

# Supplementary Information for

## Regional Wind Variability Modulates the Southern Ocean Carbon Sink

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- Supplementary text S1 to S7
- Figs. S1 to S7
- References for SI reference citations

1 **Supplementary Information Text**

2

3 **S1. The effect of Sea Level Pressure on the Southern Annular Mode (SAM)**

4 Previous studies have shown that the SAM has a zonal asymmetry that affects the mixed-  
5 layer depth and temperature<sup>1,2</sup>. Fig. S1 highlights the zonal asymmetry of the SAM, indicating  
6 that the negative correlation between the SAM and sea level pressure reaches further north in the  
7 eastern Pacific compared to the rest of the study region.

8

9 **S2. The available carbon dioxide (CO<sub>2</sub>) observations and robustness of the interpolated**  
10 **data**

11 The Southern Ocean is historically an under-sampled region, due to the harsh conditions  
12 and the relative remoteness. The available data of the partial pressure of CO<sub>2</sub> (pCO<sub>2</sub>) in the  
13 Southern Ocean come mainly from shipboard measurements, which have steadily increased in  
14 number in the past several decades, with the Drake Passage and the Tasman Sea being the best-  
15 sampled region within the Southern Ocean since the 2000s (Fig. S2A)<sup>3,4</sup>. Despite the substantial  
16 increase in observational pCO<sub>2</sub> data in the Southern Ocean, the spatio-temporal distribution  
17 remains sparse compared to other regions<sup>3</sup>. The robustness of the neural-network interpolated  
18 pCO<sub>2</sub> data until December 2011 has been demonstrated in previous studies<sup>5-7</sup>. Here, we show  
19 the robustness of the method for the most recent period.

20 Averaged over the most recent time period (2012 through 2016) the observations (Fig.  
21 S2B) are relatively well represented in the interpolated pCO<sub>2</sub> (Fig. S2C). Although some regional  
22 biases are present (Fig. S2D), they mostly cancel out when averaged over the study region (1.4  
23  $\mu\text{atm}$ ). The standard deviation at each grid point is shown in Fig. S2E, which add up to 5.6  $\mu\text{atm}$ .  
24 The Antarctic coastal areas display the largest standard deviation; however, our study mainly  
25 focuses on observations north of 65°S, due to the data availability of the temperature and salinity.

26

27 **S3. The effect of the El Niño Southern Oscillation (ENSO) on the Southern Ocean air-sea**  
28 **CO<sub>2</sub> flux**

29 We investigate the effect of the ENSO on the Southern Ocean carbon uptake, similarly  
30 as we did for the SAM (Fig. S3). We find the regional effect of the ENSO considerably smaller  
31 than the regional effect by the SAM. However, similar as the SAM, integrated over each of the  
32 three sectors, and over the whole Southern Ocean, the net effect of the ENSO on the Southern  
33 Ocean carbon uptake is ~zero.

34

35 **S4. The regional relationship between sea surface temperature (SST) and the CO<sub>2</sub> flux**

36 The air-sea CO<sub>2</sub> flux partially depends on SST, as CO<sub>2</sub> dissolves better in colder water  
37 However, other factors including biological production and vertical circulation also affect the  
38 oceanic CO<sub>2</sub> uptake. Here, we demonstrate the relationship between the SST and the air-sea CO<sub>2</sub>  
39 flux in the different sectors and interfrontal zones from 2004 through 2016 (Fig. S4).

40 In all three sectors of the Antarctic Zone, (AAZ, from 65°S to ~55°S), the SST ranges  
41 approximately from -2 to 5°C, and the carbon flux varies approximately from -5 to 3 mol m<sup>-2</sup>  
42 year<sup>-1</sup>. There is a slight trend where warmer surface waters tend to coincide with more uptake in  
43 the Atlantic and Indian sectors of the AAZ, indicating that solubility is not the main driver here.  
44 Concurrently, the Pacific sector of the AAZ does not show a considerable trend.

45 In the Polar Frontal Zone (PFZ, from ~55°S to ~40°S), the SST ranges approximately  
46 from 0 to 17°C, and the carbon flux varies from approximately -4 to 2 mol m<sup>-2</sup> year<sup>-1</sup>. In the  
47 Atlantic and Indian sectors of the PFZ, warmer surface waters tend to coincide with more uptake.  
48 However, the Pacific sector is a lot more variable and the colder surface waters have a similar

49 trend as the Atlantic and Indian sectors, but the warmer waters in this region have a reversal of  
50 this trend.

51 In the Subtropical Zone (STZ, from  $\sim 40^{\circ}\text{S}$  to  $30^{\circ}\text{S}$ ), the SST ranges approximately from  
52 10 to  $25^{\circ}\text{C}$ , and the carbon flux varies approximately from  $-5$  to  $1 \text{ mol m}^{-2} \text{ year}^{-1}$ . In all three  
53 sectors of the STZ, there is a trend of colder surface waters tending to coincide with more uptake.

