



## Comparisons of the Orbiting Carbon Observatory-2 (OCO-2) $X_{CO_2}$ measurements with TCCON

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**Abstract.** NASA's Orbiting Carbon Observatory-2 (OCO-2) has been measuring carbon dioxide column-averaged dry air mole fractions,  $X_{\text{CO}_2}$ , in the Earth's atmosphere for almost two years. In this paper, we describe the comparisons between the OCO-2 version 7Br retrievals and  $X_{\text{CO}_2}$  estimates from OCO-2's primary ground-based validation network: the Total Carbon Column Observing Network (TCCON). The OCO-2  $X_{\text{CO}_2}$  retrievals, after bias correction, agree well globally with the TCCON for nadir, glint, and target observations, with median differences less than 0.5 ppm and RMS differences typically below 1.5 ppm. Target observations over TCCON stations correlate best with the TCCON data ( $R^2 = 0.83$ ) on a global scale. At local scales, the target comparisons reveal residual biases likely related to surface properties and aerosol scattering. It is thus crucial to continue measurement comparisons with TCCON to monitor and evaluate the OCO-2  $X_{\text{CO}_2}$  data quality throughout its mission.

## 1 Introduction

NASA's Orbiting Carbon Observatory-2 (OCO-2) is NASA's first Earth-orbiting satellite dedicated to observing atmospheric carbon dioxide ( $\text{CO}_2$ ) to better understand the carbon cycle. The mission's main goal is to measure carbon dioxide with enough precision and accuracy to characterize its sources and sinks on regional scales and to quantify its seasonal and interannual variability (Crisp et al., 2008; Boland et al., 2009; Crisp, 2015). OCO-2 was successfully launched on July 2, 2014 into low-Earth orbit, and measures near infrared spectra of sunlight reflected off the Earth's surface. Carbon dioxide and oxygen ( $\text{O}_2$ ) in the Earth's atmosphere absorb sunlight at well-known wavelengths in the three spectral regions observable by OCO-2, and from those absorption features, atmospheric abundances of carbon dioxide and surface pressure, and other atmospheric and surface properties (e.g., cloud and aerosol optical depth and distribution, water vapor, temperature, and surface reflectance) are obtained (O'Dell et al., 2012; Connor et al., 2008).

The main product from the retrieved abundances of carbon dioxide and surface pressure is the column-averaged dry-air mole fraction of  $\text{CO}_2$ , called  $X_{\text{CO}_2}$ , which is a useful product for carbon cycle science, as it is directly related to surface fluxes of  $\text{CO}_2$  (Yang et al., 2007; Keppel-Aleks et al., 2011). OCO-2 measures  $X_{\text{CO}_2}$  with unprecedented precision from space (better than 1 ppm, Eldering et al., 2016) but possesses biases that the OCO-2 team have attempted to characterize and remove (Mandrake et al., 2015; O'Dell et al., 2016). To validate the OCO-2 measurements, we use the Total Carbon Column Observing Network (TCCON, Wunch et al., 2011a), a comprehensive ground-based validation network that also measures  $X_{\text{CO}_2}$ .

The OCO-2 satellite has three viewing modes: nadir mode, in which the instrument points straight down at the surface of the Earth, glint mode, in which the instrument points just off the glint spot on the surface, and target mode, in which the observatory is commanded to scan about a particular point on the ground as it passes overhead. All three modes must be independently verified using



comparisons with the ground-based TCCON data. This paper will describe the OCO-2 observation modes in §2, how the OCO-2 version 7 algorithm target-mode retrievals compare with the TCCON data in §3, and how the glint and nadir mode measurements compare with TCCON in §4.

## 2 OCO-2 Observation Modes

40 OCO-2's nadir and glint measurements are considered the nominal “science modes” of the OCO-2 measurement scheme. The nadir observations produce useful measurements only over land and near the sub-solar point over tropical oceans. The glint data are often separated into glint over land (“land glint”) and glint over water (“ocean glint”), as the two modes use different surface reflectance models: Lambertian over land, and Cox-Munk over water (O'Dell et al., 2016). Retrievals are performed  
45 over a limited latitude range in glint due to concerns about biases introduced by aerosol scattering over the largest optical path lengths; see Fig. 1. The nadir mode data can provide more reliable  $X_{\text{CO}_2}$  measurements over higher latitudes over land, which is particularly important in the northern hemisphere, where the boreal forest, the driver of the  $\text{CO}_2$  seasonal cycle, extends north of  $70^\circ\text{N}$ .

OCO-2 has a geographical “near-repeat” after 16 days or 233 orbits. The original measurement  
50 scheme alternated between glint and nadir observations on alternate 16-day ground track repeat cycles. This is not ideal, due to the loss of ocean measurements during nadir mode, and the loss of high latitude measurements during glint mode. Thus key components of the carbon cycle (e.g., the springtime draw down of  $\text{CO}_2$  due to the onset of the northern hemisphere growing season) would be poorly sampled. The observing strategy was optimized to improve the coverage of the oceans  
55 and high latitude land masses on July 2, 2015 to alternate between glint and nadir modes for each subsequent orbit. The OCO-2 observation scheme was further optimized on November 12, 2015, to assign orbits that are almost entirely over ocean to be always measured in glint mode. This change was made on 72 out of the 233 orbital paths: 15 over the Atlantic and 57 over the Pacific, resulting in higher data throughput because there are now fewer nadir soundings over ocean. Crisp et al. (2016)  
60 discuss the measurement strategy in detail.

Target mode was designed to assist the OCO-2 science team in evaluating the biases in the OCO-2  $X_{\text{CO}_2}$  product. The target locations are mostly selected to be coincident with ground validation stations, typically at TCCON sites. During a target-mode maneuver, the OCO-2 satellite rotates from its nominal science mode to point at a selected ground location. This transition takes approximately  
65 5 minutes and rotates the spacecraft's solar panels away from the sun. The spacecraft then scans across the site or “nods” as it passes overhead to sweep across the ground several times (see Fig. 2) over a period of about 4.5 minutes: these dithered measurements comprise the “target-mode data”. The spacecraft then transitions out of target mode and back into its nominal science mode over the next 5 minutes. In total, the maneuver takes about 14.5 minutes, and during this time, the spacecraft,  
70 traveling at  $7.5 \text{ km} \cdot \text{s}^{-1}$ , has traveled over 6,500 km.



The strength of target mode measurements is that thousands of spectra are obtained in a short period of time over a small region of the world (about  $0.2^\circ$  longitude  $\times$   $0.2^\circ$  latitude for the densest measurements, Fig. 2). As long as the target location is far from large emissions sources,  $X_{\text{CO}_2}$  can be assumed constant spatially and temporally within a target measurement. However, during the  
75 maneuver, many other parameters can change, such as the atmospheric path, the path length of the measurement (the “airmass”), surface reflectivity (albedo), and topography. Any variability in the retrieved  $X_{\text{CO}_2}$  in the target mode data is considered to be an artifact, and can provide insight into biases caused by the algorithm’s treatment of the parameters. With this in mind, the target locations were carefully chosen to span a wide range of latitudes, longitudes, and surface types to challenge  
80 the OCO-2 retrieval algorithm and reveal any biases caused by it.

## 2.1 Target Locations

There is a limited number of ground locations that can be targeted because the locations must be pre-programmed into the spacecraft software. For the first year after launch, there were 19 possible target locations. In July 2015, 8 additional targets slots became available, allowing for 27 target locations.  
85 At several times, target locations have been changed or replaced. A list of the ground target locations and dates is provided in Table 1, and a map of their locations is in Fig. 3. Individual locations can be targeted by OCO-2 only on specific OCO-2 orbit paths. Only one target location can be assigned to a given orbit path, and only if the OCO-2 ground track for that path is sufficiently close to the ground target location. Thus for each day, there are between one and seven ground target locations  
90 to choose from. The spacecraft power systems can handle up to three target-mode maneuvers per day due to the power constraints imposed by rotating the spacecraft solar panels away from the sun. We typically select only one target per day. On 40 occasions, we have targeted two locations on one day; on 3 occasions, we have targeted 3 locations on one day.

There are several TCCON stations that are located in regions with spatially varying topography or  
95 ground cover. For example, the Białystok TCCON station has a nearby forest, the Lauder TCCON station is in the midst of rolling hills, the Wollongong TCCON station is between the ocean and a sharp escarpment, the Darwin TCCON station is near the coast close to sea level, and the Edwards TCCON station is adjacent to a very bright playa, a land surface property previously identified in the Greenhouse Gases Observing Satellite (GOSAT, Kuze et al., 2009, 2016) project as challenging  
100 for  $X_{\text{CO}_2}$  retrievals (Wunch et al., 2011b). With target-mode measurements, the impact that local surface variability has on the  $X_{\text{CO}_2}$  retrievals becomes apparent.

Other TCCON stations (e.g., Park Falls, Lamont) have relatively uniform surface properties and are reasonably far from anthropogenic  $\text{CO}_2$  sources, but the ground cover can vary from season to season. Park Falls is snow-covered in winter, dark green in summer. Lamont is green in spring  
105 and brown in winter. The Sodankylä and Eureka sites, located at high northern latitudes, challenge the OCO-2 algorithm at very high solar zenith angles and airmasses, and with snowy scenes. Izaña,



Réunion and Ascension, all lower-latitude sites, are located on small islands remote from large land masses, but with significant topography. The Izaña TCCON station (28.3°N) is at 2.37 km altitude, whereas the Réunion (20.9°S, 0.087 km) and Ascension Island (7.9°S, 0.032 km) stations are closer  
110 to sea level.

Several TCCON target stations are near or in urban regions with varied topography and emissions sources: Pasadena is in the South Coast Air Basin of California, which contains the city of Los Angeles and is home to 17 million residents. Tsukuba is in a highly urbanized city near Tokyo (pop. ~228,000). The Paris TCCON station is located 4.5 km from the Eiffel Tower in the heart of the city  
115 (pop. ~2.24 million), and the Karlsruhe TCCON station is in a smaller, more isolated urban region (pop. ~300,000) surrounded by forest.

There are several target locations that are not TCCON stations (Fig. 3, orange stars), and although data from those targets will not be analysed in this paper, the data will help assess bias. Railroad Valley is a heavily instrumented radiometric calibration site (Kuze et al., 2011); Shanghai, São Paulo  
120 and Mexico City are geographically well-constrained urban regions with significant CO<sub>2</sub> emissions; Rosemount and Litchfield have instrumentation that will help verify the OCO-2 solar induced fluorescence observations; Boulder has frequent AirCore CO<sub>2</sub> profile measurements (Karion et al., 2010); Fairbanks is a future TCCON station. The Libya location has very specific and consistent surface properties and observations there will improve the radiometric calibration of OCO-2.

## 125 2.2 Target Selection

Target locations are selected a day or two in advance, based on the weather forecast, the operational status of the TCCON station (if the target is a TCCON station), the importance of the projected data loss in nadir or glint mode from performing the target-mode operation, and the historical statistics of successful target-mode measurements over that site. The projected data loss depends primarily  
130 on whether the nominal mode for that orbit was nadir over land, nadir over ocean, glint over land, or glint over ocean. If the nominal mode is nadir over ocean, little data loss occurs, as nadir measurements over ocean are usually too dark in the near infrared for successful retrievals: in this case, the target is almost always selected given a reasonable weather forecast. This has mostly been the case for Réunion Island, which has been targeted regularly from OCO-2 nadir orbits. For the other  
135 three cases, there will be some loss of regular science data to accommodate a target-mode operation. In these cases, the historical statistics of acquiring good target-mode data and weather forecasts are weighted more heavily before enabling the target. Often, if the weather forecast is not ideal, no target-mode measurements will be selected.

As of April 30, 2016, 227 targets have been selected, 195 of them over TCCON stations (Fig.  
140 4). Of the 195 targets of TCCON stations, about 90% were “successful”, in that both OCO-2 data and TCCON data were recorded. This 10% loss was largely due to poor weather forecasts, resulting in completely overcast skies, during which the TCCON station was unable to record any data.



There were two instances of unforeseen mechanical failures at TCCON stations. Of the 195 targets recorded, about 55% were clear enough to obtain sufficient high-quality (low warn level) OCO-2  
145 data to compare with TCCON data.

