

Supplementary Information for

Regional Wind Variability Modulates the Southern Ocean Carbon Sink

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- Figs. S1 to S7
- References for SI reference citations

1 **Supplementary Information Text**

2

3 **S1. The effect of Sea Level Pressure on the Southern Annular Mode (SAM)**

4 Previous studies have shown that the SAM has a zonal asymmetry that affects the mixed-
5 layer depth and temperature^{1,2}. Fig. S1 highlights the zonal asymmetry of the SAM, indicating
6 that the negative correlation between the SAM and sea level pressure reaches further north in the
7 eastern Pacific compared to the rest of the study region.

8

9 **S2. The available carbon dioxide (CO₂) observations and robustness of the interpolated**
10 **data**

11 The Southern Ocean is historically an under-sampled region, due to the harsh conditions
12 and the relative remoteness. The available data of the partial pressure of CO₂ (pCO₂) in the
13 Southern Ocean come mainly from shipboard measurements, which have steadily increased in
14 number in the past several decades, with the Drake Passage and the Tasman Sea being the best-
15 sampled region within the Southern Ocean since the 2000s (Fig. S2A)^{3,4}. Despite the substantial
16 increase in observational pCO₂ data in the Southern Ocean, the spatio-temporal distribution
17 remains sparse compared to other regions³. The robustness of the neural-network interpolated
18 pCO₂ data until December 2011 has been demonstrated in previous studies⁵⁻⁷. Here, we show
19 the robustness of the method for the most recent period.

20 Averaged over the most recent time period (2012 through 2016) the observations (Fig.
21 S2B) are relatively well represented in the interpolated pCO₂ (Fig. S2C). Although some regional
22 biases are present (Fig. S2D), they mostly cancel out when averaged over the study region (1.4
23 μatm). The standard deviation at each grid point is shown in Fig. S2E, which add up to 5.6 μatm .
24 The Antarctic coastal areas display the largest standard deviation; however, our study mainly
25 focuses on observations north of 65°S, due to the data availability of the temperature and salinity.

26

27 **S3. The effect of the El Niño Southern Oscillation (ENSO) on the Southern Ocean air-sea**
28 **CO₂ flux**

29 We investigate the effect of the ENSO on the Southern Ocean carbon uptake, similarly
30 as we did for the SAM (Fig. S3). We find the regional effect of the ENSO considerably smaller
31 than the regional effect by the SAM. However, similar as the SAM, integrated over each of the
32 three sectors, and over the whole Southern Ocean, the net effect of the ENSO on the Southern
33 Ocean carbon uptake is ~zero.

34

35 **S4. The regional relationship between sea surface temperature (SST) and the CO₂ flux**

36 The air-sea CO₂ flux partially depends on SST, as CO₂ dissolves better in colder water
37 However, other factors including biological production and vertical circulation also affect the
38 oceanic CO₂ uptake. Here, we demonstrate the relationship between the SST and the air-sea CO₂
39 flux in the different sectors and interfrontal zones from 2004 through 2016 (Fig. S4).

40 In all three sectors of the Antarctic Zone, (AAZ, from 65°S to ~55°S), the SST ranges
41 approximately from -2 to 5°C, and the carbon flux varies approximately from -5 to 3 mol m⁻²
42 year⁻¹. There is a slight trend where warmer surface waters tend to coincide with more uptake in
43 the Atlantic and Indian sectors of the AAZ, indicating that solubility is not the main driver here.
44 Concurrently, the Pacific sector of the AAZ does not show a considerable trend.

45 In the Polar Frontal Zone (PFZ, from ~55°S to ~40°S), the SST ranges approximately
46 from 0 to 17°C, and the carbon flux varies from approximately -4 to 2 mol m⁻² year⁻¹. In the
47 Atlantic and Indian sectors of the PFZ, warmer surface waters tend to coincide with more uptake.
48 However, the Pacific sector is a lot more variable and the colder surface waters have a similar

49 trend as the Atlantic and Indian sectors, but the warmer waters in this region have a reversal of
50 this trend.

51 In the Subtropical Zone (STZ, from $\sim 40^{\circ}\text{S}$ to 30°S), the SST ranges approximately from
52 10 to 25°C , and the carbon flux varies approximately from -5 to $1 \text{ mol m}^{-2} \text{ year}^{-1}$. In all three
53 sectors of the STZ, there is a trend of colder surface waters tending to coincide with more uptake.

