

# Massive nitrogen loss from the Benguela upwelling system through anaerobic ammonium oxidation

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In many oceanic regions, growth of phytoplankton is nitrogen-limited because fixation of  $N_2$  cannot make up for the removal of fixed inorganic nitrogen ( $NH_4^+$ ,  $NO_2^-$ , and  $NO_3^-$ ) by anaerobic microbial processes. Globally, 30–50% of the total nitrogen loss occurs in oxygen-minimum zones (OMZs) and is commonly attributed to denitrification (reduction of nitrate to  $N_2$  by heterotrophic bacteria). Here, we show that instead, the anammox process (the anaerobic oxidation of ammonium by nitrite to yield  $N_2$ ) is mainly responsible for nitrogen loss in the OMZ waters of one of the most productive regions of the world ocean, the Benguela upwelling system. Our *in situ* experiments indicate that nitrate is not directly converted to  $N_2$  by heterotrophic denitrification in the suboxic zone. In the Benguela system, nutrient profiles, anammox rates, abundances of anammox cells, and specific biomarker lipids indicate that anammox bacteria are responsible for massive losses of fixed nitrogen. We have identified and directly linked anammox bacteria to the removal of fixed inorganic nitrogen in the OMZ waters of an open-ocean setting. We hypothesize that anammox could also be responsible for substantial nitrogen loss from other OMZ waters of the ocean.

anammox | denitrification | oceanic nitrogen cycle | oxygen-minimum zone

Large amounts of dissolved inorganic carbon, fixed inorganic nitrogen ( $NH_4^+$ ,  $NO_2^-$ , and  $NO_3^-$ ), and phosphate in ocean-surface waters are consumed by algae and cyanobacteria during photosynthesis. The N:C and P:C ratios of the organic products are similar throughout the marine realm (1, 2). On average, the C:N:P molar ratio of this phytoplanktonic biomass is 106:16:1 (3). Carbon dioxide, ammonium, and phosphate are released in the same ratio, the so-called Redfield ratio, when this phytoplankton-derived organic matter is remineralized (3). Under oxic conditions, nitrifying organisms (mainly chemoautotrophic bacteria) will consume  $O_2$  and oxidize ammonium through nitrite to nitrate. The nitrate can be reutilized to produce new phytoplanktonic biomass. Alternatively, in oxygen-deficient marine environments, predominately facultative anaerobic prokaryotes can use nitrate as an electron acceptor while oxidizing organic matter. During this process (i.e., heterotrophic denitrification) nitrate is reduced to  $N_2$ . Some ammonium is also released and, in the absence of oxygen, is expected to remain as ammonium (4).

For some time, however, oceanographers have known that far less ammonium accumulates in anoxic fjords and basins than would be expected from the stoichiometry of heterotrophic denitrification (5, 6). To explain this shortfall, it was suggested that microorganisms can combine ammonium and nitrate to yield  $N_2$  (5). Direct evidence for the anaerobic oxidation of ammonium was found in wastewater bioreactors, where so-called “anammox” bacteria belonging to the order *Planctomycetales* directly oxidize ammonium to  $N_2$  with nitrite as the electron acceptor (7–9). Nitrite is an intermediate in both heterotrophic denitrification (7) and aerobic ammonium oxidation (10, 11). Experiments with enrichment cultures and purified cells indicate

that anammox bacteria can also reduce nitrate to nitrite while oxidizing short-chain fatty acids (e.g., acetate and propionate) (12).

The anaerobic oxidation of ammonium with nitrite to  $N_2$  (anammox) was recently recognized as a major sink for fixed inorganic nitrogen in coastal sediments and the anoxic waters of basins isolated from oxygenated deep circulation (4, 13–19). Globally, 30–50% of the total nitrogen loss occurs in oxygen-minimum zones (OMZs) and is commonly attributed to heterotrophic denitrification (20, 21). There has been, until now, no published direct evidence for anammox in OMZs. However, the extremely low concentration of ammonium could indicate that anammox bacteria also play an important role in the nitrogen removal from OMZ waters (e.g., refs. 4, 16, and 18).

Upwelling of nutrient-rich South Atlantic midwaters in the Benguela current system along the southwest African continental margin sustains some of the highest rates of primary production in the ocean (Fig. 1a) (22, 23). Although the upwelling water is generally well oxygenated ( $>200 \mu M O_2$ ), bottom waters become severely oxygen-depleted ( $<10 \mu M O_2$ ; Fig. 1b) over large areas of the southwest African shelf. This condition results from the consumption of  $O_2$  during the decomposition of settling algal biomass (23). A strong N deficit (i.e., a decrease in the ratio of fixed inorganic N to P) (24) in the bottom waters (Fig. 1b) has been attributed to denitrification (23, 24). Water in the Benguela OMZ is exchanged rapidly, and concentrations of oxygen vary significantly (23). One of the aims of this study was to investigate whether anammox bacteria, with their slow growth rate and anaerobic physiology, are able to thrive under such highly dynamic conditions.