### 3 Target Mode Comparisons to TCCON and Residual Bias Assessment

All current space-based  $X_{\text{CO}_2}$  measurements have systematic biases caused by uncertainties in the spectroscopy, by limitations in the information content of the measurements and uncertainties or oversimplifications in the optical properties of the atmosphere and surface, and by uncertainties  
150 in the instrument characterization and calibration (e.g., Crisp et al., 2016; Wunch et al., 2011a; Guerlet et al., 2013; Schneising et al., 2012). Considerable effort is dedicated to creating robust “bias corrections”. There are three key types of biases considered here, and the bias correction addresses them all: the first is any constant scaling factor from the World Meteorological Organization (WMO) trace-gas standard scale (a “constant” bias), the second is spurious correlations of the re-  
155 trieved  $X_{\text{CO}_2}$  with other retrieval parameters such as the surface pressure retrieval error, signal level, airmass, surface albedo (a “parameter-dependent” bias), and the third is any footprint-dependent biases. Footprint-dependent biases are corrected with calibration observations compiled using a special subset of OCO-2 data collected while the spectrometer slit is oriented parallel to the ground track and the eight OCO-2 footprints are geographically coincident (Mandrake et al., 2015). The  
160 parameter-dependent biases must be removed before the constant scaling factor can be established, as the parameter-dependent bias correction can impact the magnitude of the constant scaling factor. The procedure by which the OCO-2 data have been bias-corrected is documented in detail by Mandrake et al. (2015) and O’Dell et al. (2016) and is briefly described here. The parameter-dependent bias correction uses a genetic algorithm to determine which retrieval parameters account for the  
165 largest fraction of the spurious variability found in the estimated  $X_{\text{CO}_2}$  on large spatial scales. The algorithm uses several subsets of the OCO-2 data for this task: a “southern hemisphere approximation” which exploits the low spatial and temporal variability of  $X_{\text{CO}_2}$  in the the southern hemisphere south of 25°S (e.g., Wunch et al., 2011b); a “small area analysis” which exploits the low spatial variability of  $X_{\text{CO}_2}$  within small regions (0.89° latitude on a single orbit track) and can be applied  
170 at all latitudes; and data near coastlines wherein differences in  $X_{\text{CO}_2}$  between land and ocean can be clearly diagnosed. A multivariate regression is performed between spurious  $X_{\text{CO}_2}$  variability and the parameters. The resulting slopes of the regressions allow us to then subtract the predicted bias from the  $X_{\text{CO}_2}$  values. In the results that follow, the parameter-dependent biases in the OCO-2 target-mode data have been removed following Mandrake et al. (2015), allowing us to determine the  
175 constant scaling factor.

Placing the OCO-2 data on the World Meteorological Organization’s (WMO) trace-gas standard scale is crucial to obtaining flux inversions consistent with the state-of-the-art inversions of surface



in situ CO<sub>2</sub> data, which are carefully calibrated to the WMO scale (Zhao and Tans, 2006). To achieve this, we compare the OCO-2 target mode data with the latest version of the TCCON data (GGG2014, Wunch et al., 2015), which have been tied to the WMO trace-gas standard scale through comparisons with in situ CO<sub>2</sub> profiles from aircraft (Wofsy, 2011; Pan et al., 2010; Singh et al., 2006) and AirCore (Karion et al., 2010), following the methods described by Wunch et al. (2010). We consider TCCON data to be coincident with the OCO-2 target-mode measurements when they have been recorded within ±30 minutes of the time at which the spacecraft is closest to nadir during the maneuver. If there are fewer than 5 TCCON data points recorded within that time, the window is extended to ±120 minutes, but this is required in only 14% of cases. We use the full OCO-2 version 7B retrospective data (i.e., 7Br), available from GES-DISC (2016, <http://disc.sci.gsfc.nasa.gov/OCO-2>), and manually apply the filters listed in Table 2. These filters are consistent with the “warn level 15” scheme described by Mandrake et al. (2015), except that the filter on the surface pressure difference from the prior in the A-band pre-processor is loosened, and we have added an additional outlier filter.

Figure 5 shows the OCO-2 X<sub>CO<sub>2</sub></sub> target-mode median data comparisons with coincident TCCON data. The best fit lines were computed using a method that accounts for uncertainties in the dependent and independent variables as described by York et al. (2004). Panel (a) shows the results prior to applying the parameter-dependent bias correction and has a correlation coefficient of  $R^2 = 0.75$ . Panel (b) shows the relationship after the correction has been applied and an improved correlation coefficient ( $R^2 = 0.83$ ). The improvement indicates that the parameter-dependent bias correction is effective at removing spurious variability in the OCO-2 data. The slope in panel (b), which has a y-intercept that is forced through 0, is used to derive the constant scaling factor between TCCON and OCO-2 target observations ( $m = 0.9975 \pm 0.04$ , which represents ~1 ppm) for the time period spanning September 8, 2014 through April 30, 2016. The time-dependence of the difference between the OCO-2 target-mode data and the coincident TCCON data ( $\Delta X_{\text{CO}_2}$ ), after the constant bias is removed, is plotted in Fig. 6. The algorithm, calibration and instrument cause no apparent time-dependent drift in  $\Delta X_{\text{CO}_2}$  nor their errors. Thus, the bias correction is successful at reducing both the parameter-dependent and constant global biases.

However, the target mode measurements are sensitive enough to point to some residual biases that are currently under investigation by the OCO-2 algorithm, calibration and validation teams. These residual biases are more geographically localized in nature, and as such might not be expected to be captured by the standard bias correction, which is designed to minimize globally-relevant biases. The residual biases seem to fall into two classes: biases in high southern latitude ocean glint data (§3.1), and biases caused by the surface properties at a target location (§3.2).

### 3.1 Southern hemisphere winter glint measurements over water

In the southern hemisphere winter, there is a significant high bias in the retrieved X<sub>CO<sub>2</sub></sub> from the OCO-2 ocean glint data. There were three target-mode measurements recorded in the southern hemi-



sphere during that time: two points over Wollongong, and a third point over Réunion recorded during  
215 late July/early August 2015 (Fig. 6). Figure 7 illustrates this problem by showing the divergence of  
the latitude gradients of the OCO-2  $X_{\text{CO}_2}$  in all viewing modes and the TCCON  $X_{\text{CO}_2}$  during Au-  
gust 2015 and the agreement in October 2015. Appendix Fig. A1 panels (r) and (s) also show this  
problem as a function of time. This bias does not impact the overall one-to-one line within the uncer-  
tainty but does impact the latitudinal gradients (and hence fluxes) inferred by the OCO-2 data. While  
220 the cause of the bias in the southern winter is currently unclear, there is a working hypothesis related  
to the OCO-2 algorithm's misrepresentation of stratospheric aerosols, exacerbated by the eruption  
of Mount Calbuco in Chile on April 22, 2015 (Romero et al., 2016). This is described in detail in  
O'Dell et al. (2016).

### 3.2 Surface Properties

225 Site-dependent differences from the one-to-one plot, shown in Fig. 8, reveal significant location-  
dependent biases. Any differences with magnitudes less than 0.3 ppm could be attributable to TC-  
CON station site-to-site biases (Wunch et al., 2010), so we focus on the biases that are significantly  
larger and thus attributable to the OCO-2 data. Two clear examples of site-dependent biases are at  
Edwards, with a mean low bias of  $\sim 1.3$  ppm, and Wollongong, with a mean high bias of  $\sim 0.9$  ppm.  
230 The spatial dependence of the target-mode measurements reveal that small-scale variability in sur-  
face properties (e.g., albedo, altitude, surface roughness) can cause significant and spurious variabil-  
ity in the  $X_{\text{CO}_2}$ .

The Edwards TCCON station is situated in the bright California high desert on the edge of a  
very bright playa with little topographic change. On some, but not all targets over Edwards, the  
235  $X_{\text{CO}_2}$  appears dependent on surface brightness (Fig. 9). All mean target-mode measurements of  
 $X_{\text{CO}_2}$  at Edwards are biased lower than the TCCON measurements. Conversely, the Wollongong  
station, which is situated near the East coast of Australia, is a very dark surface in the visible, and  
lies between the Tasman Sea to the East and the Illawarra Escarpment to the West (Fig. 10). The  
OCO-2 retrievals of  $X_{\text{CO}_2}$  are systematically higher than the TCCON measurements for all target-  
240 mode measurements, and are particularly high (up to 5 ppm) in July and August (Fig. 11), due to  
the problem discussed above in §3.1. OCO-2 data over Białystok, located in a dark, forested region,  
also has a persistent high bias (on the order of 1.3 ppm) compared with TCCON.

Even at sites which do not appear to have a significant bias with respect to TCCON can show spu-  
rious spatially-correlated errors. The Lauder TCCON station is situated in a valley between rolling  
245 hills (Fig. 12). The surface altitude is spatially well-correlated with changes in  $X_{\text{CO}_2}$  during each  
target-mode maneuver.

In the standard OCO-2 retrieval algorithm, land surfaces are assumed to have a Lambertian bidi-  
rectional reflectance distribution function (BRDF, O'Dell et al., 2012). A more complex BRDF  
that can better model changes in surface topography and vegetation is currently undergoing testing





250 and shows promise for reducing the problems addressed above (Natraj et al., 2015). An improved treatment of the surface properties will be included in the next version of the OCO-2 data.

#### 4 Nadir and Glint Mode Comparisons to TCCON

In this section, we evaluate the bias-corrected OCO-2 glint and nadir modes against ground-based TCCON data to reveal mode-dependent biases that were not eliminated using the standard version  
255 7 bias correction. We use the version 7B retrospective OCO-2 “lite” files here, which have had the footprint-dependent, parameter-dependent, and constant scaling biases removed, and we limit ourselves to data for which the “warn level” is less than or equal to 11 (Mandrake et al., 2013, 2015), and the “outcome\_flag” is zero. These data are available from the CO<sub>2</sub> Virtual Science Data Environment (JPL-Caltech, 2016, <http://co2.jpl.nasa.gov>) and from GES-DISC (2016). For these comparisons, we choose the following coincidence criteria: a box centred around the TCCON station that  
260 spans 5° in latitude and 10° in longitude on the same day as a TCCON measurement, with the exceptions mentioned below. In the southern hemisphere south of 25°S, we use a larger box spanning 20° in latitude and 120° in longitude because the geographical variance in X<sub>CO<sub>2</sub></sub> in the southern hemisphere is low (e.g., Wunch et al., 2011b). The Edwards and Pasadena boxes are constructed  
265 differently because they are geographically very close to each other, but the Pasadena site is within the polluted, mountain-contained South Coast Air Basin, and Edwards is in the clean desert north of the mountains. Thus, we limit the Edwards latitudes to north of Edwards, but allow the longitudes to span 5° further west over the Pacific Ocean. The Pasadena coincidence box is constrained to the South Coast Air Basin, which significantly limits the number of coincident points (see Appendix  
270 Figs. A1(a-t)). The median of the measurements within the coincidence box on the same day as the TCCON measurement is compared with the TCCON daily median. The more complicated dynamical coincidence criteria used to increase the number of coincident measurements between TCCON and GOSAT in Wunch et al. (2011b) and Nguyen et al. (2014) are not required for OCO-2, due to OCO-2’s much higher data density.

275 The overall comparisons between the OCO-2 data and TCCON data are reported in Tables 3 and 4, and shown in Figs. 13–15 for data from land glint mode, ocean glint mode, and nadir mode. The biases between OCO-2 and TCCON are all less than 0.5 ppm, and the RMS of the difference is less than about 1.5 ppm. The nadir mode data show the best correlation of the three science modes ( $R^2 = 0.75$ ), followed by land glint ( $R^2 = 0.72$ ), and finally ocean glint ( $R^2 = 0.50$ ). The low  
280 correlation coefficient in the ocean glint data is partially driven by the high anomalies in the southern hemisphere winter, most obviously in the data over Wollongong (Fig. 7), and also because a high proportion of ocean glint data are in the southern hemisphere, where X<sub>CO<sub>2</sub></sub> variability is lower. If the southern hemisphere winter data (June–September) are excluded from the ocean glint correlations, the  $R^2$  improves to 0.71. The slopes of all three curves are significantly different from 1.0. The



285 agreement between the science-mode OCO-2 data and TCCON is poorer than that for the target-  
mode measurements. Halving the spatial coincidence criteria over each site does not significantly  
improve the correlation coefficients ( $R^2 = 0.78, 0.73, 0.58, 0.77$  for nadir, land glint, ocean glint,  
and ocean glint excluding the southern hemisphere winter, respectively) and the slopes increase, but  
well within the uncertainties. This suggests that it is not solely our definition of the coincidence  
290 criteria that causes the low correlation coefficients, and that perhaps the surface properties within  
the coincidence boxes contain sufficient variability to degrade the comparisons. This highlights the  
importance of the target-mode data for assessing local, site-to-site, and overall bias.