54

55 **S5. Mean sea surface properties and air-sea CO_2 flux of the sectors**

56 Here, we show the zonal mean sea SST, sea surface salinity (SSS) and the air-sea CO_2
57 flux in the Southern Ocean sectors to provide context (Fig. S5); the anomalies are shown in S6.

58

59 **S6. Sea surface property anomalies and the air-sea CO_2 flux anomalies of the sectors**

60 We analyse the sea-air CO_2 flux anomalies in comparison with the physical sea surface
61 properties of the Southern Ocean (Fig. S6). Generally, we find that warmer and saltier surface
62 waters tend to coincide with a stronger carbon sink. Therefore, it is evident that solubility is not
63 the dominant mechanism driving the sink variability, but rather reflects the reduction in vertical
64 mixing that usually brings cool but carbon-rich deep water to the surface. However, strong
65 differences exist between both the different sectors and the interfrontal zones (see also S4).

66 Similar to the sea-air CO_2 flux, the SST and SSS anomalies are strongest in the Atlantic
67 sector (Fig. S6A-C). Here, the temporal evolution is dominated by a warm anomaly in 2011,
68 north of the Polar Front (PF), which coincides with saltier surface waters and a stronger carbon
69 sink. A similar, but weaker anomalous period can be observed around 2009, north of the PF.
70 After this event, the Atlantic sector becomes cooler and fresher, while the carbon sink becomes
71 weaker again, especially north of the PF.

72 In the Pacific sector, the signal varies on shorter time scales compared to the Atlantic
73 sector in all three variables (Fig. S6D-F), which strongly coincides with the ENSO (see S3),
74 suggesting the influence of remote modes of variability on the Southern Ocean CO₂ sink. While
75 throughout the time period, warm SST anomalies usually coincide with saltier phases and vice
76 versa, the carbon sink appears disconnected from this pattern. Warmer and saltier waters
77 coincide with less carbon uptake in the PFZ from 2004 through 2011, which was previously⁷
78 considered as evidence that the CO₂ flux trends in the Pacific sector are solubility dominated⁸.
79 However, subsequently colder and fresher waters coincide with less carbon uptake in the PFZ
80 after 2011, challenging this view.

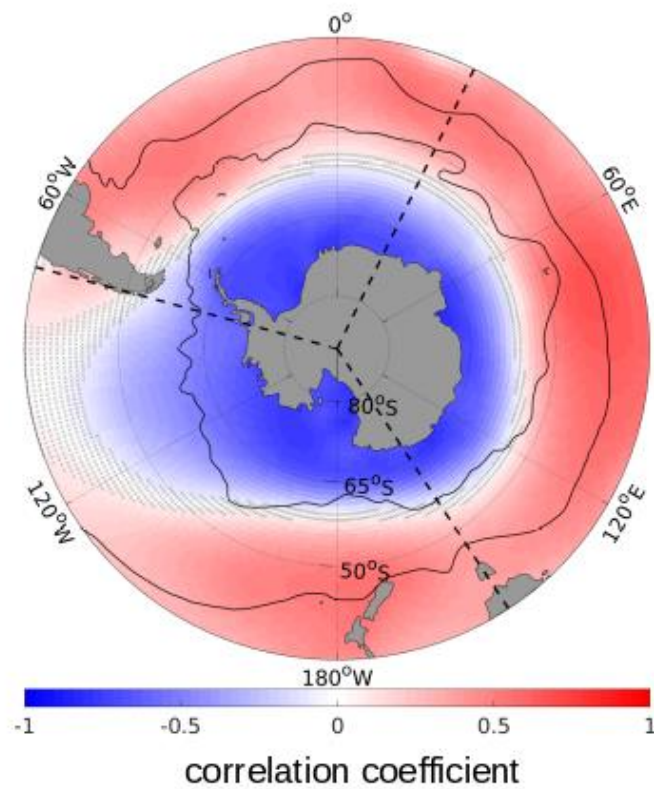
81 The Indian sector behaves similarly to the Atlantic sector in terms of SST, similar to the
82 Pacific sector in terms of SSS, and like a mixture between the Atlantic and Pacific sectors in
83 terms of its CO₂-flux (Fig. S6G-I). For example, the period of much higher SST around 2011
84 north of the PF that was observed in the Atlantic, is also present in the Indian sector, albeit with
85 less intensity. In the following years, the surface waters in the STZ of both the Pacific and the
86 Indian sectors are saltier than the mean. This trend moves further south in ~2016.