## Methods

Salinity, temperature, density, turbidity, and oxygen profiles were obtained by a conductivity–temperature–depth (CTD) system equipped with an oxygen sensor (Sea-Bird Electronics, Bellevue, WA). The oxygen sensor was calibrated by manual Winkler titration (in duplicate) of 100-ml water samples collected from four depths throughout the water column for each station.

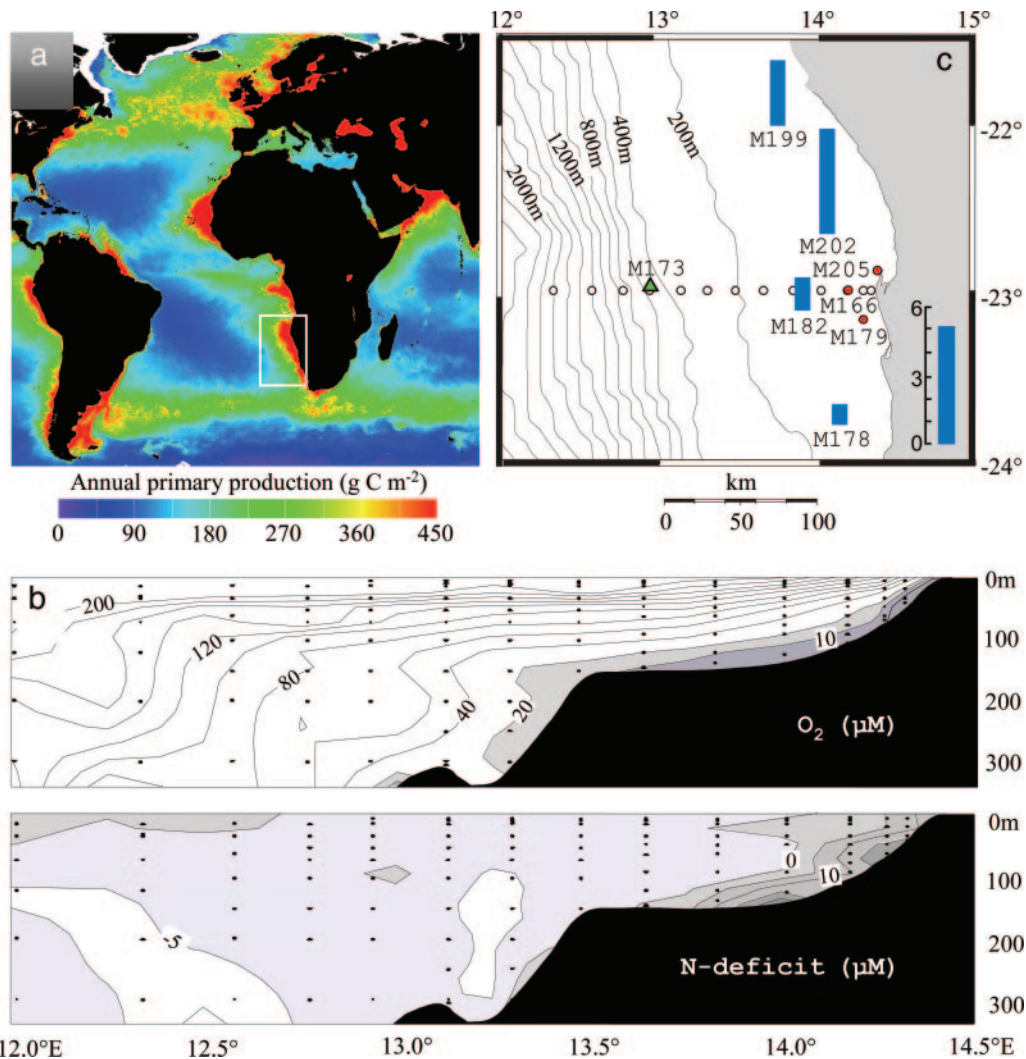
**Nutrient Analyses.** Water samples for nutrient analyses were obtained by a pump-CTD system or go-flow bottles (Hydrobios, Loudeac, France). Nitrate, nitrite, ammonium, and phosphate concentrations (detection limits 0.1, 0.03, 0.3, and  $0.1 \mu M$ , respectively) were determined on board with an autoanalyzer (TRAACS 800, Bran & Lubbe, Hamburg, Germany) immediately after sampling. The fixed-inorganic-nitrogen deficit was calculated from nutrient concentrations by using the equation  $N \text{ deficit} = 16 \times [\text{phosphate}] - ([\text{nitrate}] + [\text{nitrite}] + [\text{ammonium}])$ , in which 16 is the Redfield ratio of fixed inorganic nitrogen to phosphate (24).

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Abbreviations: CTD, conductivity–temperature–depth; OMZ, oxygen-minimum zone.