## 5 Conclusions

The OCO-2  $X_{CO_2}$  estimates generally compare well with coincident TCCON data at global scales,  
295 with mean biases less than 0.5 ppm and RMS differences less than about 1.5 ppm. The best com-  
parisons are with the target-mode data, which is expected because the target mode measurement  
scheme was designed for this purpose. The target-mode data are also particularly well-suited to ty-  
ing the OCO-2 data to the WMO trace gas scale through comparisons with the TCCON data, because  
they best represent coincident measurements under invariant atmospheric conditions. While the bias  
300 correction clearly improves the relationship between TCCON and OCO-2 globally, smaller-scale  
biases, typically  $< 2$  ppm, remain. Examples of these biases include spurious local  $X_{CO_2}$  variability  
correlated with topography and surface brightness, and ocean glint measurements at southern lati-  
tudes during the southern hemisphere winter. Remedying these residual biases is the current focus  
of the OCO-2 algorithm development and validation teams and we anticipate that the next version of  
305 the OCO-2 data will represent a significant improvement. It is imperative to continue measurement  
comparisons with TCCON in all three modes (target, glint and nadir) to monitor and evaluate the  
OCO-2 data quality throughout its entire mission.

## 6 Data Availability

Unfiltered, uncorrected OCO-2 data are available from the Goddard Data Center (GES-DISC, 2016).  
310 The filtered and bias-corrected data are contained in “lite” files, which are available from JPL’s CO<sub>2</sub>  
portal (JPL-Caltech, 2016). TCCON data are available from the TCCON data archive, hosted by  
CDIAC: <http://tcon.ornl.gov>. Each TCCON dataset used in this paper is cited independently in  
Table 1 or in the captions of Fig. A1.

## Appendix A: Site Plots

315 The ocean glint, land glint and nadir mode plots for each TCCON station are shown in Fig. A1. In  
each plot, there are four panels. The top left panel shows the time series of the TCCON daily median



data (black circles) and the OCO-2 data (triangles coloured differently for each mode). The bottom left panel shows the difference between OCO-2 and TCCON measurements (OCO-2 – TCCON). The top right panel shows the correlations between the TCCON data and the OCO-2 data. The  
320 bottom right panel shows the coincidence area for the OCO-2 measurements.

### Appendix B: Author contributions

DW wrote the manuscript and produced the main analysis and results with significant input from GO, CV, and POW. DW, GO, BF, BN, CMR, CO, AE, LM, CV, MRG, DC, and POW contributed to the experiment design and analysis of data. DWTG, NMD, VAV, JN, TW, CP, MdM, MKS, RS,  
325 MR, DP, JR, IM, OU, FH, TB, MK, DGF, SGA, KS, JM, RK, PH, LI, JP, PH, SK, MD, HAP, ES, OEGR, YT, PJ, provided TCCON data. All authors read and provided comments on the manuscript.

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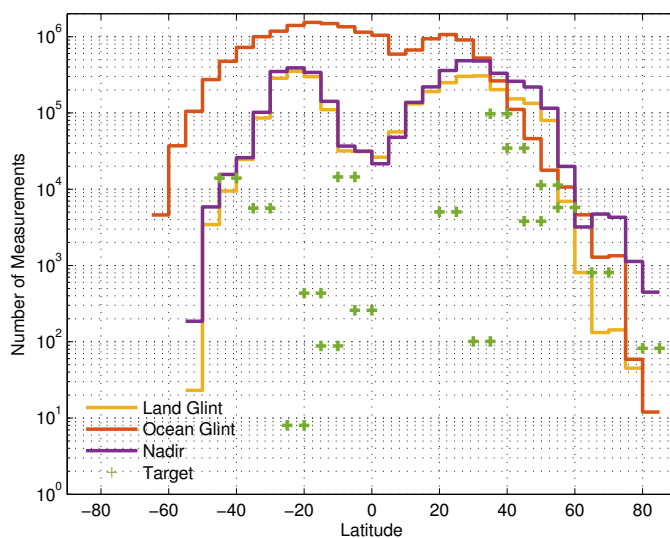


Figure 1: Nadir, and glint, and target-mode measurement density in 5° bins as a function of latitude from the beginning of the mission through April 30, 2016. These are from the “lite” files applying “warn level” 11 filters and requiring that the “outcome\_flag” is zero.

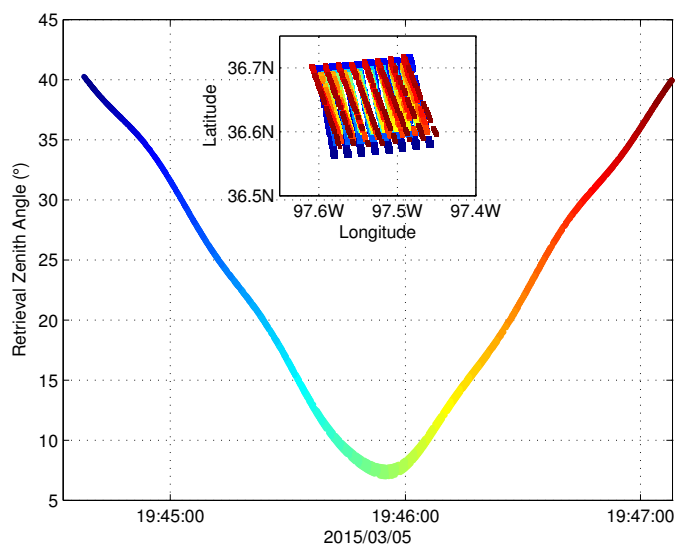


Figure 2: The zenith angles viewed during a target mode measurement over Lamont. The spacecraft “nods” across the ground target as it rotates overhead. The colours represent the time of the measurement. The top inset shows the locations of the measurements in latitude and longitude.



Figure 3: Map of OCO-2 target locations. Gold circles show the locations of the targets that coincide with TCCON stations; orange stars show the locations of targets that do not have co-located TCCON stations.

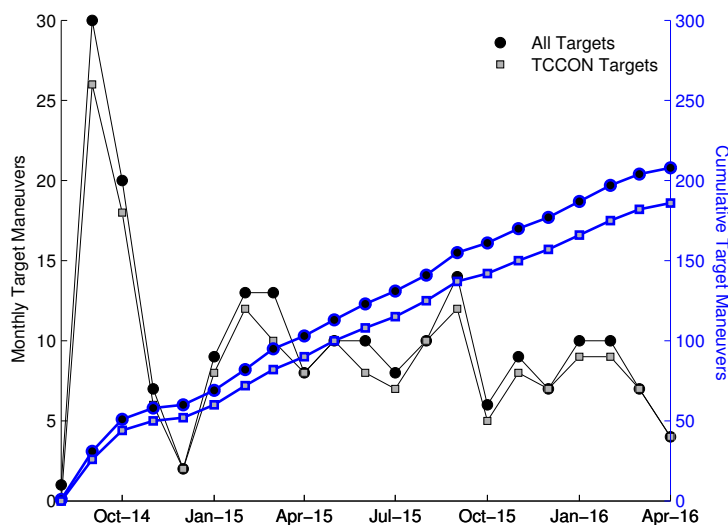
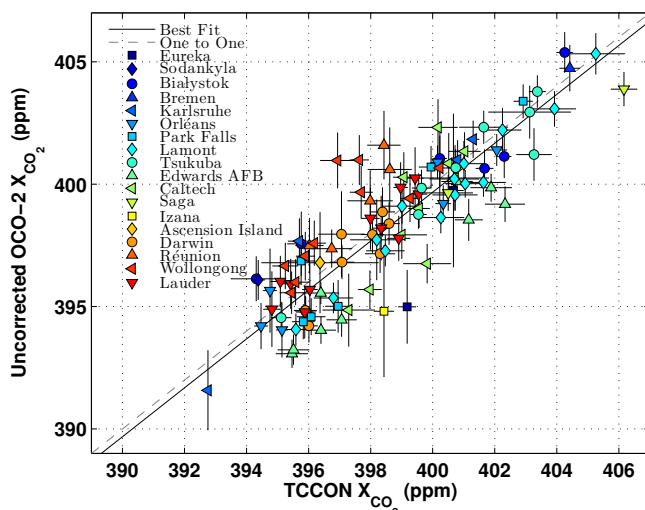
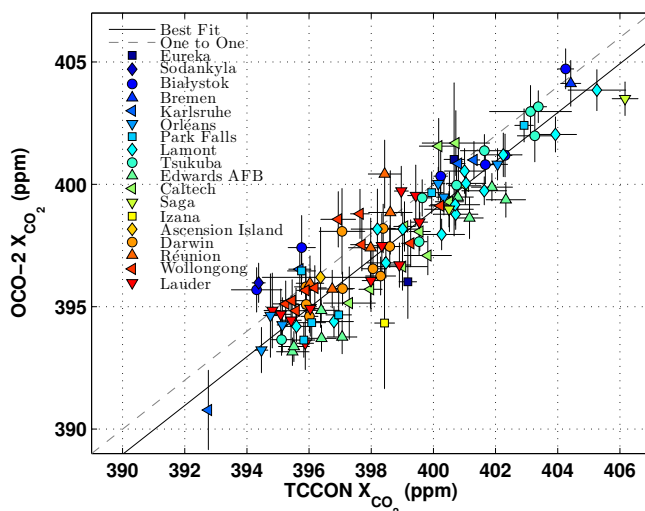


Figure 4: The number of target mode maneuvers attempted per month (left axis), and the cumulative number of target mode maneuvers attempted to date (right axis). The TCCON targets are in grey squares; all targets (including Railroad Valley, which does not have a TCCON station) are marked in black circles.



(a) Without bias correction.



(b) With bias correction.

Figure 5: The relationship between the OCO-2 target-mode data and the coincident TCCON data. The top plot (a) does not have the Mandrake et al. (2015) bias correction applied, the bottom plot (b) is after bias correction, but before the scaling is applied. The one-to-one line is indicated by the dashed black line, and the best fit is marked in the solid black line.

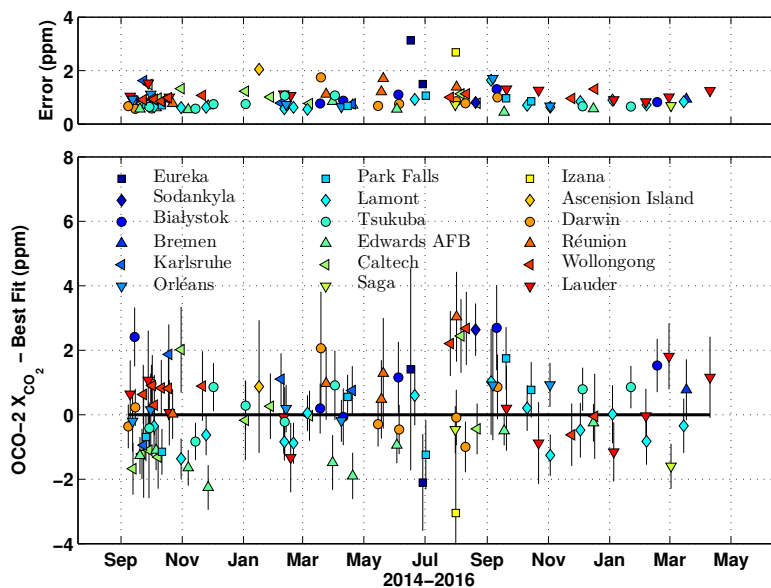


Figure 6: The time series of the differences between the OCO-2 target-mode data and the best fit line in Fig. 5(b). The top panel shows the magnitude of the sum in quadrature of the standard deviation of the OCO-2 data during the target and the standard deviation of the coincident TCCON data. Those values are plotted as the error bars in the lower panel.