54

#### 55 **S5. Mean sea surface properties and air-sea $\text{CO}_2$ flux of the sectors**

56 Here, we show the zonal mean sea SST, sea surface salinity (SSS) and the air-sea  $\text{CO}_2$   
57 flux in the Southern Ocean sectors to provide context (Fig. S5); the anomalies are shown in S6.

58

#### 59 **S6. Sea surface property anomalies and the air-sea $\text{CO}_2$ flux anomalies of the sectors**

60 We analyse the sea-air  $\text{CO}_2$  flux anomalies in comparison with the physical sea surface  
61 properties of the Southern Ocean (Fig. S6). Generally, we find that warmer and saltier surface  
62 waters tend to coincide with a stronger carbon sink. Therefore, it is evident that solubility is not  
63 the dominant mechanism driving the sink variability, but rather reflects the reduction in vertical  
64 mixing that usually brings cool but carbon-rich deep water to the surface. However, strong  
65 differences exist between both the different sectors and the interfrontal zones (see also S4).

66 Similar to the sea-air  $\text{CO}_2$  flux, the SST and SSS anomalies are strongest in the Atlantic  
67 sector (Fig. S6A-C). Here, the temporal evolution is dominated by a warm anomaly in 2011,  
68 north of the Polar Front (PF), which coincides with saltier surface waters and a stronger carbon  
69 sink. A similar, but weaker anomalous period can be observed around 2009, north of the PF.  
70 After this event, the Atlantic sector becomes cooler and fresher, while the carbon sink becomes  
71 weaker again, especially north of the PF.

72 In the Pacific sector, the signal varies on shorter time scales compared to the Atlantic  
73 sector in all three variables (Fig. S6D-F), which strongly coincides with the ENSO (see S3),  
74 suggesting the influence of remote modes of variability on the Southern Ocean CO<sub>2</sub> sink. While  
75 throughout the time period, warm SST anomalies usually coincide with saltier phases and vice  
76 versa, the carbon sink appears disconnected from this pattern. Warmer and saltier waters  
77 coincide with less carbon uptake in the PFZ from 2004 through 2011, which was previously<sup>7</sup>  
78 considered as evidence that the CO<sub>2</sub> flux trends in the Pacific sector are solubility dominated<sup>8</sup>.  
79 However, subsequently colder and fresher waters coincide with less carbon uptake in the PFZ  
80 after 2011, challenging this view.

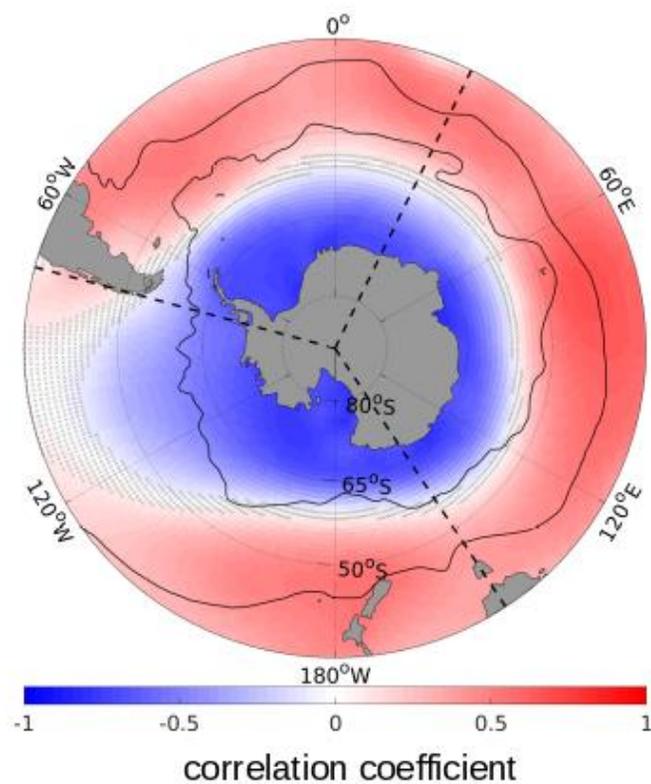
81 The Indian sector behaves similarly to the Atlantic sector in terms of SST, similar to the  
82 Pacific sector in terms of SSS, and like a mixture between the Atlantic and Pacific sectors in  
83 terms of its CO<sub>2</sub>-flux (Fig. S6G-I). For example, the period of much higher SST around 2011  
84 north of the PF that was observed in the Atlantic, is also present in the Indian sector, albeit with  
85 less intensity. In the following years, the surface waters in the STZ of both the Pacific and the  
86 Indian sectors are saltier than the mean. This trend moves further south in ~2016.