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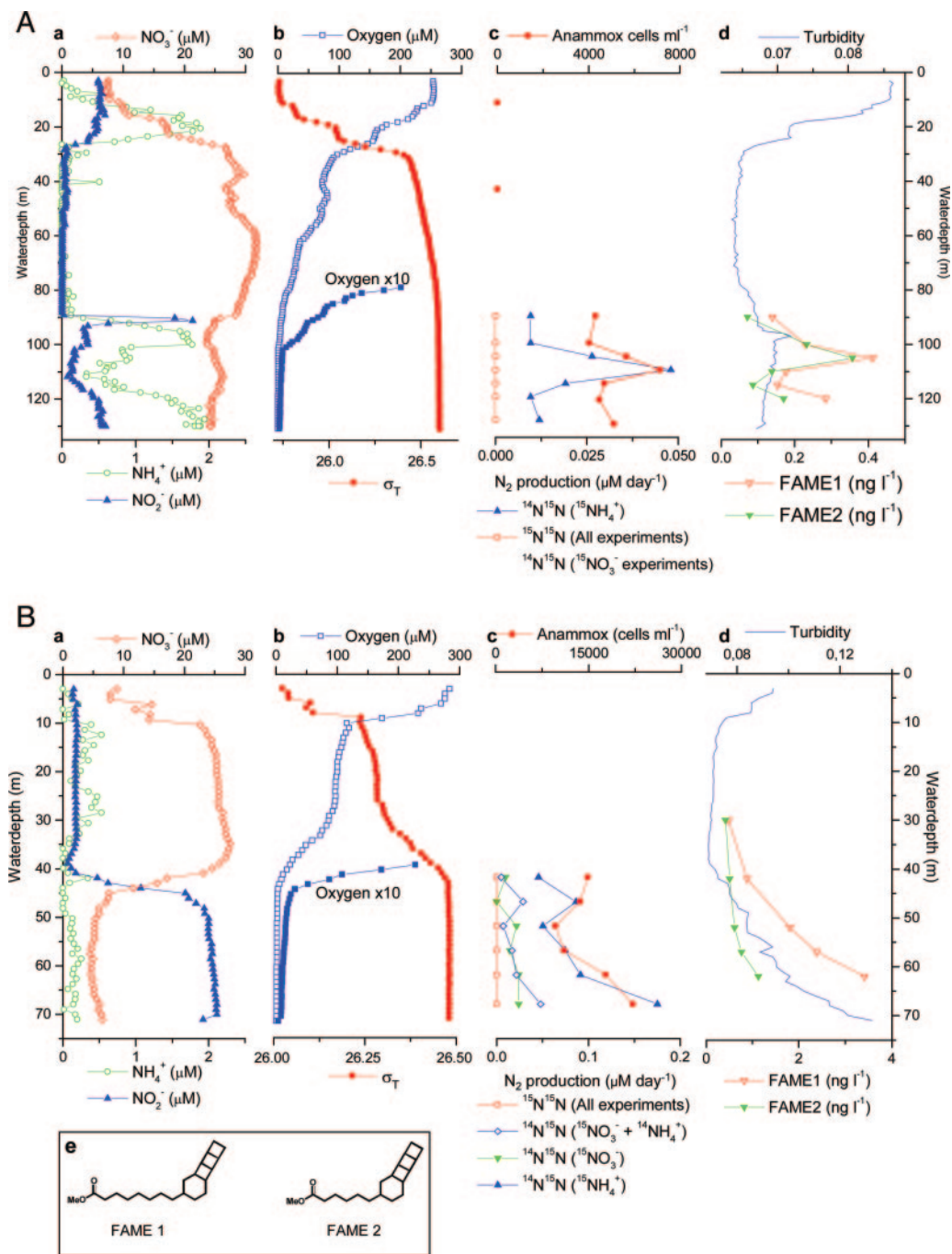
**Fig. 1.** The Benguela system. (a) Distribution of annual primary production (source: <http://marine.rutgers.edu/opp/swf/Production/results/all2.swf.html>). The white box indicates the extent of the Benguela upwelling system. (b) Vertical transect showing the lateral extension and fixed-inorganic-nitrogen deficit in the OMZ off Namibia at 23° south (see *Nutrient Analyses in Methods*). (c) Sites and nitrogen losses. The open circles represent sites used to construct the lateral transect in b. The blue bars represent the depth-integrated nitrogen loss ( $\text{mmol}\cdot\text{m}^{-2}\cdot\text{d}^{-1}$ ) through anammox determined from anaerobic  $^{15}\text{N}$  incubations of water collected from six or seven depths throughout the suboxic zone at sites M178, M182, M199, and M202. For sites M179 and M205 (red circles) anaerobic  $^{15}\text{N}$  incubations indicate anammox activity, but rates were not measured. The integrated nitrogen loss for site M166 (red circle) is not shown because anaerobic  $^{15}\text{N}$  incubations were performed for only one depth (107 m, Fig. 4). Anammox activity was undetectable at site M173 (green triangle), where oxygen concentrations in the bottom waters exceed 20  $\mu\text{M}$ .