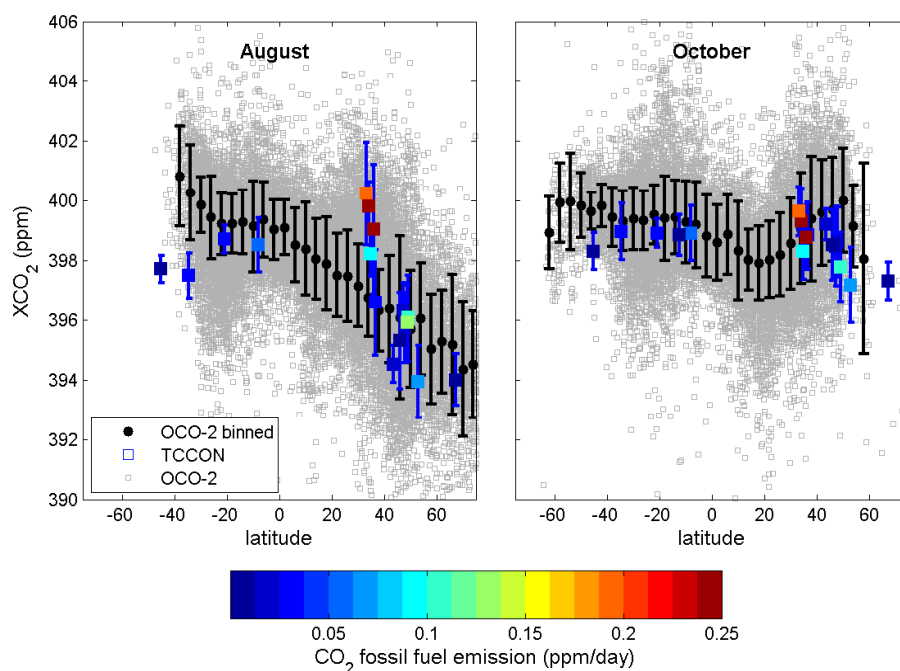


Figure 7: The latitudinal gradient of OCO-2 data in August 2015 (left), and in October 2015 (right). The grey squares are monthly mean OCO-2 data from glint and nadir mode gridded in  $0.5^\circ \times 0.5^\circ$  bins. The black circles are the OCO-2 zonal means after gridding onto a  $4^\circ \times 5^\circ$  grid. The error bars indicate the  $1\sigma$  standard deviation. The TCCON monthly medians are marked by the squares with  $1\sigma$  standard deviation error bars. The colours represent the  $\text{CO}_2$  fossil fuel emissions within 50 km of the TCCON location. The fossil fuel emission source is from the European Commission, Joint Research Centre (JRC)/Netherlands Environmental Assessment Agency (PBL). Emission Database for Global Atmospheric Research (EDGAR), <http://edgar.jrc.ec.europa.eu>. The high bias of the OCO-2 data in the higher latitude southern hemisphere (which is dominated by glint measurements over water) is clearly evident in the August plot south of  $-20^\circ$ .



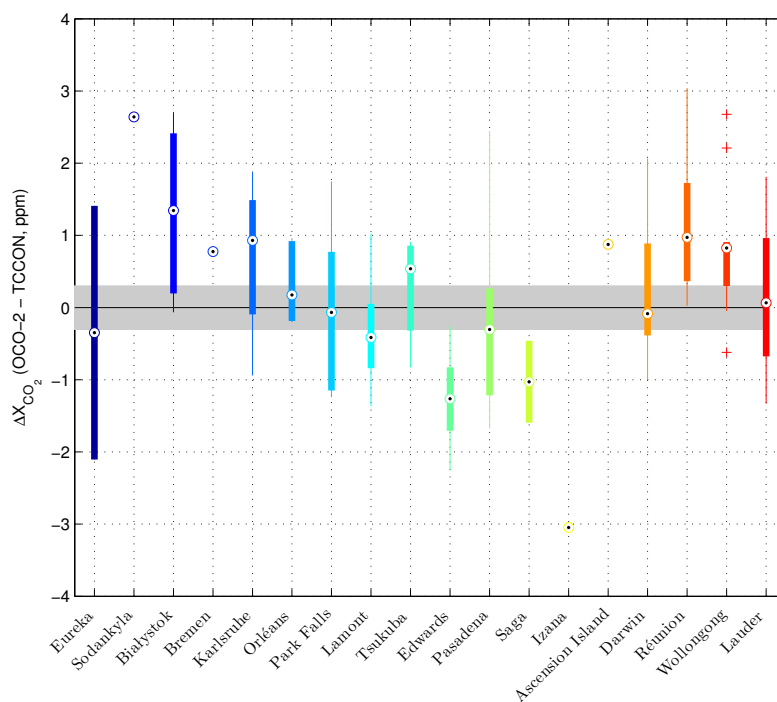


Figure 8: The site-to-site differences between the OCO-2 target-mode data and the coincident TCCON data. This is a “box plot”: the circles indicate the median value of the difference, the thick bars indicate the 25 and 75 percentile limits, the thin bars represent the full range of the data, excluding the outliers (McGill et al., 1978). The outliers are represented by plus (+) symbols. The grey shaded area indicates the  $\pm 0.3$  ppm uncertainty in the TCCON values: deviations beyond the shading can be attributed to uncertainties in the OCO-2 data.

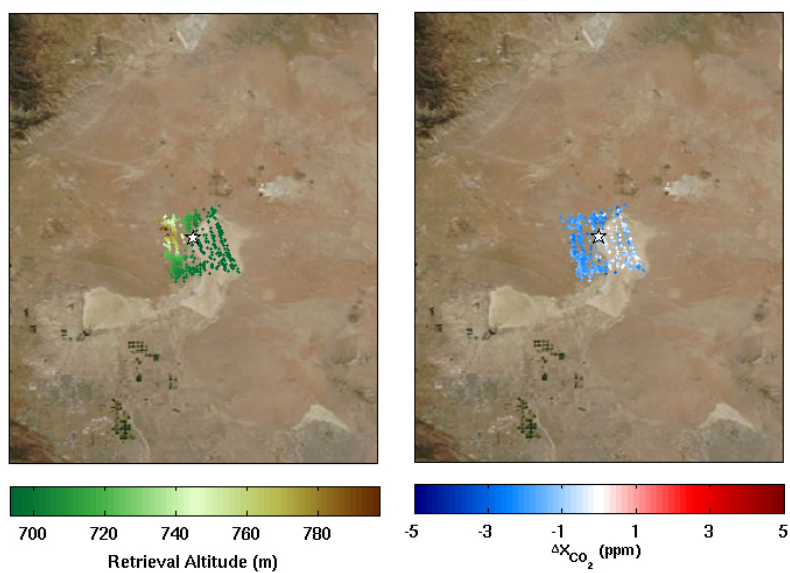


Figure 9: Edwards target on April 19, 2015. The background is the MODIS true-colour image of the Edwards area at the time of the target-mode measurements. The white star indicates the location of the Edwards TCCON station. The left panel shows the elevation model of the surface and the right panel shows the difference in OCO-2  $X_{\text{CO}_2}$  from the value recorded by the TCCON instrument. A spatial bias related to the surface brightness is clearly present in this target-mode measurement. In other Edwards target mode measurements, this surface brightness-correlated bias is not as apparent.



Figure 10: The retrieval altitudes near the Wollongong TCCON station during target measurements. The sharp Illawarra escarpment can be clearly seen inland from the Tasman Sea (to the East).

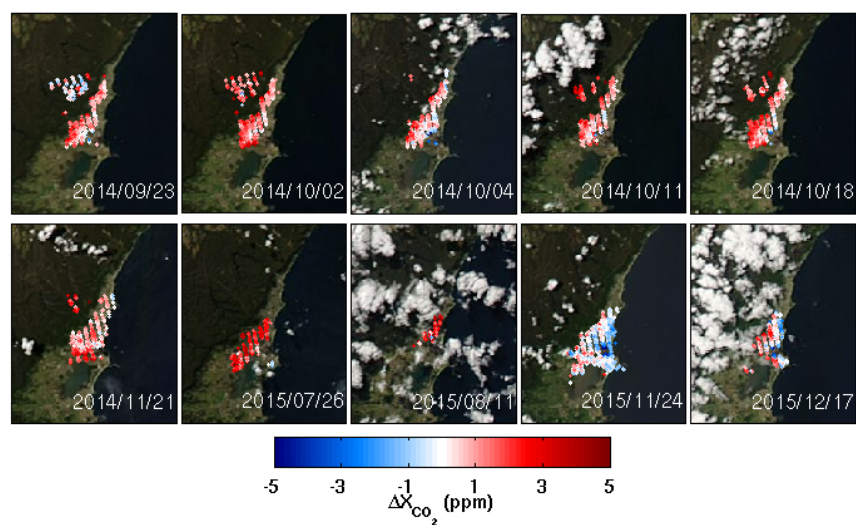


Figure 11: The filtered target-mode measurements over Wollongong. The colours represent the difference between the OCO-2 measurement and the coincident TCCON measurement. The OCO-2 data over Wollongong are generally higher (redder) than the TCCON measurements, and significantly high in the July and August 2015 target-mode maneuvers.

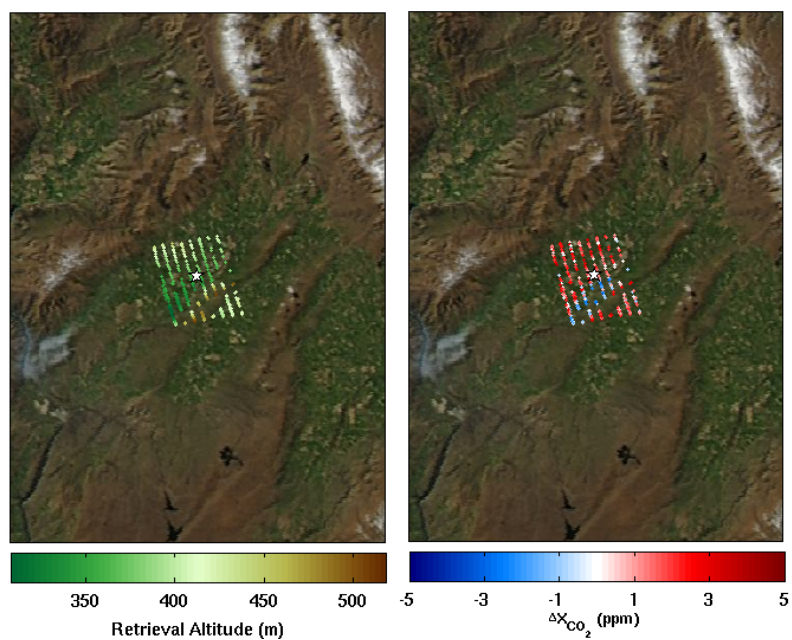


Figure 12: Lauder target on September 28, 2014. The background is the MODIS true-colour image of the Lauder area at the time of the target-mode measurements. The white star indicates the location of the Lauder TCCON station. The left panel shows the elevation model of the surface and the right panel shows the difference in  $X_{\text{CO}_2}$  from the value recorded by the TCCON instrument. A spatial bias is clearly present, related to the surface elevation.

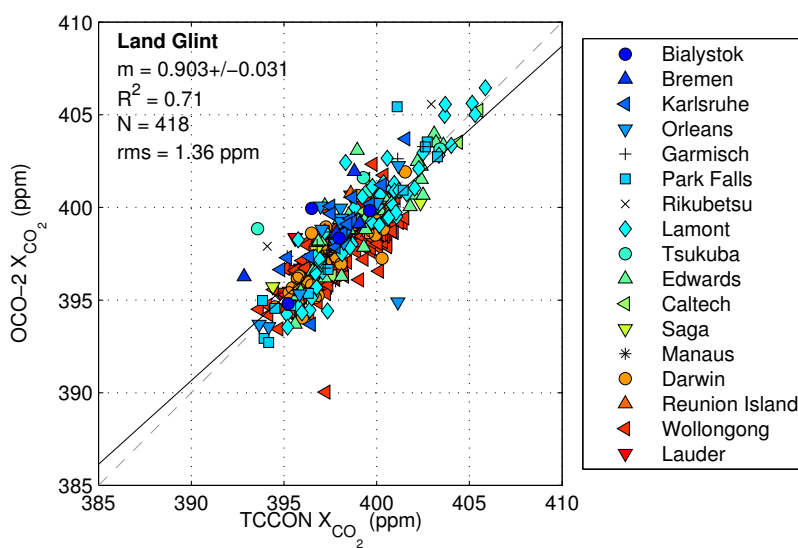


Figure 13: Land glint OCO-2 one-to-one plot against TCCON. The slope of the relationship is represented by “m” in the figure, and the coefficient of determination is represented by “R<sup>2</sup>”. The number of points on the graph is indicated by “N” and the root-mean-square value (rms) of the differences between OCO-2 and TCCON X<sub>CO<sub>2</sub></sub> is also shown.