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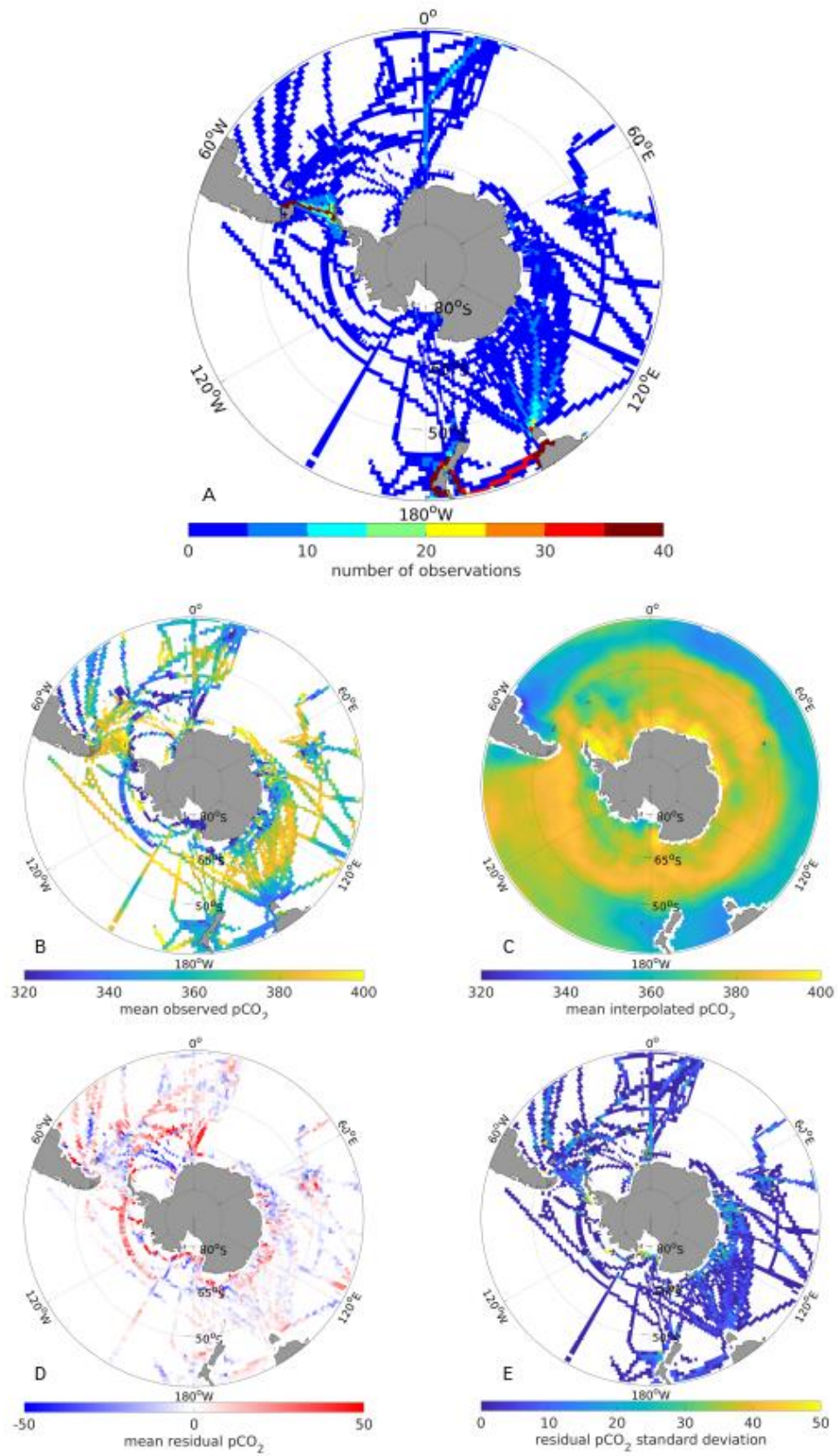
88 **S7. Trends of the pCO₂, its components, and drivers in the reinvigoration period**

89 Here, we show the trends of the reinvigoration period to put our findings on the trends
90 during the most recent period in context. As discussed in Landschützer et al. [2015]⁸, in the
91 reinvigoration period, the $\Delta p\text{CO}_2$ decreased in the Southern Ocean (Fig. S7A), resulting in
92 enhanced CO₂ uptake by the ocean. The authors demonstrated that in this period the westerly
93 winds were stronger in the Pacific and weaker in the Atlantic due to a dipole in sea level pressure
94 (Fig. S7D). This change in surface wind patterns is thought to have caused enhanced
95 downwelling and warmer surface waters in the Atlantic sector. The non-thermal component (Fig.
96 S7C) dominated over the thermal component (Fig. S7B), resulting in an overall decrease in