**$^{15}\text{N}$  Incubations and Analysis.** For the  $^{15}\text{N}$  incubations, we slightly modified the method previously published by Dalsgaard *et al.* (16). Briefly, 250 ml of Namibian shelf water collected from specific water depths with the pump-CTD were flushed with helium for 15 min after the addition of 5  $\mu\text{mol}$  of  $\text{Na}^{15}\text{NO}_3$ , 2.5  $\mu\text{mol}$  of  $^{15}\text{NH}_4\text{Cl}$ , or 5  $\mu\text{mol}$  of  $\text{Na}^{15}\text{NO}_3$  and 2.5  $\mu\text{mol}$  of  $^{14}\text{NH}_4\text{Cl}$  (reagents from Campro Scientific, Berlin). The  $^{15}\text{N}$  incubations were started immediately after sampling. The water was transferred into 12-ml Exetainers (Labco, High Wycombe, Buckinghamshire, U.K.) and incubated for up to 48 h at *in situ* temperatures. Samples were taken after 0, 12, 24, and 48 h by removing 1 ml of water while replacing it with helium. Mercuric chloride was added to the samples to stop biological activity, and the samples were stored at 4°C until analysis.  $^{15}\text{N}^{14}\text{N}:^{14}\text{N}^{14}\text{N}$  and  $^{15}\text{N}^{15}\text{N}:^{14}\text{N}^{14}\text{N}$  ratios were determined by GC isotope ratio MS.

**Particulate Organic Carbon and Lipid Analysis.** Particulate matter was collected from specific water depths by using *in situ* pumps

(Challenger Oceanic, Haslemere, Surrey, U.K.) to filter large volumes of water ( $\approx 500$  liters per sample) through glass-fiber filters (GFF) (Whatman, Maidstone, U.K.) precombusted at 450°C (GFF nominal pore size, 0.7  $\mu\text{m}$ ). Contents of particulate organic carbon were determined by using a carbon and nitrogen (CN) analyzer (Carlo Erba, Milan). The GFF filters were Soxhlet-extracted for 24 h to obtain the total lipid extracts. Aliquots of the total extracts were saponified after the addition of an internal standard and separated into fatty-acid and neutral-lipid fractions. The fatty-acid fractions were methylated and analyzed by GC-MS for identification and quantification of ladderane lipids. Repeated concentration measurements agreed within  $\pm 10\%$ . Because filters with a pore size of 0.7  $\mu\text{m}$  may undersample anammox cells, the calculated ladderane-lipid concentrations represent minimum values.

**Phylogenetic Analysis.** Extraction, isolation, and cloning of DNA followed procedures described in ref. 25. The software package ARB (26) was used for phylogenetic analyses.



**Fig. 2.** Chemical zonation and distribution of anammox indicators at sites M182 on March 21, 2003 (A) and M202 on March 27, 2003 (B). (a) Concentrations of fixed inorganic nitrogen species. (b) Water density ( $\sigma_T$ , the density of seawater in  $\text{kg}\cdot\text{m}^{-3}$  – 1,000) and oxygen concentrations (notice the 10-fold expanded  $\text{O}_2$  gradient). (c) Anammox cells per milliliter and rates of production of  $\text{N}_2$ . The isotopic species pertain to the incubations indicated. At M182, production of  $^{15}\text{N}^{15}\text{N}$  in all incubations and of  $^{14}\text{N}^{15}\text{N}$  with added  $^{15}\text{NO}_3^-$  and  $^{14}\text{NH}_4^+$  or with added  $^{15}\text{NO}_3^-$  was undetectable. Accordingly, all are represented by a single line. (d) Turbidity and depth distributions of membrane lipids specific for anammox bacteria. FAME, fatty acid methyl ester. (e) Molecular structures of the two ladderane fatty acid methyl esters represented in d.

**FISH.** Particulate organic matter for FISH analyses was collected from specific water depths by filtration of 30 ml of water through polycarbonate filters (0.2- $\mu\text{m}$  pore size). The filter material was fixed with paraformaldehyde and hybridized with fluorescently labeled oligonucleotide probes specific to *Planctomycetes* (27) and to anammox bacteria (25, 28). Total cell counts were based on DAPI staining. The average number of anammox bacteria was determined by epifluorescence microscopy (28).