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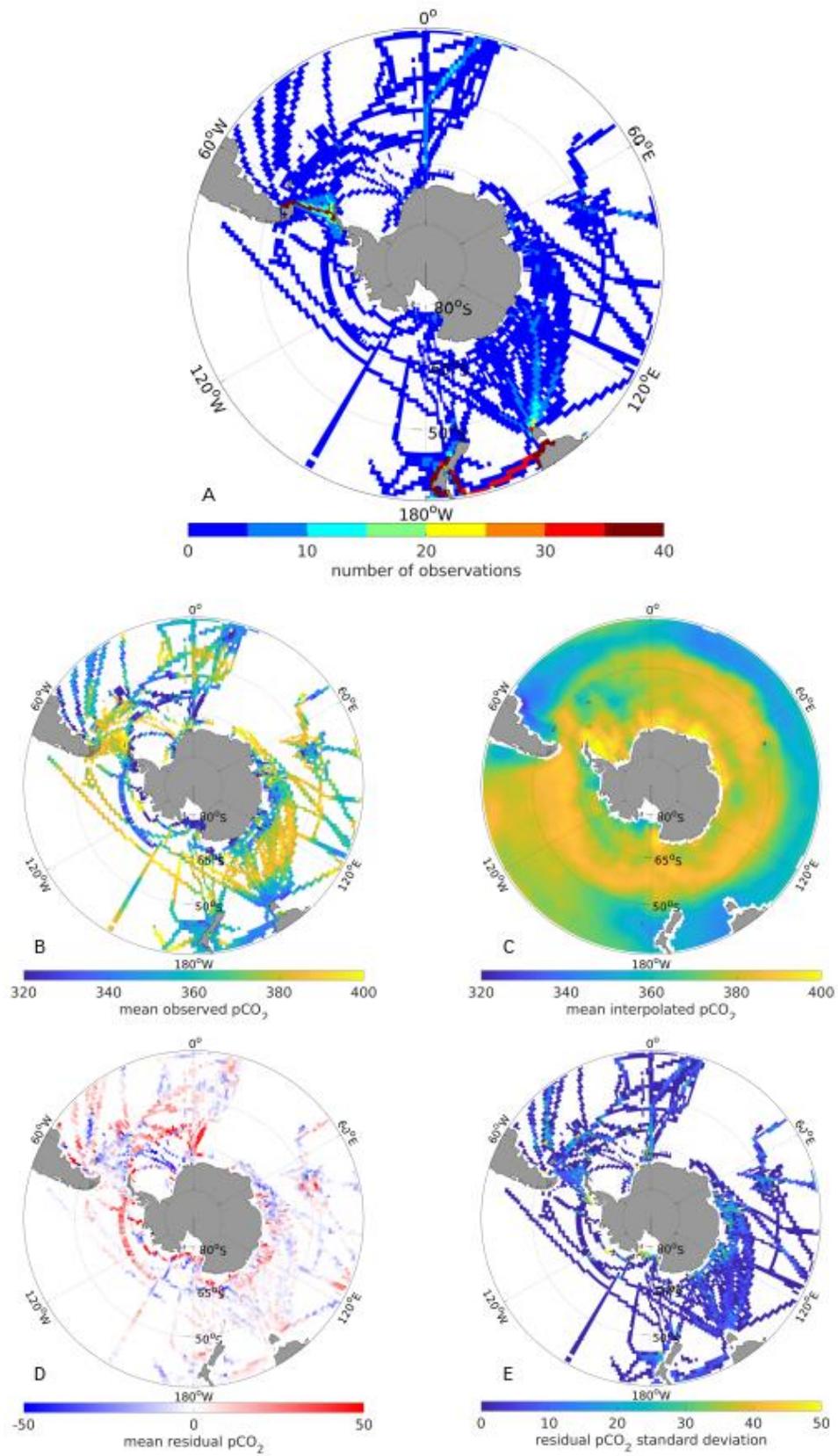
## 88 **S7. Trends of the pCO<sub>2</sub>, its components, and drivers in the reinvigoration period**

89 Here, we show the trends of the reinvigoration period to put our findings on the trends  
90 during the most recent period in context. As discussed in Landschützer et al. [2015]<sup>8</sup>, in the  
91 reinvigoration period, the ΔpCO<sub>2</sub> decreased in the Southern Ocean (Fig. S7A), resulting in  
92 enhanced CO<sub>2</sub> uptake by the ocean. The authors demonstrated that in this period the westerly  
93 winds were stronger in the Pacific and weaker in the Atlantic due to a dipole in sea level pressure  
94 (Fig. S7D). This change in surface wind patterns is thought to have caused enhanced  
95 downwelling and warmer surface waters in the Atlantic sector. The non-thermal component (Fig.  
96 S7C) dominated over the thermal component (Fig. S7B), resulting in an overall decrease in