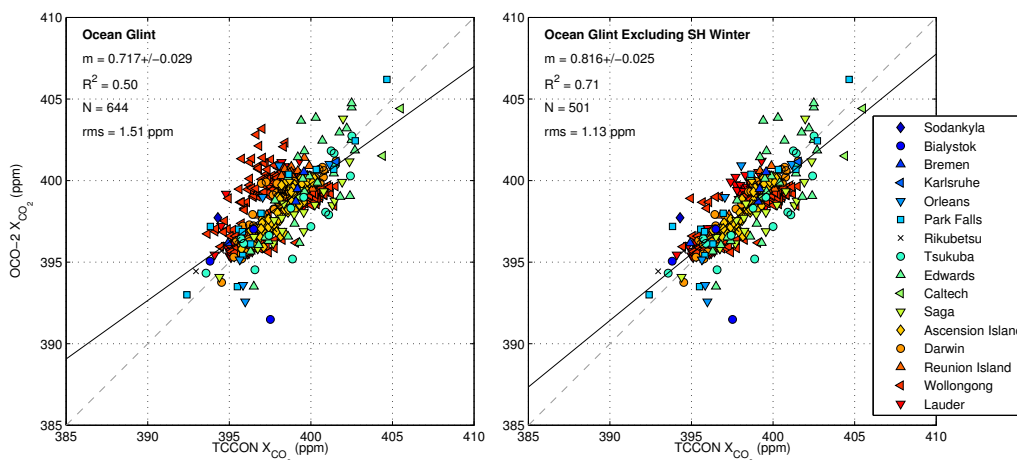


Figure 14: Ocean glint OCO-2 one-to-one plot against TCCON. The left panel shows all the glint-mode data. The right panel removes the southern hemisphere wintertime (June through September) glint data that has a known high bias.

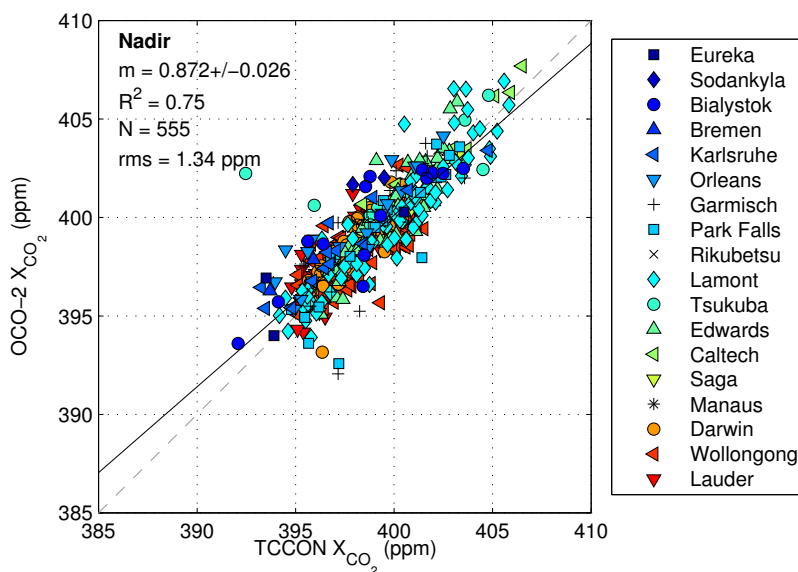
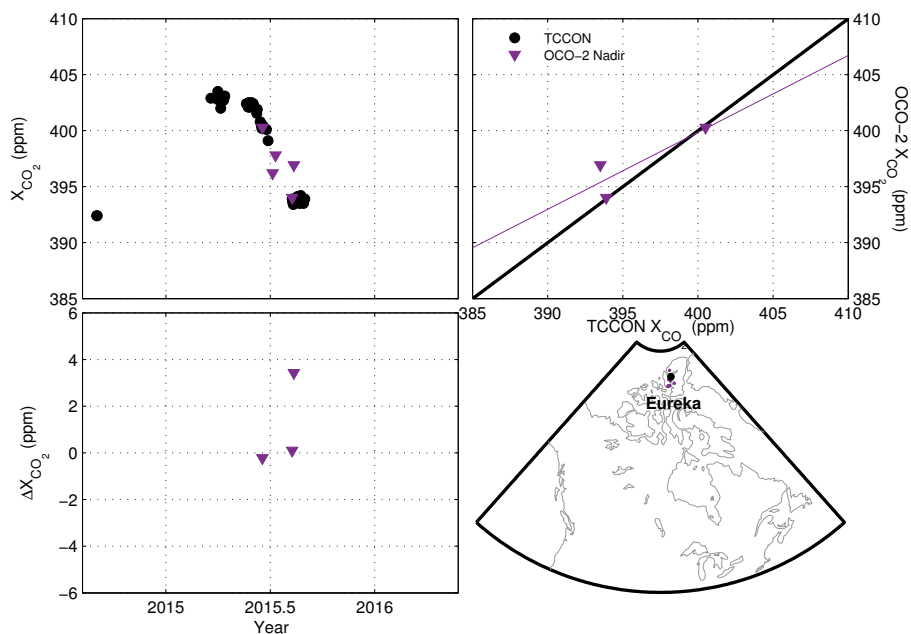
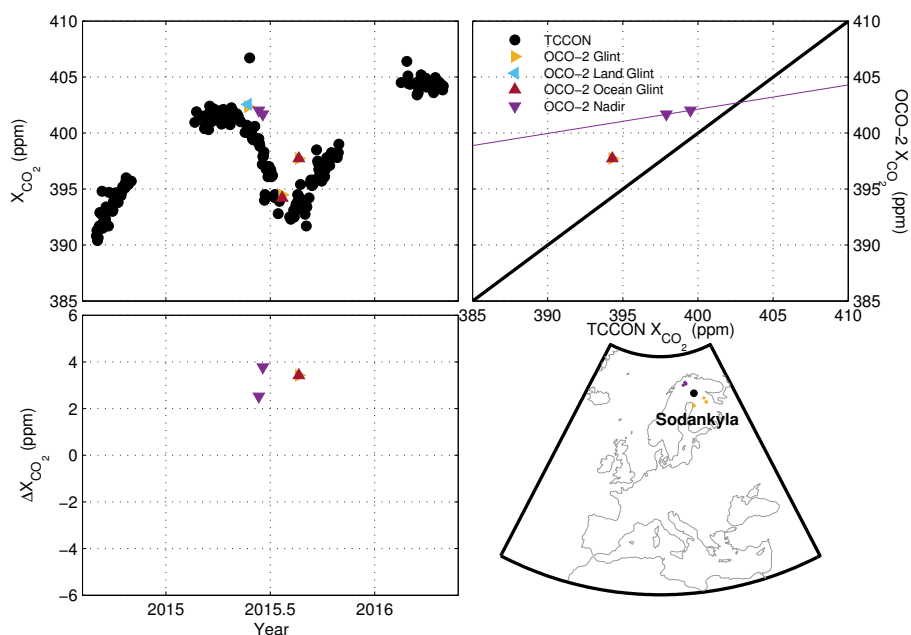


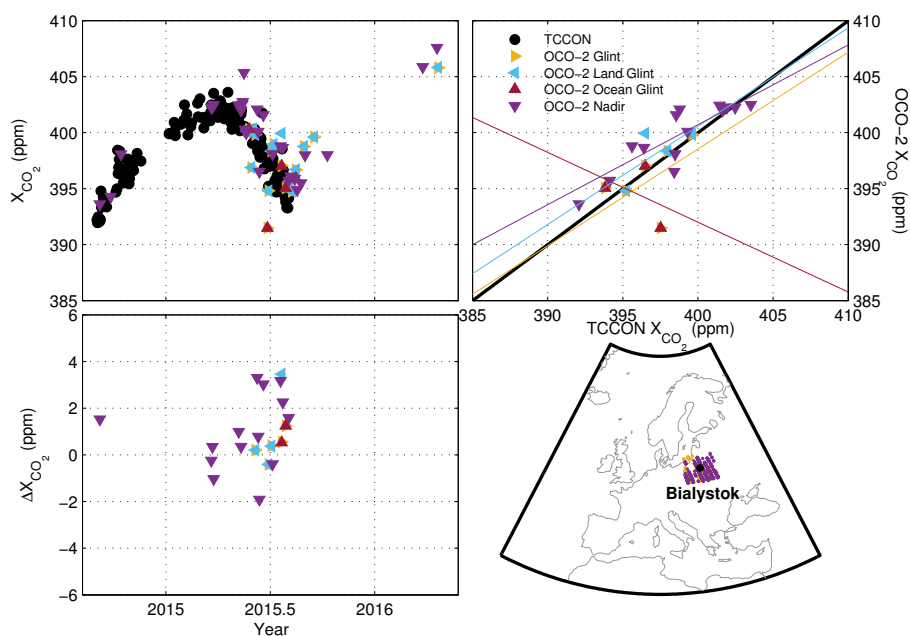
Figure 15: Nadir OCO-2 one-to-one plot against TCCON.



