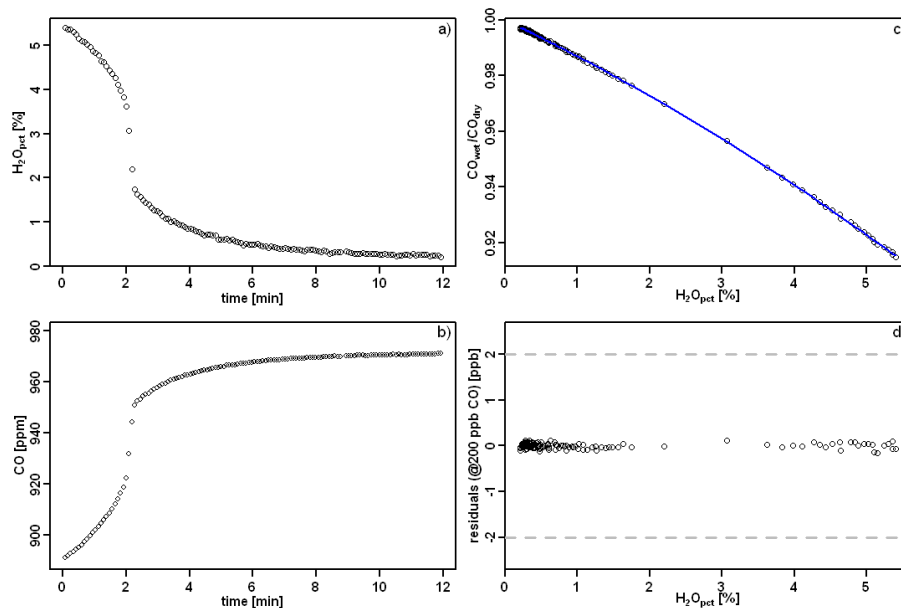


Fig. 2. Experimental data and analysis results during a water test using air with a high mole fraction of CO ~ 1000 ppm: **(a, b)** time series of H_2O_{pct} and mole fractions of measured CO during the test, respectively; x-axis is given in minute relative to the start of the experiment. **(c)** plots of CO_{wet}/CO_{dry} vs. H_2O_{pct} and a quadratic fit (curve in blue) **(d)** The black circles are residuals of the curve fit in **(c)** when applied to 200 ppb dry mole fraction of CO, with a standard deviation of 0.06 ppb; dashed lines indicate the comparability goals set by the WMO for CO measurements.

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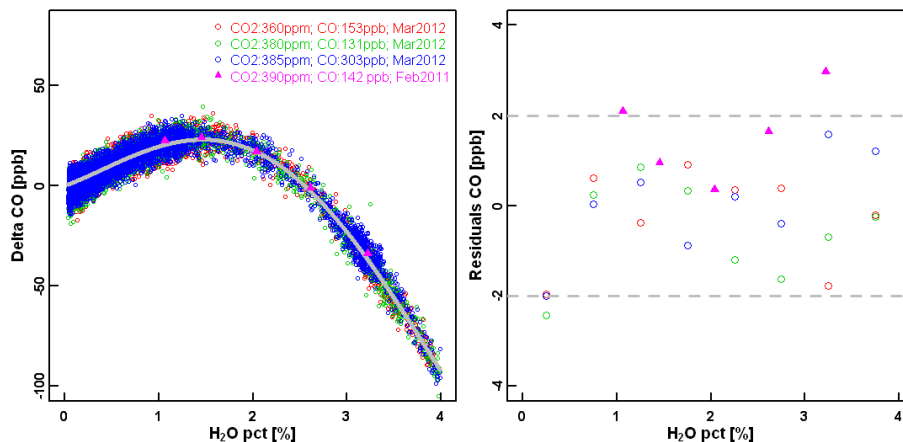


Fig. 3. Fourth-order fit of water droplet test data with 360–390 ppm CO₂ (red, green, and blue), Δ CO vs. H₂O mole fractions for the range of 0–4%. The experimental data were weighted based on the density of the measurements in the water vapor domain so that equal water vapor range contributed equally to the curve fit. The pink solid triangles are averaged values of Δ CO obtained using water test Method#1 in February 2011, and cover a water vapor range of 1–3.2%. The residuals from the fit for each water test, averaged over 0.5% H₂O intervals, are shown on the right figure.