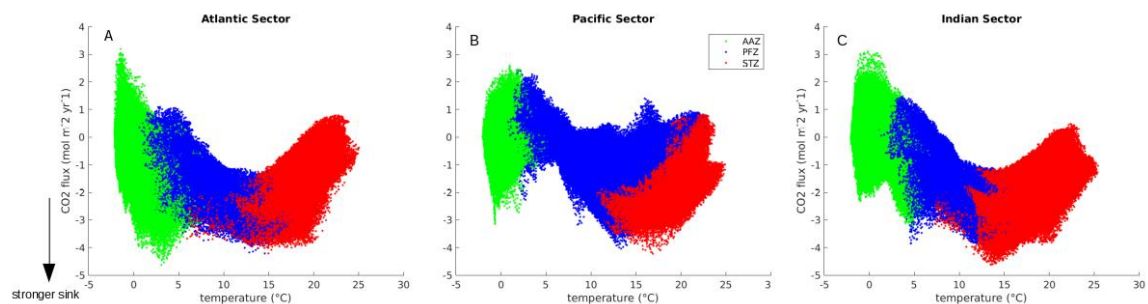
97 $\Delta p\text{CO}_2$ (Fig. S7A). Concurrently, the stronger westerlies in the Pacific sector caused enhanced
98 upwelling and colder surface waters. Here, the thermal component (Fig. S7B) dominated over
99 the non-thermal component (Fig. S7C), resulting in an enhanced CO_2 uptake by the ocean.



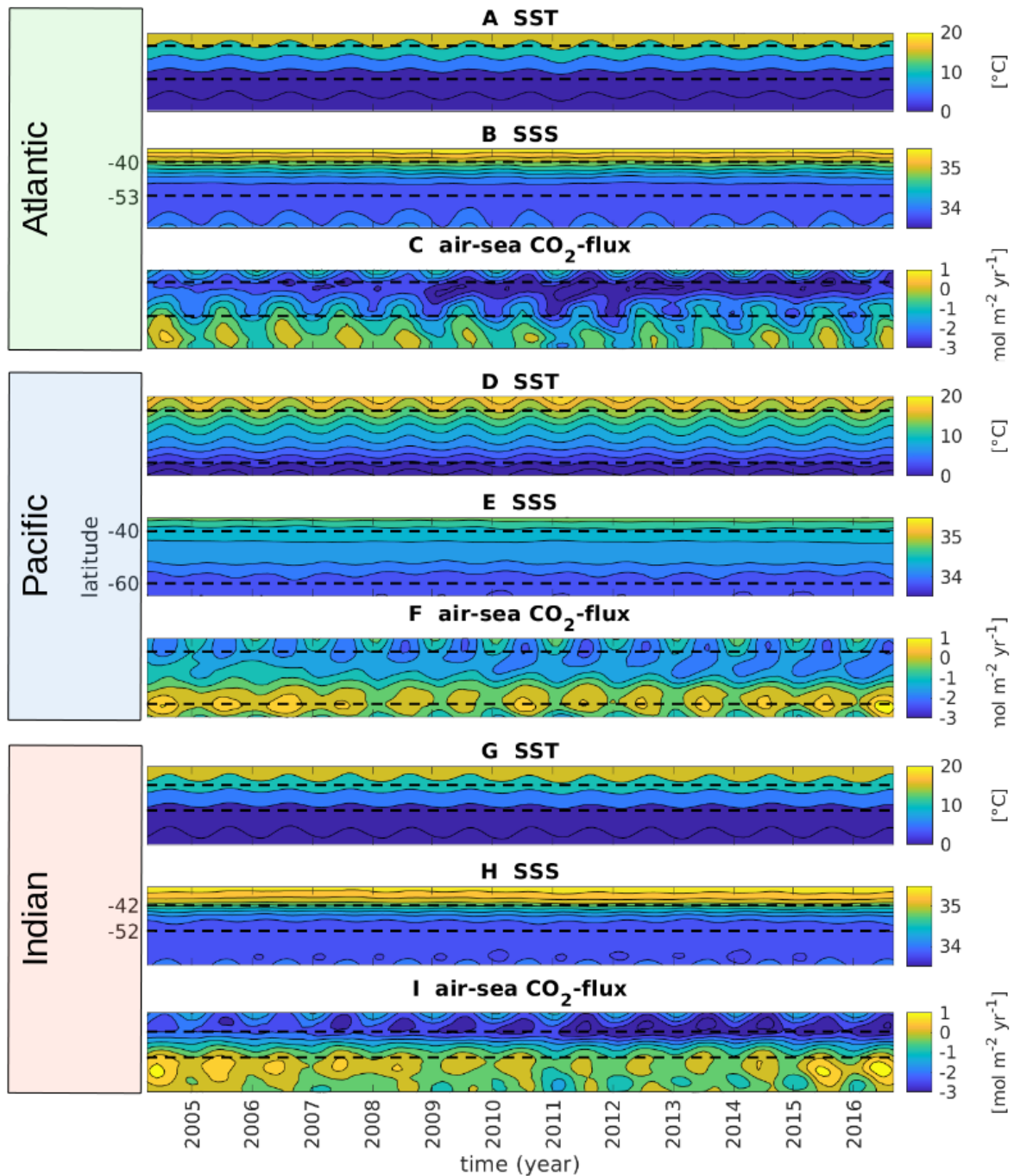
100 **Fig. S1.** As Fig. 2B but with the correlation between the SAM and sea level pressure: Correlation
 101 coefficients between the sea level pressure [hPa] and the standardized SAM index, smoothed
 102 with a 3-month running average, between January 1982 and December 2016. Coefficients with
 103 significance <95% are hatched. The mean positions of the PF (~55°S) and the STF (~40°S) are
 104 illustrated as thin black lines, and the three Southern Ocean sectors are delimited by dashed black
 105 lines.