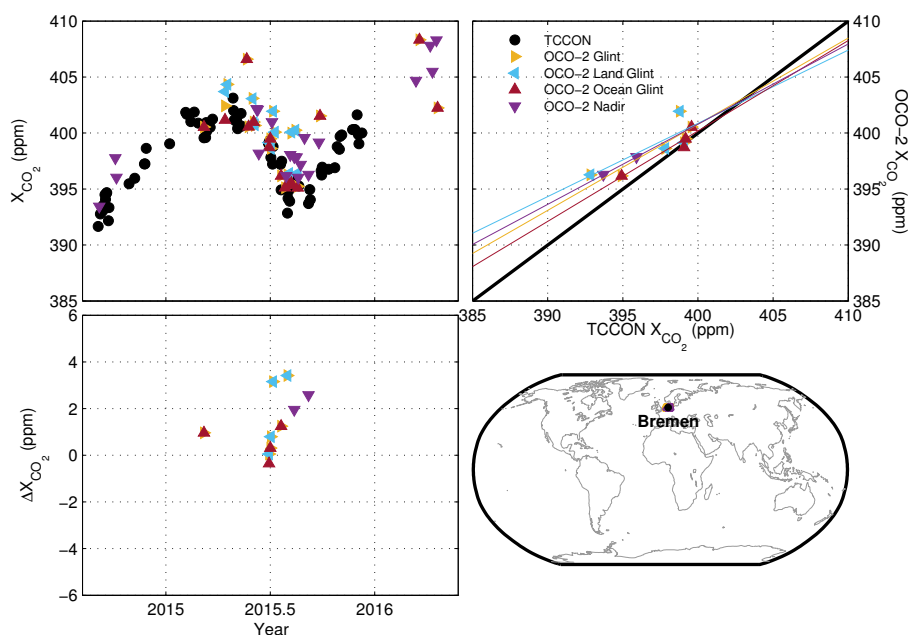
(a) Eureka



(b) Sodankylä

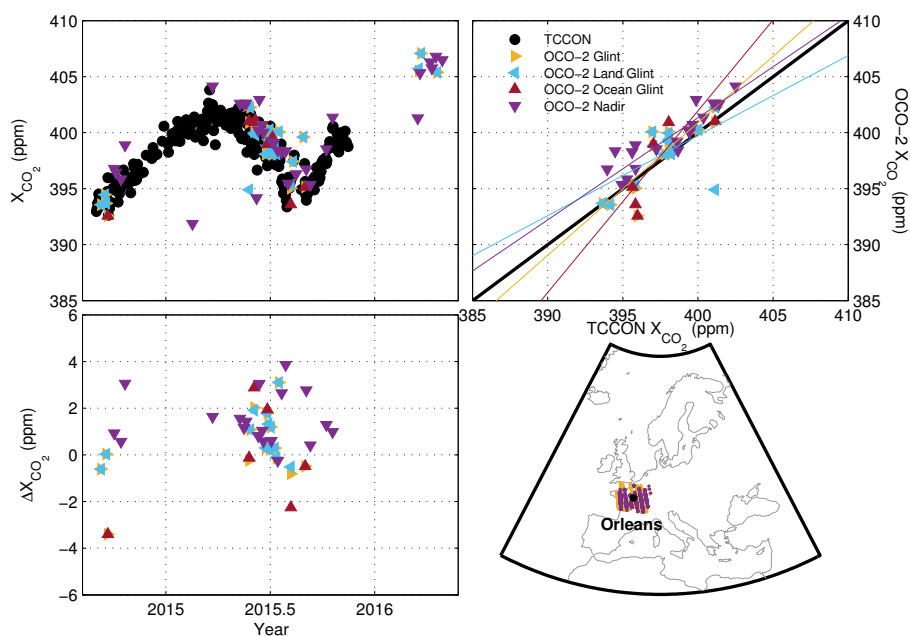


(c) Bialystok

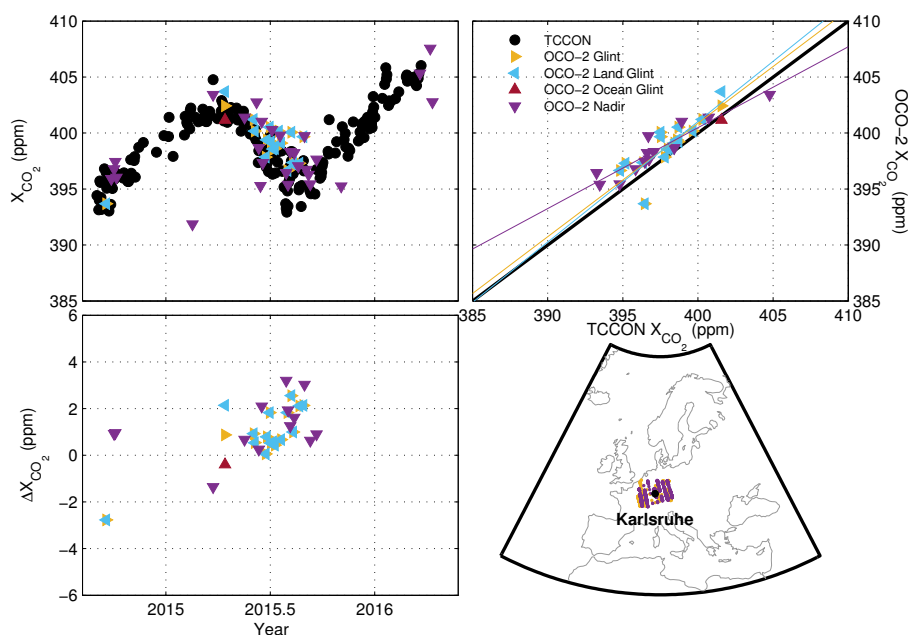


(d) Bremen

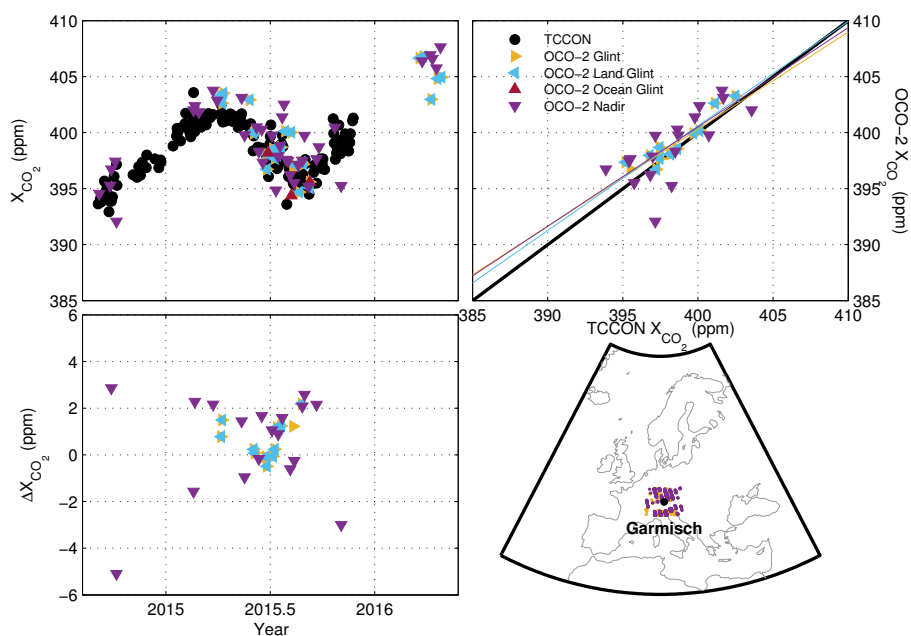




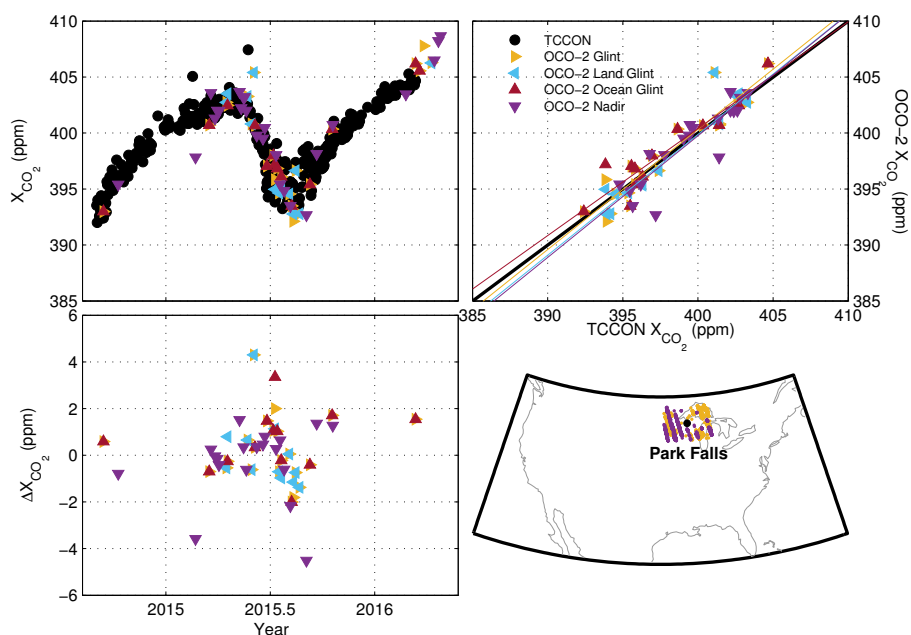
(e) Orléans



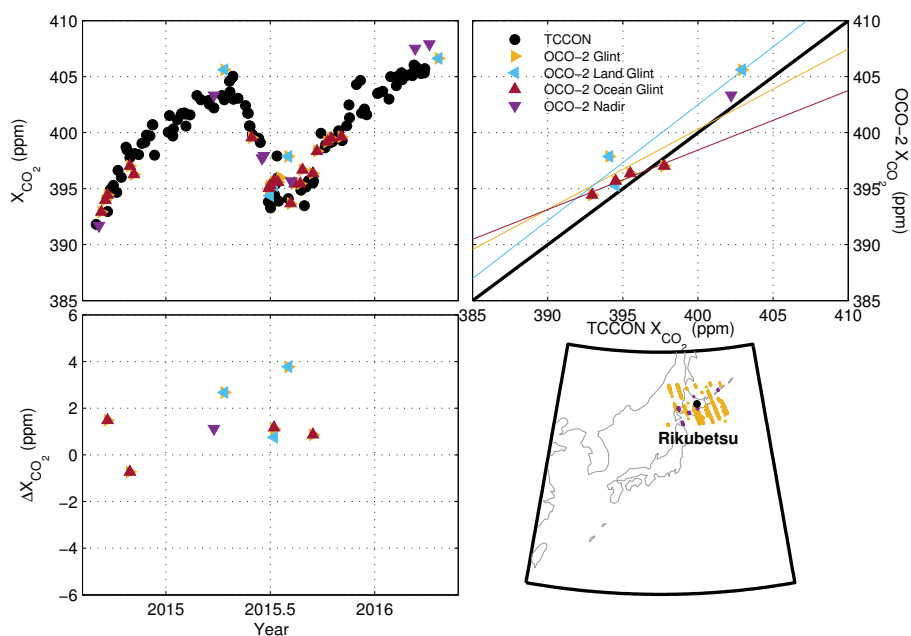
(f) Karlsruhe



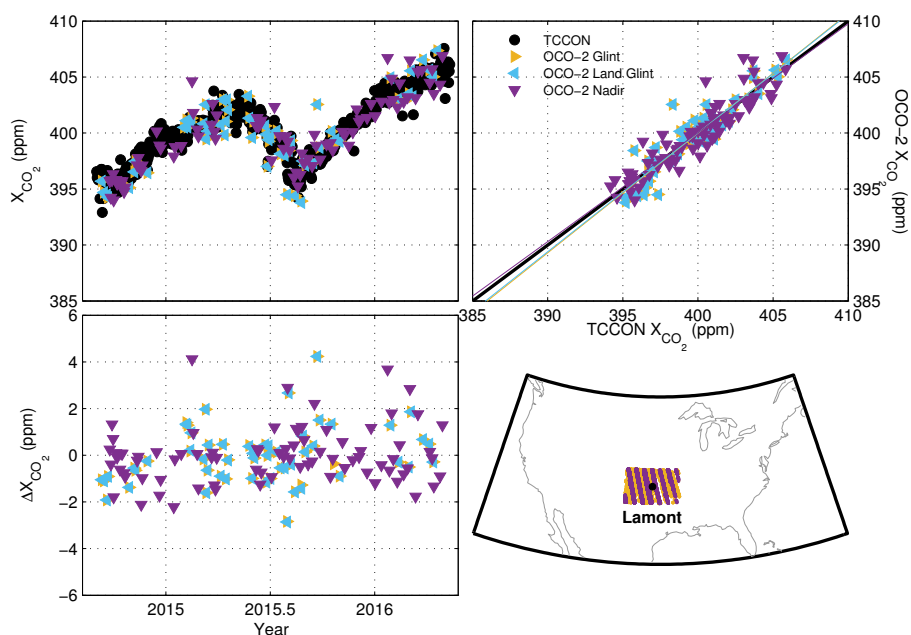
(g) Garmisch: Sussmann and Rettinger (2014)