## Results and Discussion

During an R/V *Meteor* cruise in March and April, 2003, we investigated the role of denitrification and anammox in Namibian shelf waters by combining microbiological and biogeochemical techniques. As observed in ref. 23, concentrations of nitrate maximized at 25–30  $\mu\text{M}$  in the midwaters of the Namibian shelf (Fig. 2) and, because of uptake of nitrate during phytoplankton growth, were substantially lower in the surface

waters. Vertical distributions of nitrate, nitrite, and ammonium differ substantially between sites M182 and M202 (Fig. 2), reflecting the heterogeneous and dynamic nature of the Benguela upwelling system (23). However, at both sites, concentrations of nitrate drop at the base of the oxic zone ( $<10 \mu\text{M O}_2$ ; Fig. 2). This decrease has been attributed to the conversion of nitrate to  $\text{N}_2$  by denitrifying bacteria (23, 24). Nitrite is an intermediate in this process, and nitrite maxima were associated with the decrease in nitrate concentrations at both sites (Fig. 2). The N:P ratios for dissolved inorganic nutrients are well below Redfield values and indicate extensive loss of nitrogen from the oxygen-deficient waters at sites M182 and M202 (see Tables 1 and 2, which are published as supporting information on the PNAS web site). Ammonium should have accumulated in these waters if the nitrogen loss were due solely to heterotrophic denitrification (5, 6). The low ammonium concentrations (below detection limit) in the suboxic zone at site M202 (Fig. 2B) could indicate that anammox bacteria play an important role in nitrogen removal from the waters of the Benguela OMZ (4, 16, 18).

The N deficit (Fig. 1b) requires that large amounts of fixed inorganic nitrogen are being removed from the Namibian suboxic waters. To elucidate the microbial process responsible for this loss of N, we incubated water samples from various depths after the addition of [ $^{15}\text{N}$ ]nitrate or [ $^{15}\text{N}$ ]ammonium. Denitrification combines two nitrate ions to form one molecule of  $\text{N}_2$ . Because all water samples contained a natural abundance of [ $^{14}\text{N}$ ]nitrate in addition to the added [ $^{15}\text{N}$ ]nitrate, denitrification would produce  $^{14}\text{N}^{14}\text{N}$ ,  $^{14}\text{N}^{15}\text{N}$ , and  $^{15}\text{N}^{15}\text{N}$  through random isotope pairing (13, 14, 29, 30). Even though the water samples were incubated with 45–80%  $^{15}\text{N}$ -labeled nitrate, production of  $^{15}\text{N}^{15}\text{N}$  was undetectable in all 0-, 12-, and 24-h incubations and in 96 of 100 48-h incubations (Figs. 2 and 3). The potential to respire nitrate is widespread among bacteria and archaea (31). The production of  $^{15}\text{N}^{15}\text{N}$  in 4 of 100 48-h incubations (Fig. 4) indicates that facultatively anaerobic heterotrophs that can use nitrate in the absence of oxygen are present in the Namibian shelf waters but does not indicate that nitrate is directly converted to  $\text{N}_2$  by denitrifying microorganisms in the suboxic shelf water.

Denitrification requires electron donors (e.g., organic matter, sulfide). No sulfide was detected in Namibian shelf waters during the 2003 R/V *Meteor* cruise. However, concentrations of particulate organic carbon were  $30\text{--}95 \mu\text{g}\cdot\text{liter}^{-1}$  at M182 and  $70\text{--}300 \mu\text{g}\cdot\text{liter}^{-1}$  at M202, indicating that organic matter is not limiting. Apparently, the expression or activity of denitrifying enzymes was suppressed in the shelf waters. We attribute this suppression to regular incursions of oxygen into the suboxic zone (23). Although some bacterial strains can denitrify under aerobic conditions in the laboratory (31), denitrifying enzymes are generally repressed or inhibited by free oxygen, and induction of full denitrifying capacity after reestablishment of anaerobic conditions can take at least 20 h (31, 32).

In contrast, anammox-enrichment cultures that have been exposed to oxygen resume activity immediately after the reestablishment of anaerobic conditions (33). Water samples from site M202 that were anaerobically incubated after the addition of [ $^{15}\text{N}$ ]nitrate and [ $^{14}\text{N}$ ]ammonium produced significant amounts of  $^{14}\text{N}^{15}\text{N}$  (Figs. 2B and 3), indicating either that the reduction of nitrate to nitrite by heterotrophic denitrifiers was coupled to anammox (7, 14) or that anammox bacteria could be using nitrate as an electron acceptor concurrently with nitrite. The latter process, in which anammox bacteria reduce nitrate to nitrite while oxidizing short-chain fatty acids, was recently discovered in anammox bioreactors (12). The anaerobic incubation of water samples after the addition of [ $^{15}\text{N}$ ]ammonium resulted in substantial  $^{14}\text{N}^{15}\text{N}$  production in samples from the suboxic zone (Figs. 2 and 3), providing direct evidence for removal of fixed inorganic nitrogen by anammox