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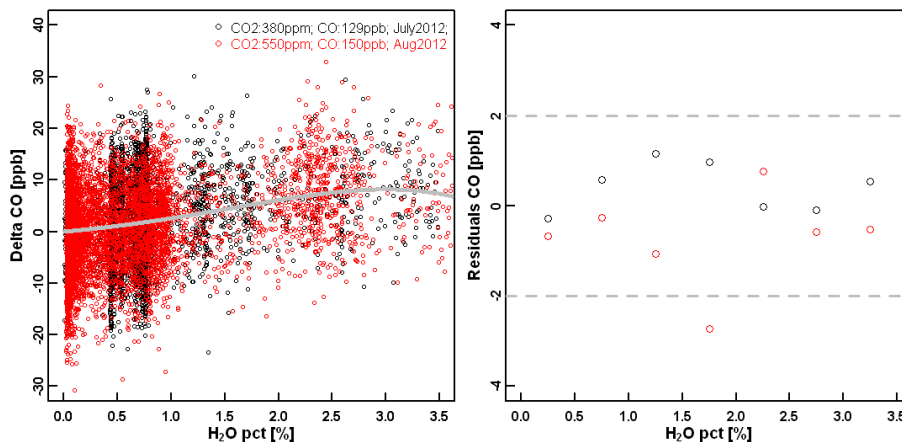


Fig. 4. Fourth-order fit of water droplet test data for analyzer with new fitting algorithm, ΔCO vs. H_2O mole fractions for the range of 0–3.5%. The experimental data were weighted in the same way as for those presented in Fig. 3. The residuals from the fit are shown on the right figure, and dashed lines indicate the comparability goals set by the WMO for CO measurements.

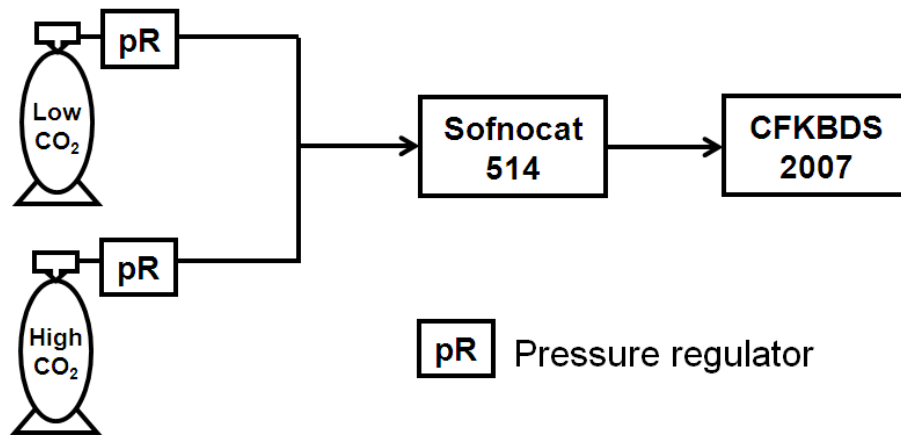


Fig. 5. Schematics of the CO₂ dependence test.

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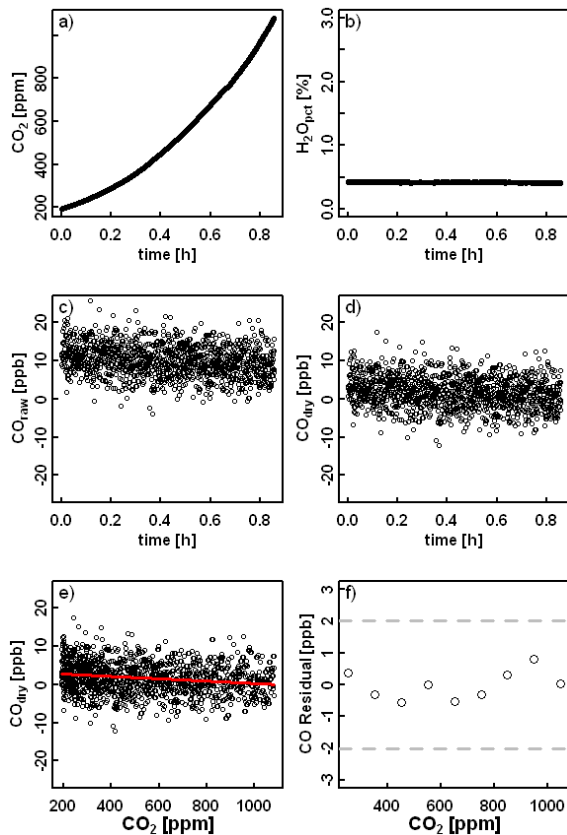


Fig. 6. Experimental data of a CO_2 dependence test: **(a)** the slow change in CO_2 over time; **(b)** the time series of water vapor; **(c)** the time series of raw CO; **(d)** the time series of water corrected dry CO; **(e)** dry CO vs. CO_2 , and a linear fit shown by the red line; **(f)** the 100 ppm bin averaged residuals of the linear fit. Dashed lines indicate the comparability goals set by the WMO for CO measurements.