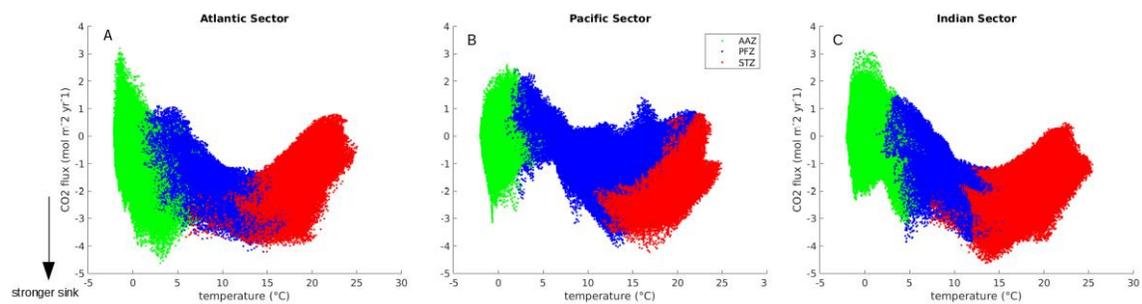
97  $\Delta p\text{CO}_2$  (Fig. S7A). Concurrently, the stronger westerlies in the Pacific sector caused enhanced  
98 upwelling and colder surface waters. Here, the thermal component (Fig. S7B) dominated over  
99 the non-thermal component (Fig. S7C), resulting in an enhanced  $\text{CO}_2$  uptake by the ocean.



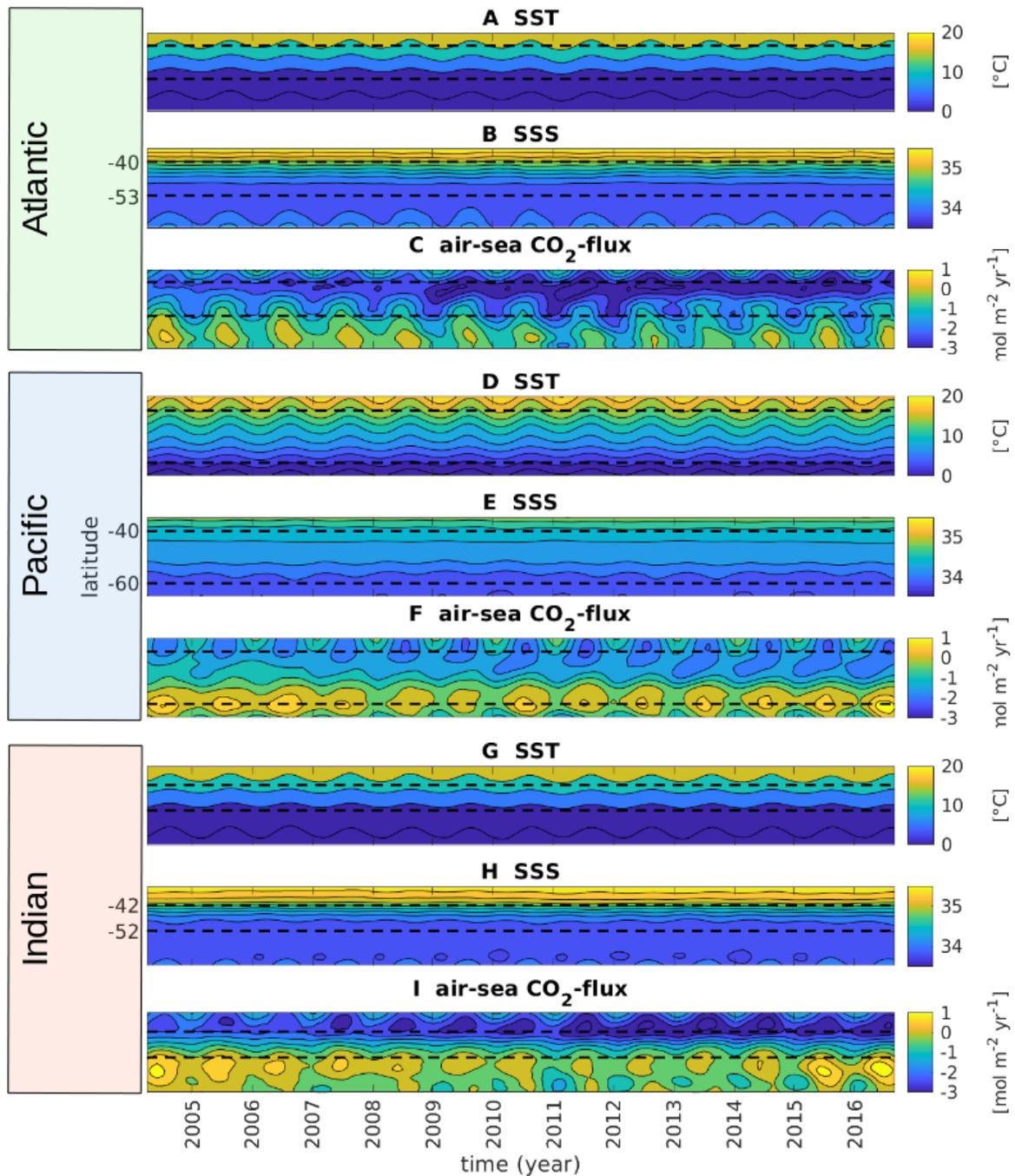
100 **Fig. S1.** As Fig. 2B but with the correlation between the SAM and sea level pressure: Correlation  
 101 coefficients between the sea level pressure [hPa] and the standardized SAM index, smoothed  
 102 with a 3-month running average, between January 1982 and December 2016. Coefficients with  
 103 significance <95% are hatched. The mean positions of the PF (~55°S) and the STF (~40°S) are  
 104 illustrated as thin black lines, and the three Southern Ocean sectors are delimited by dashed black  
 105 lines.