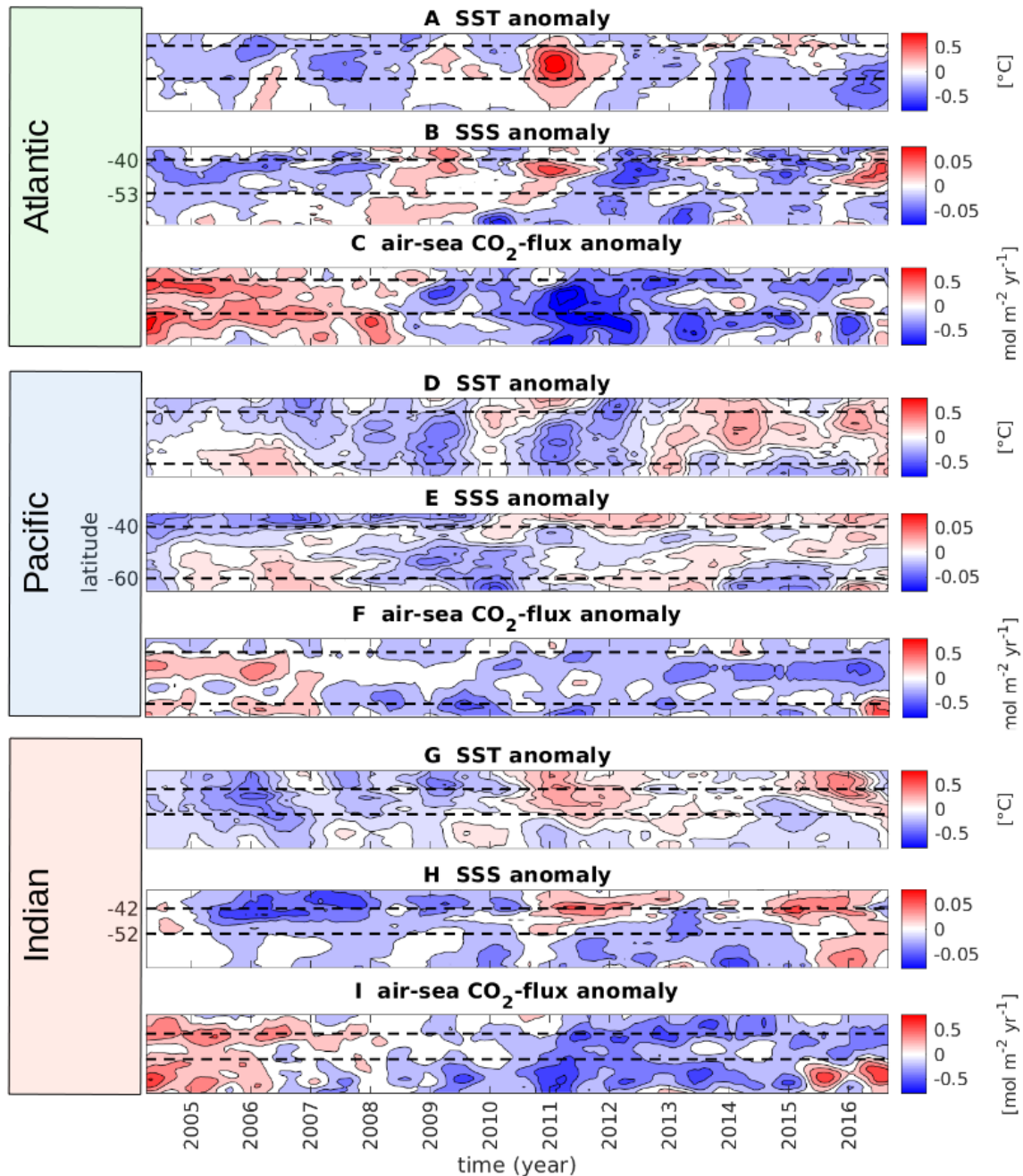
107 **Fig. S2.** Available pCO₂ observations [μ atm] and robustness of the interpolated data in the most
108 recent period (2012 through 2016): (A) The distribution of the shipboard pCO₂ observations in
109 the Southern Ocean from the SOCATv5 database for each 1°x1° grid point (B) The mean
110 observed pCO₂ from the SOCATv5 database (C) The mean interpolated pCO₂ from
111 Landschützer et al. [2015] (D) The mean residual pCO₂ (interpolated minus observed pCO₂). (E)
112 The standard deviation of the residual pCO₂.