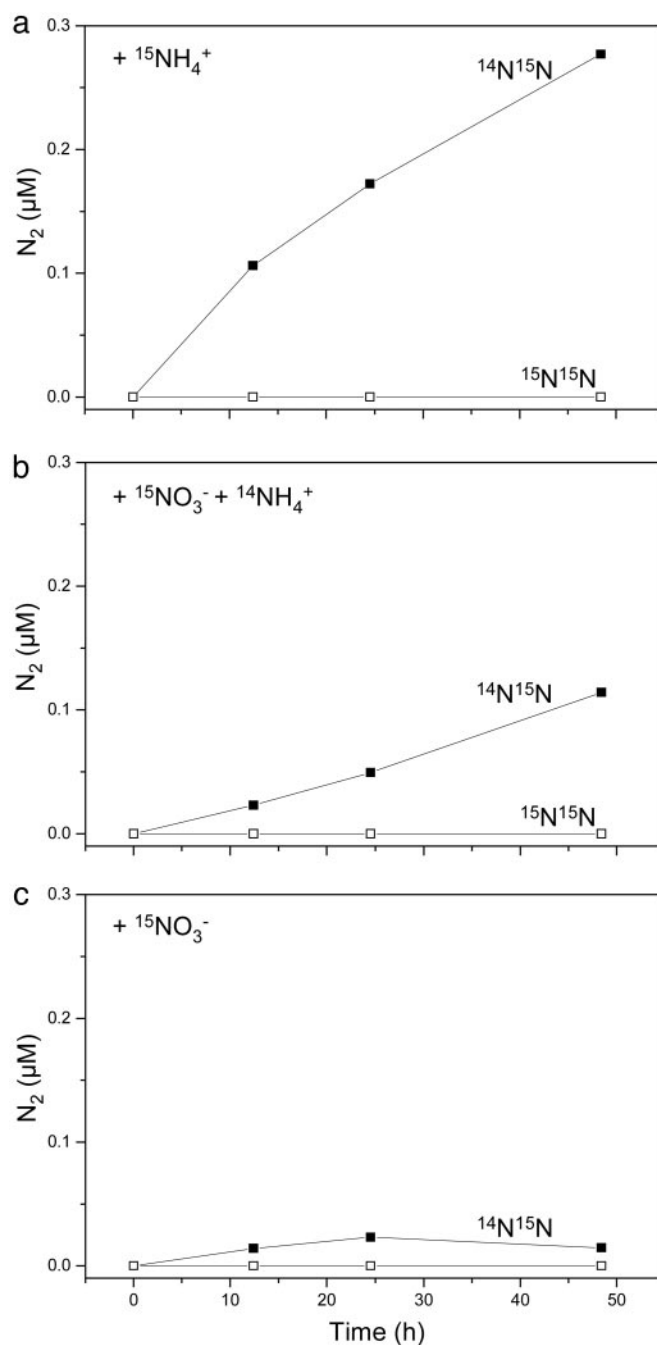
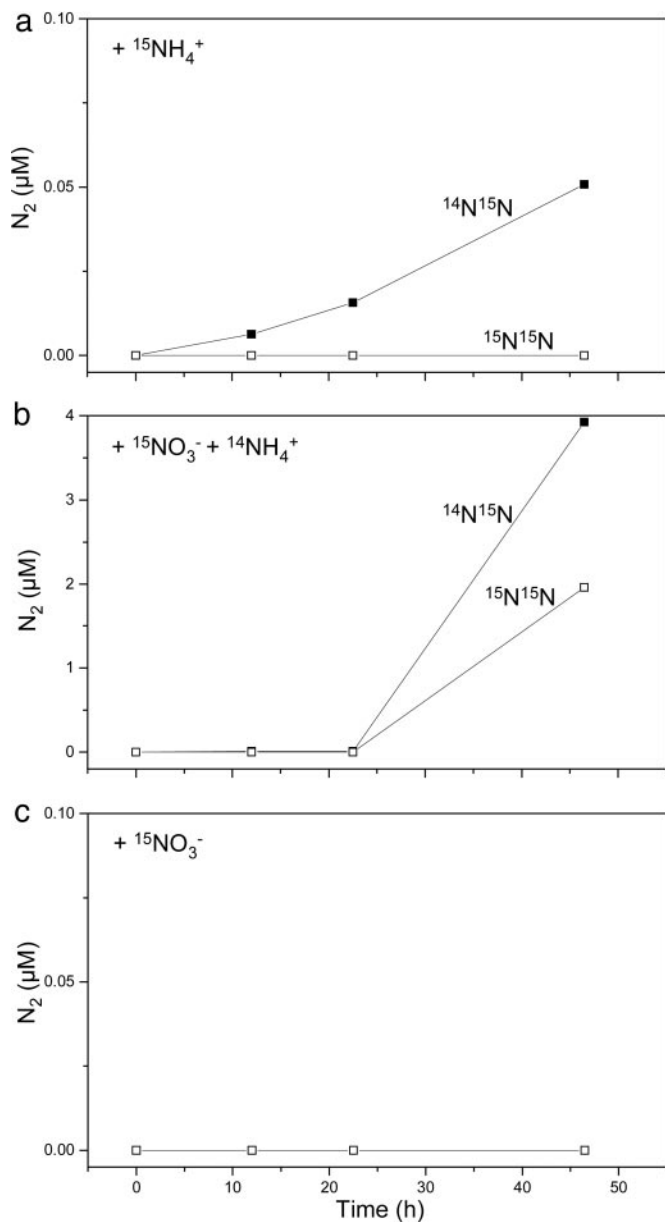


Fig. 3. Changes in concentrations of isotopically labeled  $\text{N}_2$  species vs. time during incubations of samples from a depth of 68 m at site M202.

bacteria. The time-dependent data from M182 (data not shown) and M202 (Fig. 3) show rather linear accumulations of  $^{14}\text{N}^{15}\text{N}$  for both the  $^{15}\text{NO}_3^-$  and  $^{15}\text{NH}_4^+$  additions. This lack of a lag time indicates that the anammox bacteria are active from time 0 and are likely active *in situ*.