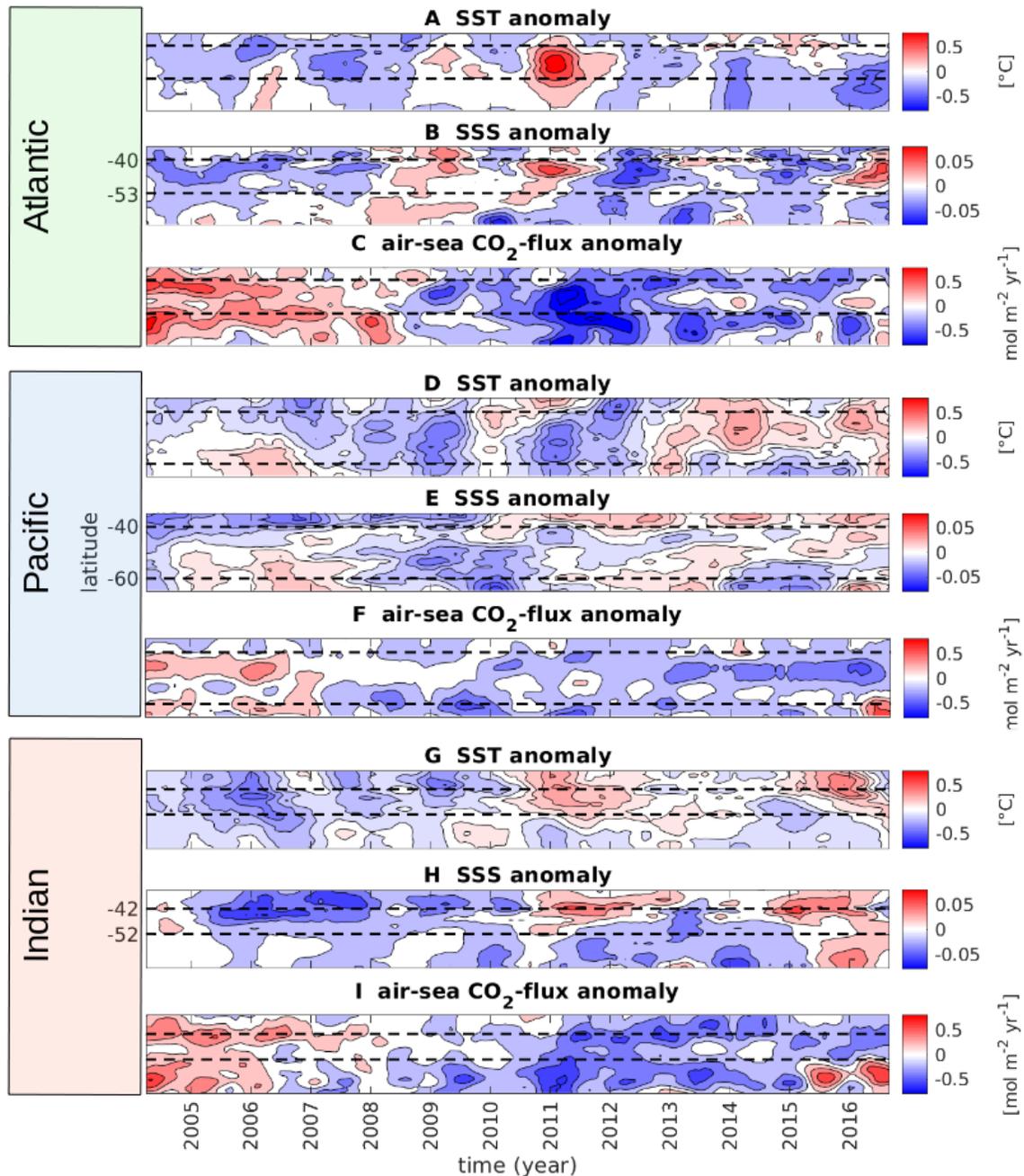
107 **Fig. S2.** Available pCO<sub>2</sub> observations [ $\mu$ atm] and robustness of the interpolated data in the most  
108 recent period (2012 through 2016): (A) The distribution of the shipboard pCO<sub>2</sub> observations in  
109 the Southern Ocean from the SOCATv5 database for each 1°x1° grid point (B) The mean  
110 observed pCO<sub>2</sub> from the SOCATv5 database (C) The mean interpolated pCO<sub>2</sub> from  
111 Landschützer et al. [2015] (D) The mean residual pCO<sub>2</sub> (interpolated minus observed pCO<sub>2</sub>). (E)  
112 The standard deviation of the residual pCO<sub>2</sub>.