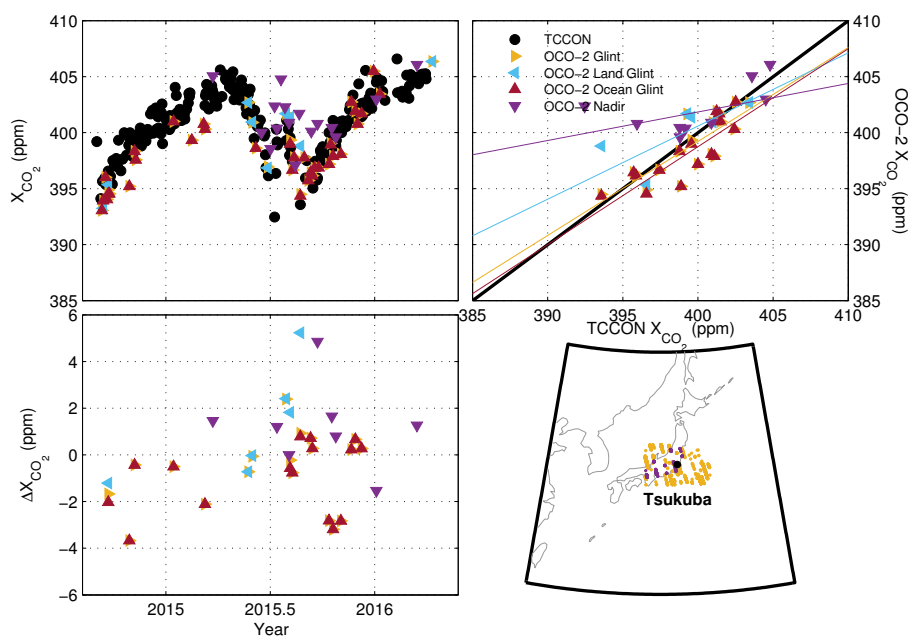
(h) Park Falls



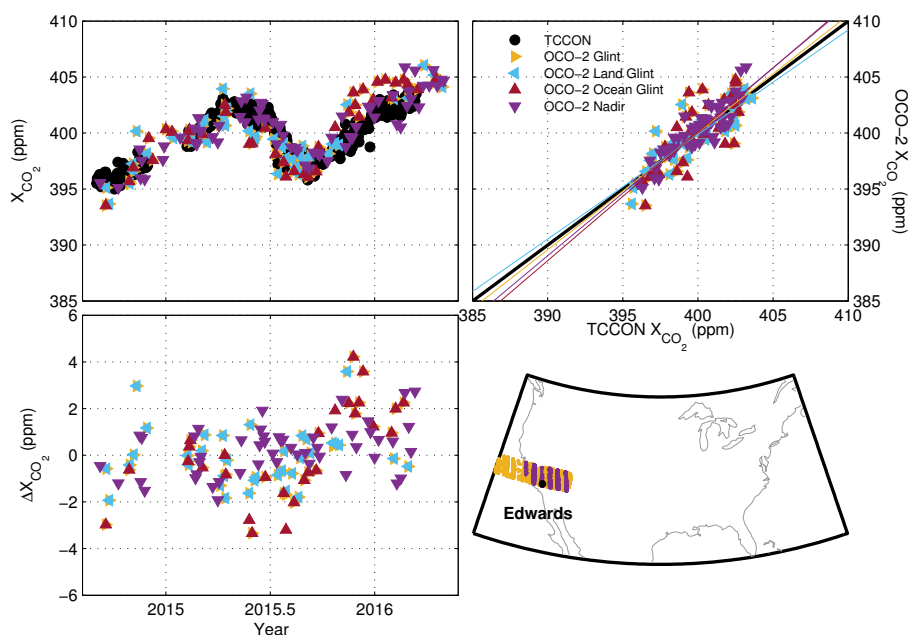
(i) Rikubetsu



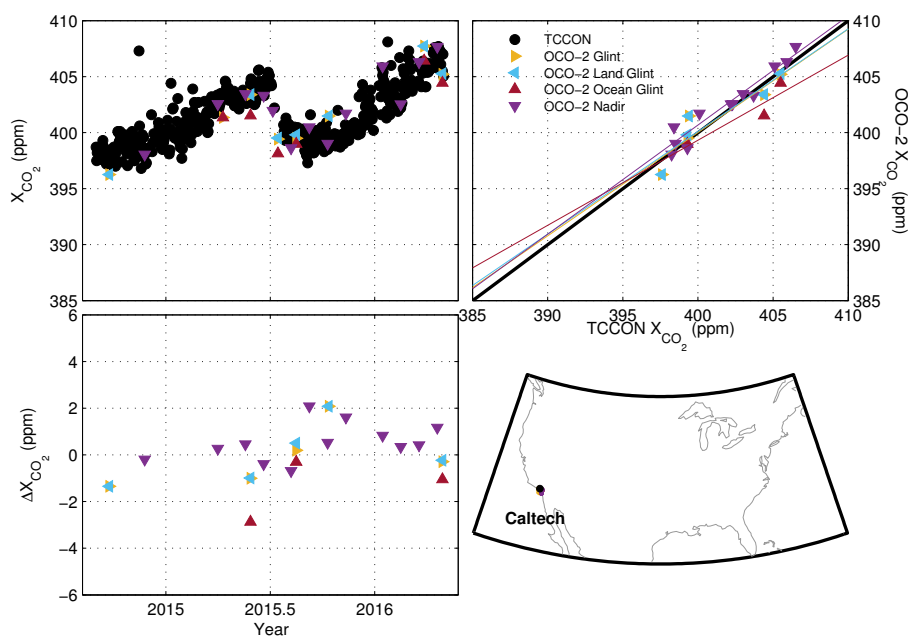
(j) Lamont



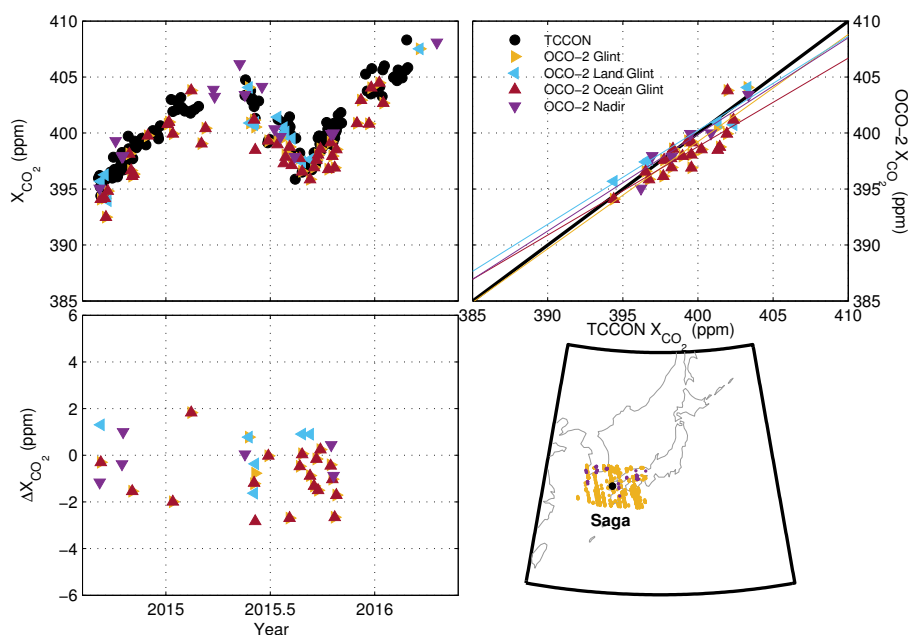
(k) Tsukuba



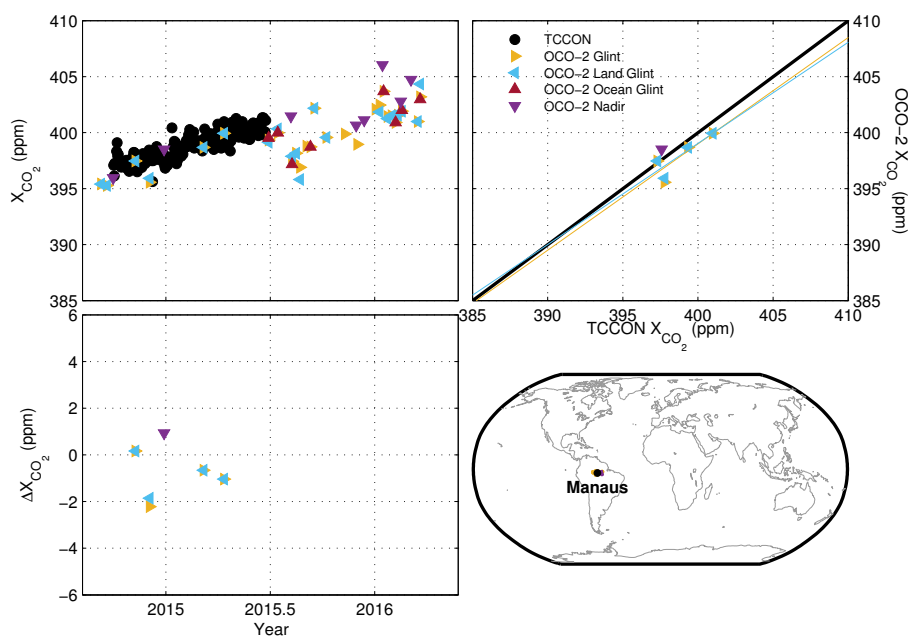
(l) Edwards



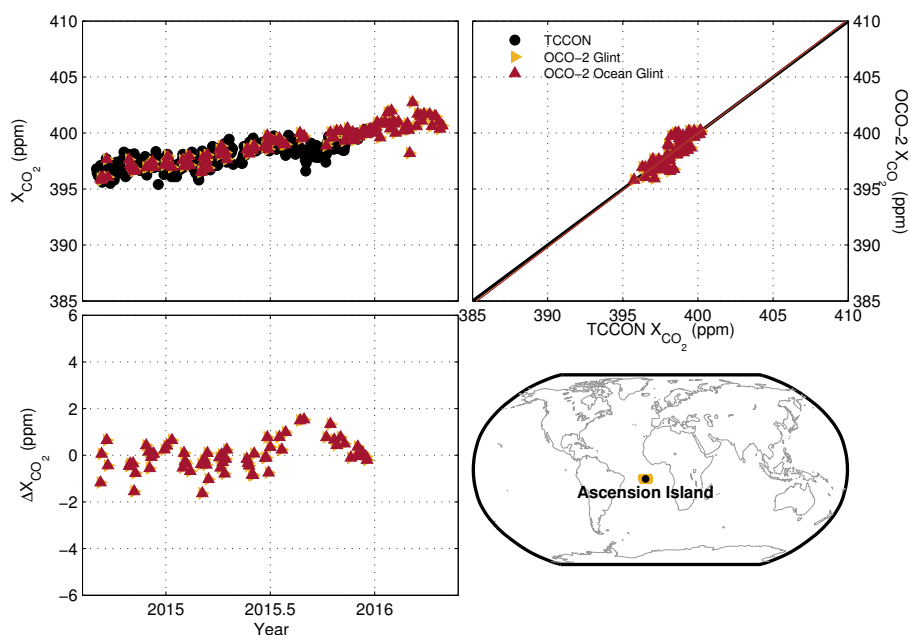
(m) Pasadena



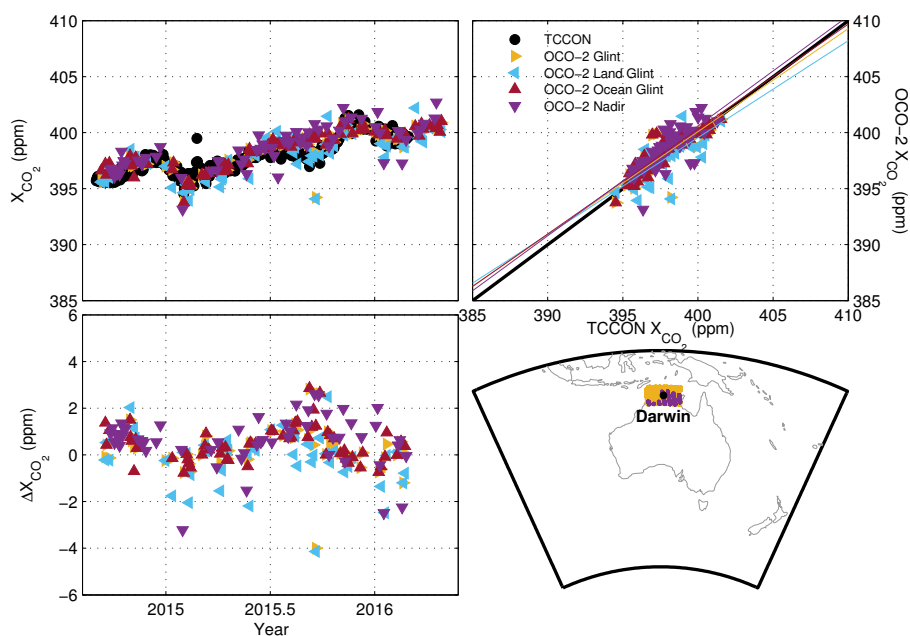
(n) Saga



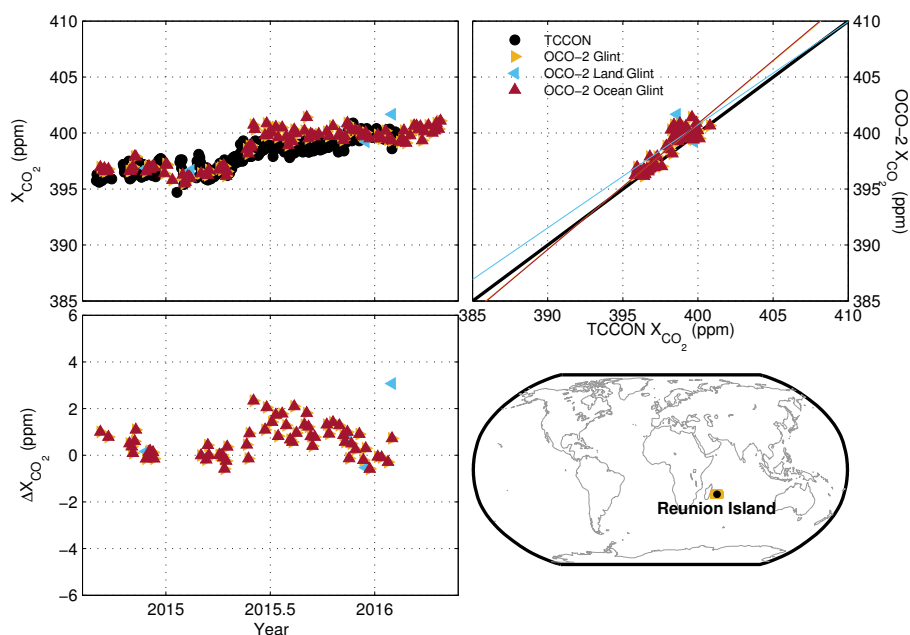
(o) Manaus



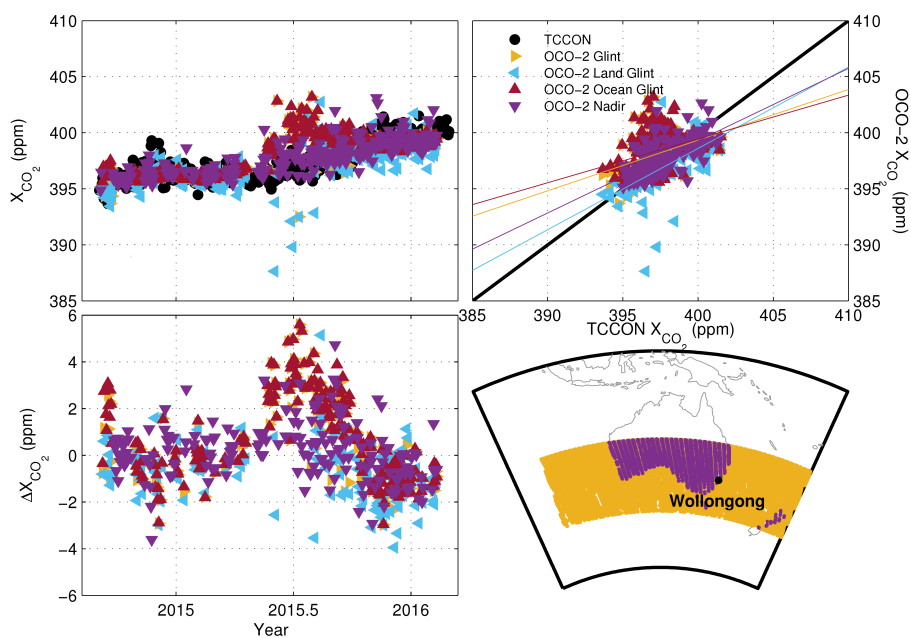
(p) Ascension



(q) Darwin

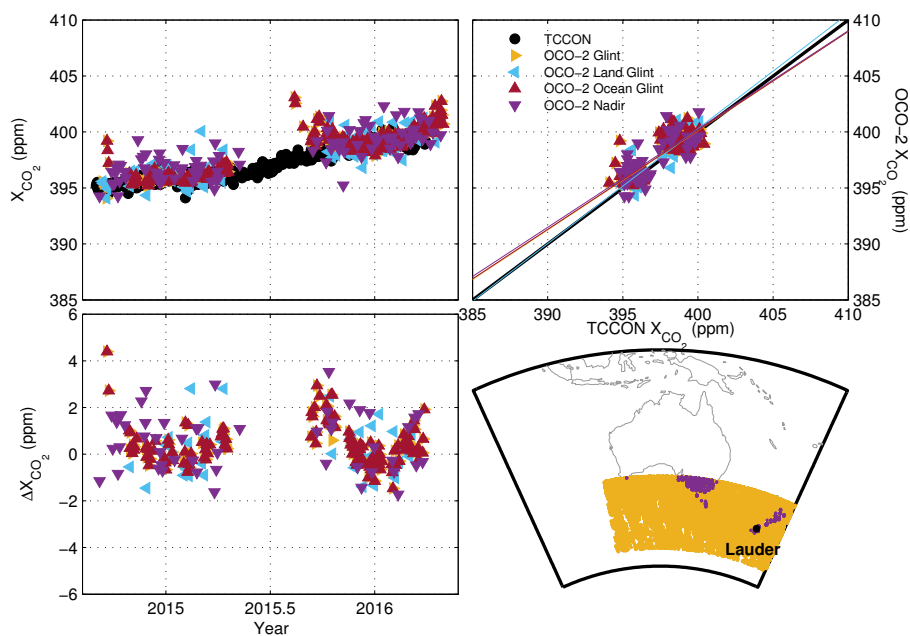


(r) Réunion Island



(s) Wollongong





(t) Lauder

Figure A1: The top left panel of each plot (a–t) shows the time series of the TCCON daily medians (black circles), and the daily medians of the OCO-2 glint mode (gold triangles), split into land glint (blue triangles) and ocean glint (red triangles), and OCO-2 nadir mode (purple triangle). The bottom left panel shows the difference between the OCO-2 data and TCCON data as a function of time. The top right panel shows the one-to-one correspondence between the OCO-2  $X_{CO_2}$  values and the TCCON values, and the best fit lines in the colours corresponding to the symbols. The one-to-one line is marked in black. The lower right panel shows the location of the TCCON station (black circle), and the locations of the OCO-2 data, showing glint-mode data in gold and nadir-mode data in purple. The lower right panel is intended to give a sense of the spatial coincidence criteria applied to the OCO-2 data for each TCCON station.



Table 1: Available targets. Those available only prior to July 2015 are marked with a natural symbol ( $\natural$ ). Those only available after July 2015 are marked with a sharp symbol ( $\sharp$ ). Note that the Target Location may not be exactly centered on a TCCON site location. Targets without a corresponding TCCON station are marked with a flat symbol ( $\flat$ ) and are not discussed in this paper.

Target Name	Target Location (Lat, Lon, Alt)	Target Active Dates	Data Reference
Anmyeondo, South Korea $\sharp$	36.624, 126.373, 0.006	July 2015 - Present	
Ascension Island	-7.947, -14.387, 0.165	July 2014 - Present	Feist et al. (2014)
Bialystok, Poland	53.196, 23.0758, 0.124	July 2014 - Present	Deutscher et al. (2014)
Boulder, CO $\flat$	40.014, -105.104, 1.61	July 2015 - Present	
Bremen, Germany	53.104, 8.850, 0.004	July 2014 - Present	Notholt et al. (2014)
Caltech, Pasadena, CA	34.123, -118.073, 0.157	July 2014 - Present	Wennberg et al. (2014c)
ARM TWP - Darwin, Aus	-12.375, 130.917, 0.0049	July 2014 - Present	Griffith et al. (2014a)
Edwards FRC, CA	34.958, -117.882, 0.699	July 2014 - Present	Iraci et al. (2014)
Izaña, Tenerife, Spain	28.297, -16.518, 2.2317	July 2014 - Present	Blumenstock et al. (2014)
Karlsruhe, Germany	49.100, 8.438, 0.11	July 2014 - Present	Hase et al. (2014)
Eureka, Canada $\natural$	80.053, -86.417, 0.601	July 2014 - June 2015	Strong et al. (2014)
SGP ARM Site, Lamont OK	36.604, -97.486, 0.3179	July 2014 - Present	Wennberg et al. (2014b)
Lauder, NZ	-45.002, 169.685, 0.384	July 2014 - Present	Sherlock et al. (2014)
Libya	28.550, 23.390, 0.108	June 2016 - Present	
Litchfield, Aus	-17.151, 139.795, 0.233	June 2016 - Present	
Manaus, Brazil	-3.213, -60.598, 0.04877	July 2014 - June 2016	Dubey et al. (2014)
Mexico City, Mexico $\flat$ $\sharp$	19.429, -99.138, 2.239	July 2015 - Present	
Orléans, France	47.965, 2.113, 0.1308	July 2014 - Present	Warneke et al. (2014)
Paris, France $\sharp$	48.846, 2.356, 0.034	July 2015 - Present	Te et al. (2014)
Park Falls, WI	45.945, -90.273, 0.474	July 2014 - Present	Wennberg et al. (2014a)
Fairbanks, Alaska $\flat$ $\sharp$	64.859, -147.844, 0.501	July 2015 - Present	
Railroad Valley $\flat$	38.497, -115.690, 1.4359	July 2014 - Present	