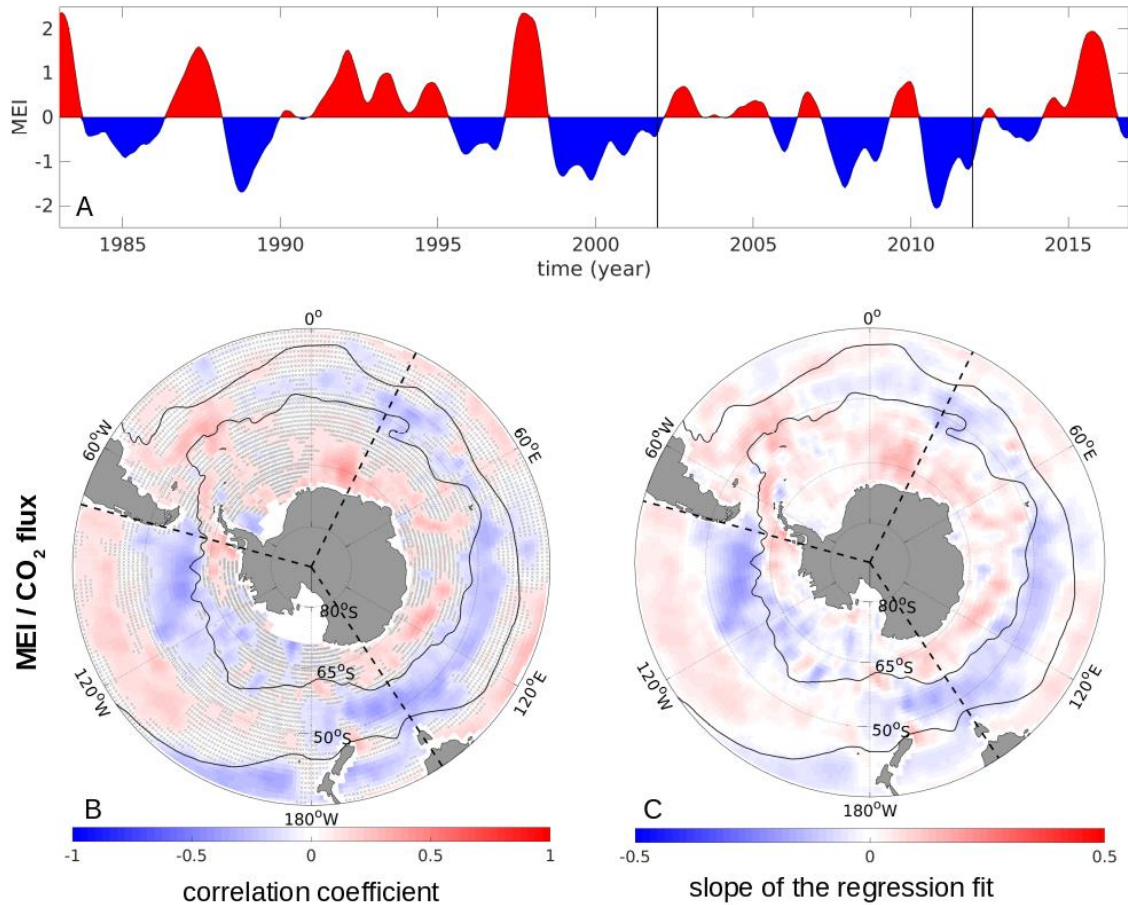
113 **Fig. S3.** CO₂ flux as a function of SST in each of the Southern Ocean sectors and interfrontal
 114 zones, using monthly means from 2004 through 2016: Atlantic (A), Pacific (B), and Indian sector
 115 (C), for the AAZ (green), PFZ (blue), and STZ (red).