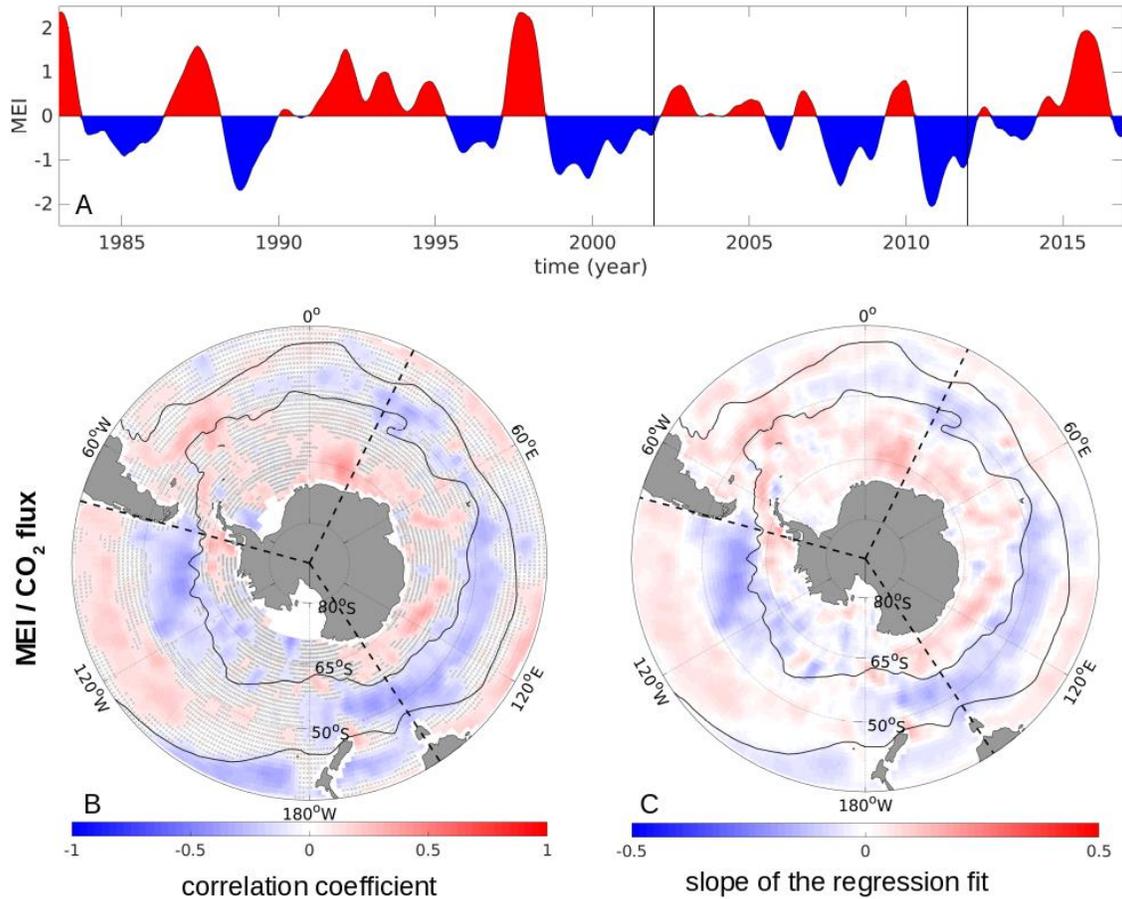
113 **Fig. S3.** CO<sub>2</sub> flux as a function of SST in each of the Southern Ocean sectors and interfrontal  
 114 zones, using monthly means from 2004 through 2016: Atlantic (A), Pacific (B), and Indian sector  
 115 (C), for the AAZ (green), PFZ (blue), and STZ (red).



116 **Fig. S4.** As Fig. 2, but the mean instead of the anomalies: Hovmöller plots of the zonal means of the Southern Ocean sectors as a function of time (x-axis) and latitude (y-axis) from 35°S to 65°S. SST [°C] (A,D,G) and SSS (B,E,H) in comparison to the carbon flux [mol m<sup>-2</sup> yr<sup>-1</sup>] (C,F,I) for the Atlantic (A-C), Pacific (D-F), and Indian sectors (G-I). The mean positions of the STF (40°S, 40°S, 42°S) and PF (53°S, 60°S, 52°S) are shown for the Atlantic, Pacific, and Indian sectors respectively as dashed black lines. Negative values in the CO<sub>2</sub>-flux indicate oceanic uptake. The seasonal cycle is not removed, but we smoothed with a 3-month running mean.