Production of  $\text{N}_2$  by anammox at site M182 maximizes in the suboxic zone between 100- and 115-m depth, where ammonium and nitrite disappear (Fig. 2A). Production of  $^{14}\text{N}^{15}\text{N}$  was below the detection limit for sea water incubated for up to 48 h with [ $^{15}\text{N}$ ]nitrate, indicating that aerobic oxidation of ammonium, rather than nitrate reduction, is the source of nitrite for anammox at site M182 ( $\approx 110\text{-m}$  depth; Fig. 2A). At site M202, the production of  $\text{N}_2$  by anammox increases with depth, with

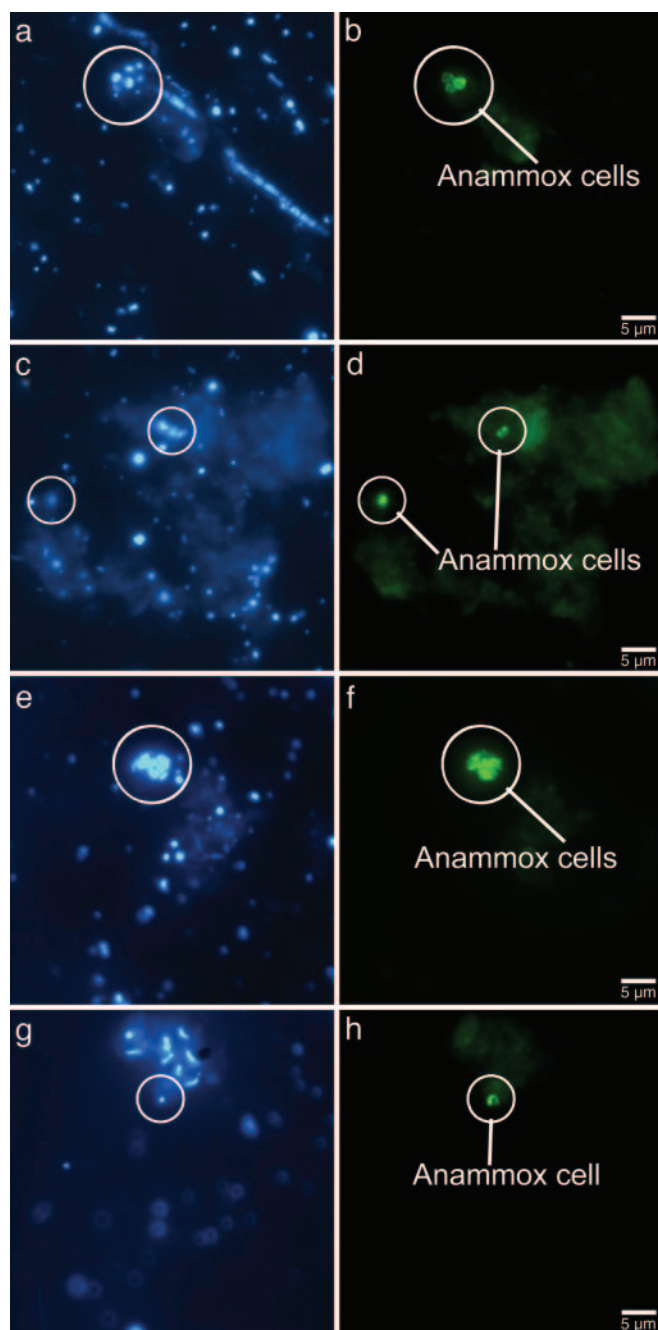


**Fig. 4.** Changes in concentrations of isotopically labeled  $N_2$  species vs. time during incubations of samples from a depth of 107 m at site M166.

the highest rates near the sea floor. Concentrations of nitrite remain constant, indicating transient conditions (Fig. 2B).

Specific biomarkers, so-called ladderane lipids, were used to trace anammox bacteria in particulate organic matter collected from various depths across the suboxic zone. Ladderane lipids (17, 34, 35) are the main building blocks of a unique bacterial membrane that surrounds the “anammoxosome,” a special compartment in the anammox cell where  $N_2$  is produced. Two different ladderane lipids were found in the fatty-acid fraction obtained after saponification of the total lipid extract. Their distribution vs. depth was similar to that of anammox activity (Fig. 2).