116 **Fig. S4.** As Fig. 2, but the mean instead of the anomalies: Hovmöller plots of the zonal means of the Southern Ocean sectors as a function of time (x-axis) and latitude (y-axis) from 35°S to 65°S. SST [°C] (A,D,G) and SSS (B,E,H) in comparison to the carbon flux [mol m⁻² yr⁻¹] (C,F,I) for the Atlantic (A-C), Pacific (D-F), and Indian sectors (G-I). The mean positions of the STF (40°S, 40°S, 42°S) and PF (53°S, 60°S, 52°S) are shown for the Atlantic, Pacific, and Indian sectors respectively as dashed black lines. Negative values in the CO₂-flux indicate oceanic uptake. The seasonal cycle is not removed, but we smoothed with a 3-month running mean.

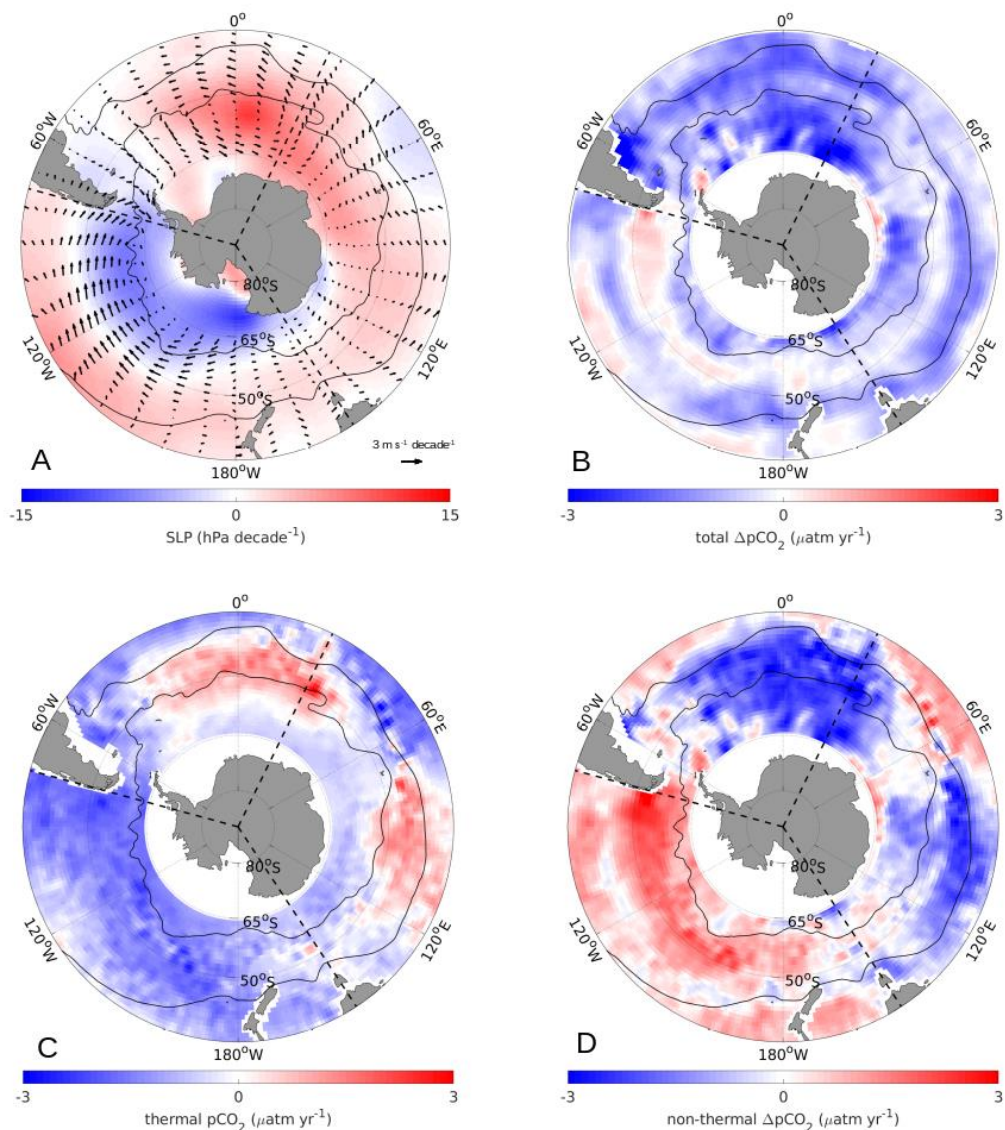


123 **Fig. S5.** Hovmöller plots of the zonal mean anomalies of the Southern Ocean sectors as a function of time
 124 (x-axis) and latitude (y-axis) from 35°S to 65°S. SST anomalies [°C] (A,D,G), SSS anomalies (B,E,H) and
 125 the carbon flux anomalies [mol m⁻² yr⁻¹] (C,F,I) for the Atlantic (A-C), Pacific (D-F), and Indian sectors
 126 (G-I). The mean positions of the Subtropical Front (STF, 40°S, 40°S, 42°S) and PF (53°S, 60°S, 52°S) are
 127 illustrated for the Atlantic, Pacific, and Indian sectors respectively as dashed black lines. The anomalies
 128 are based on the mean between 2004 and 2016, and the first and last 3 months are removed in the
 129 smoothing. Negative values in the CO₂-flux anomalies indicate a stronger sink. Note that while Fig. 1
 130 extends until the Antarctic coast (~77°S), Fig. 2 only extends until 65°S due to the data availability of the
 131 SST and SSS. See also S5 for the mean values instead of the anomalies.



133

134 **Fig. S6.** As Fig. 4, but with the Multivariate ENSO Index (MEI) instead of the SAM index: The
 135 relationship between the MEI and the CO₂ flux anomaly between January 1982 and December
 136 2016. (A) Standardized MEI smoothed with a 3-month running mean. Positive is shown in red,