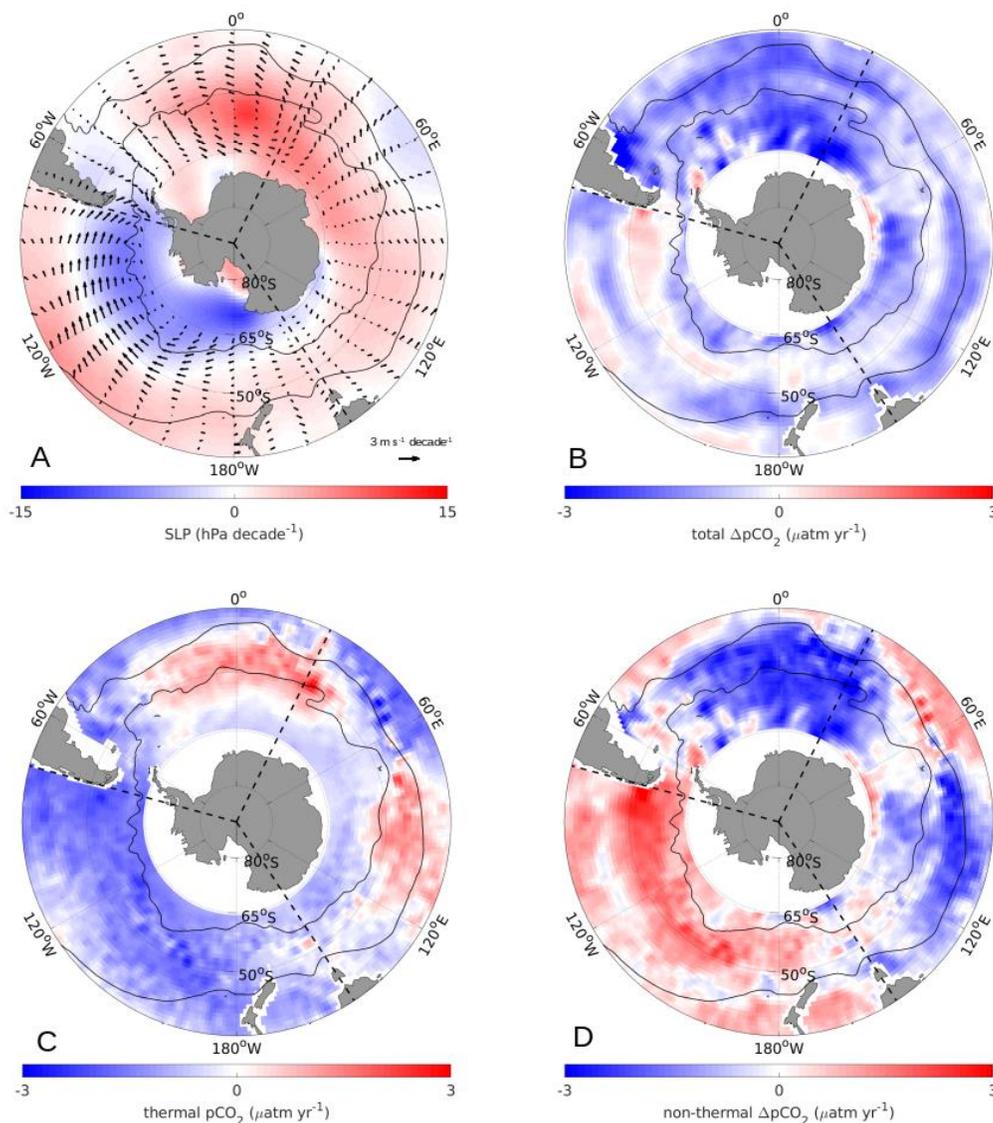


123 **Fig. S5.** Hovmöller plots of the zonal mean anomalies of the Southern Ocean sectors as a function of time  
 124 (x-axis) and latitude (y-axis) from 35°S to 65°S. SST anomalies [°C] (A,D,G), SSS anomalies (B,E,H) and  
 125 the carbon flux anomalies [mol m<sup>-2</sup> yr<sup>-1</sup>] (C,F,I) for the Atlantic (A-C), Pacific (D-F), and Indian sectors  
 126 (G-I). The mean positions of the Subtropical Front (STF, 40°S, 40°S, 42°S) and PF (53°S, 60°S, 52°S) are  
 127 illustrated for the Atlantic, Pacific, and Indian sectors respectively as dashed black lines. The anomalies  
 128 are based on the mean between 2004 and 2016, and the first and last 3 months are removed in the  
 129 smoothing. Negative values in the CO<sub>2</sub>-flux anomalies indicate a stronger sink. Note that while Fig. 1  
 130 extends until the Antarctic coast (~77°S), Fig. 2 only extends until 65°S due to the data availability of the  
 131 SST and SSS. See also S5 for the mean values instead of the anomalies.