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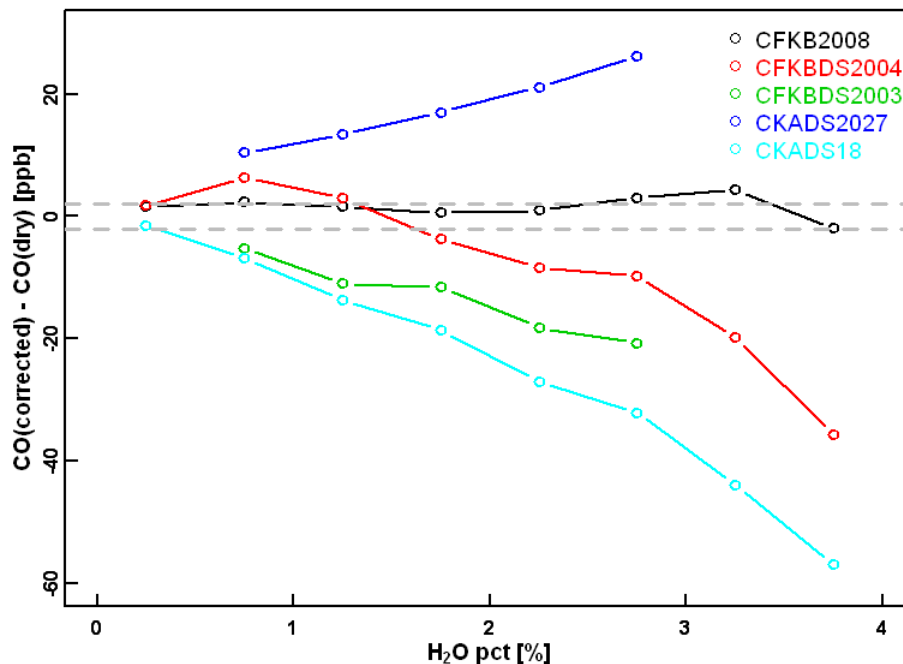


Fig. 7. Differences of corrected wet measurements using the coefficients ($A = 16.628\%^{-1}$, $B = 12.460\%^{-2}$, $C = -10.692\%^{-3}$, $D = 1.275\%^{-4}$) derived in Sect. 2 for CFKBDS2007 and dry values for water tests on each analyzer. Dashed lines indicate the comparability goals set by the WMO for CO measurements.

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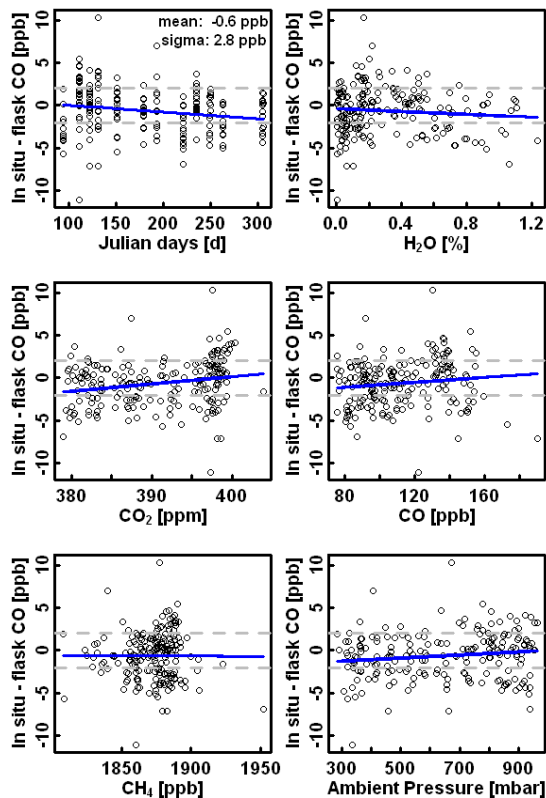


Fig. 8. Comparisons of in situ and flask measurements of CO made by NOAA/ESRL on board an aircraft in Alaska in 2011. **(a)** The time series of the differences between integrated in situ measurements and flask measurements; **(b–f)** the differences between in situ measurements and flask measurements plotted as a function of H_2O , CO_2 , CO, CH_4 , and ambient pressure, with linear fits shown in blue lines. The WMO recommended accuracy of ± 2 ppb is shown on each plot in grey dashed lines.

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