Réunion Island	-21.049, 55.285, 0.504	July 2014 - Present	De Maziere et al. (2014)
Rikubetsu, Japan	43.452, 143.700, 0.236	July 2015 - Present	Morino et al. (2014b)
Rosemount, MN $\flat$ $\sharp$	44.689, -93.027, 0.289	June 2016 - Present	
Saga, Japan	33.241, 130.288, 0.003	July 2015 - Present	Kawakami et al. (2014)
São Paulo, Brazil $\flat$ $\sharp$	-23.539, -46.634, 0.76	July 2015 - June 2016	
Shanghai, China $\flat$ $\sharp$	31.22, 121.456, 0.12	July 2015 - June 2016	
Sodankylä, Finland	67.368, 26.633, 0.18	July 2014 - Present	Kivi et al. (2014)
Tsukuba, Japan	36.051, 140.122, 0.0277	July 2014 - Present	Morino et al. (2014a)
Wollongong, Aus	-34.451, 150.855, 0.008	July 2014 - Present	Griffith et al. (2014b)



Table 2: Filters applied to the target-mode OCO-2 data from the standard OCO-2 files (i.e., not the “lite” files). Parameters for which there is only one limit are marked with a ‘—’. The units are listed where applicable. The parameter “blended\_albedo” is defined as  $2.4 \times albedo_{o2\_fph} - 1.13 \times albedo_{strong\_co2\_fph}$ .

Parameter	Lower Bound	Upper Bound	Units
surface_pressure_delta_abp	-4000	583	Pa
retrieval_surface_roughness	—	26.50	
relative_residual_mean_square_weak_co2	—	0.00250	
retrieval_zenith	—	40	°
outcome_flag	—	2	
blended_albedo	—	0.8	
h2o_ratio_idp	0.7	1.02	
co2_ratio_idp	0.995	1.025	
surface_pressure_delta_fph	-5	10	hPa
dof_co2_profile	1.8	—	
ice_aod	—	0.03	
dust_aod	0.001	0.3	
co2_grad_del	-70	70	
sulfate_aod	0.4	—	
albedo_weak_co2_fph	0.1	—	
airmass	—	3.6	
surface_type	‘Coxmunk’	‘Coxmunk’	



Table 3: Glint and nadir statistics. The median bias (OCO-2 - TCCON) and its RMS,  $R^2$  and number of daily median comparison points (N) are listed below for each TCCON station.

TCCON Site	Land Glint				Ocean Glint				Nadir			
	Bias	RMS	$R^2$	N	Bias	RMS	$R^2$	N	Bias	RMS	$R^2$	N
Eureka									0.10	1.99	0.739	3
Sodankylä									3.15	3.21	1.000	2
Białystok	0.29	1.75	0.485	4	0.53	3.57	0.181	3	0.89	1.83	0.766	14
Bremen	1.98	2.36	0.665	4	0.63	0.82	0.903	4	2.27	2.29	1.000	2
Karlsruhe	0.86	1.55	0.676	16					0.96	1.69	0.871	13
Orléans	0.29	2.08	0.369	14	-0.31	2.20	0.256	6	1.15	1.82	0.833	19
Garmisch	0.24	1.00	0.852	11					1.25	2.13	0.528	18
Park Falls	-0.50	1.44	0.919	13	0.60	1.42	0.868	13	0.27	1.55	0.800	19
Rikubetsu	2.64	2.70	0.917	3	1.02	1.10	0.932	4				
Lamont	-0.20	1.23	0.854	52					-0.13	1.18	0.865	81
Tsukuba	0.71	2.47	0.530	6	-0.56	1.79	0.698	16	1.34	3.80	0.288	9
Edwards	-0.08	1.27	0.732	35	-0.13	1.97	0.579	28	0.34	1.17	0.780	51
Pasadena	-0.26	1.20	0.856	5	-1.08	1.79	0.851	3	0.49	0.99	0.929	12
Saga	0.00	1.11	0.916	7	-1.12	1.50	0.715	20	-0.18	0.62	0.936	6
Manaus	-0.82	1.06	0.795	4								
Ascension Island					-0.09	0.63	0.736	66				
Darwin	0.10	0.96	0.676	46	0.40	0.88	0.808	60	0.58	1.04	0.711	56
Réunion Island	0.27	1.16	0.582	4	0.50	0.90	0.824	63				
Wollongong	-0.69	1.40	0.638	154	0.18	2.01	0.183	220	0.03	1.02	0.680	197
Lauder	0.03	0.98	0.731	41	0.27	0.97	0.749	136	0.69	1.30	0.599	51
<b>Total</b>	<b>-0.21</b>	<b>1.36</b>	<b>0.716</b>	<b>419</b>	<b>0.19</b>	<b>1.51</b>	<b>0.501</b>	<b>644</b>	<b>0.29</b>	<b>1.34</b>	<b>0.754</b>	<b>555</b>



Table 4: Glint, nadir and target relationships with TCCON. The slope and its uncertainty,  $R^2$  and number of daily median comparison points (N) are listed below for each OCO-2 viewing mode. The values for ocean glint data with and without the southern hemisphere winter data are included on separate rows. Note that the land glint, ocean glint and nadir slopes are computed after the global bias has been removed from the data. The global bias is determined by the target data value listed in the table.

	<b>slope</b>	$R^2$	<b>N</b>
<b>Land Glint</b>	0.9074±0.02	0.72	419
<b>Ocean Glint</b>	0.7167±0.03	0.50	644
<b>Excluding SH Winter</b>	0.8164±0.03	0.71	501
<b>Nadir</b>	0.8690±0.03	0.75	555
<b>Target</b>	0.9975±0.04	0.81	106