In addition, DNA was isolated from various depths through the suboxic zone. A clone library was generated after the 16S ribosomal RNA gene was amplified with primers specific for *Planctomycetes* (25). Phylogenetic analysis of the resulting 16S rRNA sequences showed that *Planctomycetes* closely related



**Fig. 5.** *In situ* identification of aggregate-associated anammox cells (encircled) from station M202 (62 m). Single xy images of the same section were combined. (a, c, e, and g) Micrographs show aggregates stained with DAPI. (b, d, f, and h) Micrographs show results of hybridization with an oligonucleotide probe (AmxB5820) specific for anammox bacteria.

(98% sequence similarity) to known anammox bacteria, *Candidatus “Scalindua sorokinii”* and *Candidatus “Scalindua brodae”* (17, 25), were present in the Namibian shelf waters.

FISH with oligonucleotide probes specific for anammox bacteria (25) was used to quantify anammox bacteria in the Namibian shelf waters. Up to 22,000 anammox cells per ml ( $\approx 1\%$  of the total number of cells counted after DAPI staining) were detected in the suboxic zone. The depth distribution was similar to that of the anammox activity. No anammox cells were found in surface waters, where concentrations of  $O_2$  exceed  $20 \mu M$  (Fig. 2). A strong linear correlation between cell numbers and

activity ( $R^2 > 0.8$ ) indicates that anammox bacteria are responsible for the anaerobic ammonium oxidation in the suboxic Namibian shelf waters. The cell-specific anammox activity calculated ( $\approx 4.5$  fmol of ammonium per cell per day) for the Namibian shelf water is comparable with the rates estimated for the suboxic zone of the Black Sea (3–4 fmol of ammonium per cell per day) (17) and is well within the range found for anammox bacteria in laboratory bioreactors (2–20 fmol of ammonium per cell per day) (9).

Intriguingly, chemical profiles,  $^{15}\text{N}$ -labeling experiments, ladder-lipid distributions, and FISH analysis indicate that active anammox bacteria are abundant at oxygen concentrations up to  $9\ \mu\text{M}$  (Fig. 2). The abundance of anammox bacteria in the presence of free oxygen is surprising because the anammox metabolism is inhibited (33) by oxygen concentrations as low as  $1\ \mu\text{M}$ . The anammox bacteria present at higher oxygen concentrations could be dormant. Alternatively, marine snow aggregates are abundant in the Namibian shelf water and could provide the anammox bacteria with anaerobic microenvironments at low ambient-oxygen concentrations ( $< 25\ \mu\text{M}$ ) (36). The abundance of aggregate-associated anammox cells visualized by FISH provides evidence for a particle association of the anammox bacteria at site M202 (Fig. 5). This finding is supported by the strong covariation among anammox rates, lipid profiles, anammox cell numbers, and turbidity as a measure of particle abundance (Fig. 2B). Whether anammox bacteria are particle-associated or not, our results clearly show that these bacteria are well adapted to the dynamic conditions encountered in a coastal upwelling region.

Labeling experiments with  $^{15}\text{N}$ ammonium at a number of sampling sites showed that anammox occurred throughout the lower 20–30 m of the water column over wide areas of the Namibian shelf (Fig. 1c). Labeling experiments with  $^{15}\text{N}$ nitrate at sites M178, M199 (data not shown), M182 (Fig. 2A), and M202 (Fig. 2B) did not provide evidence for significant production of  $\text{N}_2$  by heterotrophic denitrifiers in Namibian shelf waters. Instead, our results indicate that fixed nitrogen is lost through

anammox coupled to (i) reduction of nitrate to nitrite by heterotrophic denitrifiers or anammox bacteria and (ii) aerobic ammonium oxidation. Data from the Chilean OMZ show a similar dominance of anammox over heterotrophic denitrification (B. Thamdrup, personal communication).

Depth-integrated rates indicated that 1–5 mmol of fixed inorganic nitrogen per square meter per day was lost from the suboxic waters of the Benguela upwelling system due to anammox (Fig. 1c). If we assume that the main area ( $\approx 100,000\ \text{km}^2$ ) of suboxic shelf water extends from  $28^\circ$  to  $18^\circ$  south (23),  $1.4 \pm 1\ \text{Tg}$  of fixed nitrogen per year might be lost through anammox from the Benguela system. The total estimated nitrogen loss from the OMZ waters of the world ocean ( $80\text{--}150\ \text{Tg}\cdot\text{yr}^{-1}$ ) is based on nutrient measurements and has been fully attributed to heterotrophic denitrification (20, 21, 37) because, before the discovery of anammox bacteria, there was no other process known that could transform fixed inorganic nitrogen into  $\text{N}_2$ . In fact, to the best of our knowledge, there is, so far, no published evidence from  $^{15}\text{N}$ -labeling experiments that nitrate is directly converted to  $\text{N}_2$  by heterotrophic denitrifiers in the OMZ waters of the ocean. Our combined results show that anammox bacteria are responsible for massive losses of fixed nitrogen as gaseous  $\text{N}_2$  from the Benguela OMZ water. The possibility that anammox is also a dominant process for nitrogen removal in other OMZ waters of the ocean should now be explored.

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