133

134 **Fig. S6.** As Fig. 4, but with the Multivariate ENSO Index (MEI) instead of the SAM index: The  
 135 relationship between the MEI and the CO<sub>2</sub> flux anomaly between January 1982 and December  
 136 2016. (A) Standardized MEI smoothed with a 3-month running mean. Positive is shown in red,  
 137 negative in blue. The start of the reinvigoration (Jan 2002) and the current period (Jan 2012) are  
 138 marked with thin black lines. (B) Correlation coefficients between the air-sea CO<sub>2</sub> flux anomaly  
 139 [mol m<sup>-2</sup> yr<sup>-1</sup>] and the standardized MEI. Coefficients with significance <95% are hatched. (C)  
 140 The slope of the regression fit between the air-sea CO<sub>2</sub>-flux anomalies [mol m<sup>-2</sup> yr<sup>-1</sup>] and the  
 141 standardized MEI. As the MEI is standardized to have a mean of 0 and a standard deviation of  
 142 1, (C) illustrates the change in the CO<sub>2</sub> flux [mol m<sup>-2</sup> yr<sup>-1</sup>] per standard deviation of the MEI. (B-  
 143 C) The mean positions of the PF (~55°S) and the STF (~40°S) are shown as thin black lines, the  
 144 three Southern Ocean sectors are delimited by dashed black lines, and the coastal areas are  
 145 masked white.



146 **Fig. S7.** As Fig. 4, but for the reinvigoration period (2004 through 2011) instead of the most  
 147 recent period (2012 through 2016). Trends of the pCO<sub>2</sub>, its components, and the sea level  
 148 pressure and 10 m wind velocity during the reinvigoration period (2004 through 2011). (A) trend  
 149 of the sea level pressure (hPa decade<sup>-1</sup>) (color) and trend of the 10 m wind velocity (m s<sup>-1</sup> decade<sup>-1</sup>)  
 150 (vectors). (B) Trend of the ΔpCO<sub>2</sub> (μatm year<sup>-1</sup>); (C) trend of the thermal component of the  
 151 pCO<sub>2</sub> (μatm year<sup>-1</sup>); (D) trend of the non-thermal component of the ΔpCO<sub>2</sub> (μatm year<sup>-1</sup>); The  
 152 mean positions of the PF and the STF are shown as thin black lines and dashed black lines delimit  
 153 the three Southern Ocean sectors. Note: a similar figure was shown in Landschützer et al. [2015]  
 154 <sup>8</sup> for the period from 2002 through 2011. Note that the scale is smaller than in Fig. S3 for B-D.

155       **References**

- 156   1. Sallée, J. B., Speer, K. G. & Rintoul, S. R. Zonally asymmetric response of the Southern  
157       Ocean mixed-layer depth to the Southern Annular Mode. *Nat. Geosci.* **3**, 273–279 (2010).
- 158   2. Fogt, R. L., Jones, J. M. & Renwick, J. Seasonal Zonal Asymmetries in the Southern Annular  
159       Mode and Their Impact on Regional Temperature Anomalies. *J. Clim.* **25**, 6253–6270 (2012).
- 160   3. Bakker, D. C. E. *et al.* A multi-decade record of high-quality fCO<sub>2</sub> data in version 3 of the  
161       Surface Ocean CO<sub>2</sub> Atlas (SOCAT). *Earth Syst. Sci. Data* **8**, 383–413 (2016).
- 162   4. Munro, D. R. *et al.* Recent evidence for a strengthening CO<sub>2</sub> sink in the Southern Ocean from  
163       carbonate system measurements in the Drake Passage (2002-2015). *Geophys. Res. Lett.* **42**,  
164       7623–7630 (2015).
- 165   5. Landschützer, P. *et al.* A neural network-based estimate of the seasonal to inter-annual  
166       variability of the Atlantic Ocean carbon sink. *Biogeosciences* **10**, 7793–7815 (2013).
- 167   6. Landschützer, P., Gruber, N., Bakker, D. C. E. & Schuster, U. Recent variability of the global  
168       ocean carbon sink. *Glob. Biogeochem. Cycles* **28**, 927–949 (2014).
- 169   7. Landschützer, P., Gruber, N. & Bakker, D. C. E. Decadal variations and trends of the global  
170       ocean carbon sink. *Glob. Biogeochem. Cycles* **30**, 1396–1417 (2016).
- 171   8. Landschützer, P. *et al.* The reinvigoration of the Southern Ocean carbon sink. *Science* **349**,  
172       1221–1224 (2015).

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