 137 negative in blue. The start of the reinvigoration (Jan 2002) and the current period (Jan 2012)
 138 are marked with thin black lines. (B) Correlation coefficients between the air-sea CO₂ flux anomaly
 139 [mol m⁻² yr⁻¹] and the standardized MEI. Coefficients with significance <math>< 95\%</math> are hatched. (C)
 140 The slope of the regression fit between the air-sea CO₂-flux anomalies [mol m⁻² yr⁻¹] and the
 141 standardized MEI. As the MEI is standardized to have a mean of 0 and a standard deviation of
 142 1, (C) illustrates the change in the CO₂ flux [mol m⁻² yr⁻¹] per standard deviation of the MEI. (B-
 143 C) The mean positions of the PF (~55°S) and the STF (~40°S) are shown as thin black lines, the
 144 three Southern Ocean sectors are delimited by dashed black lines, and the coastal areas are
 145 masked white.



146 **Fig. S7.** As Fig. 4, but for the reinvigoration period (2004 through 2011) instead of the most
 147 recent period (2012 through 2016). Trends of the $p\text{CO}_2$, its components, and the sea level
 148 pressure and 10 m wind velocity during the reinvigoration period (2004 through 2011). (A) trend
 149 of the sea level pressure (hPa decade^{-1}) (color) and trend of the 10 m wind velocity ($\text{m s}^{-1} \text{decade}^{-1}$)
 150 (vectors). (B) Trend of the $\Delta p\text{CO}_2$ ($\mu\text{atm year}^{-1}$); (C) trend of the thermal component of the
 151 $p\text{CO}_2$ ($\mu\text{atm year}^{-1}$); (D) trend of the non-thermal component of the $\Delta p\text{CO}_2$ ($\mu\text{atm year}^{-1}$); The
 152 mean positions of the PF and the STF are shown as thin black lines and dashed black lines delimit
 153 the three Southern Ocean sectors. Note: a similar figure was shown in Landschützer et al. [2015]
 154 ⁸ for the period from 2002 through 2011. Note that the scale is smaller than in Fig. S3 for B-D.

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