

CORONAVIRUS

Novel quantification of regional fossil fuel CO₂ reductions during COVID-19 lockdowns using atmospheric oxygen measurements

Penelope A. Pickers^{1*}, Andrew C. Manning¹, Corinne Le Quéré¹, Grant L. Forster^{1,2}, Ingrid T. Lujikx³, Christoph Gerbig⁴, Leigh S. Fleming¹, William T. Sturges¹

It is not currently possible to quantify regional-scale fossil fuel carbon dioxide (ffCO₂) emissions with high accuracy in near real time. Existing atmospheric methods for separating ffCO₂ from large natural carbon dioxide variations are constrained by sampling limitations, so that estimates of regional changes in ffCO₂ emissions, such as those occurring in response to coronavirus disease 2019 (COVID-19) lockdowns, rely on indirect activity data. We present a method for quantifying regional signals of ffCO₂ based on continuous atmospheric measurements of oxygen and carbon dioxide combined into the tracer “atmospheric potential oxygen” (APO). We detect and quantify ffCO₂ reductions during 2020–2021 caused by the two U.K. COVID-19 lockdowns individually using APO data from Weybourne Atmospheric Observatory in the United Kingdom and a machine learning algorithm. Our APO-based assessment has near-real-time potential and provides high-frequency information that is in good agreement with the spread of ffCO₂ emissions reductions from three independent lower-frequency U.K. estimates.

INTRODUCTION

Fossil fuel combustion and industrial processes are responsible for the majority of anthropogenic carbon dioxide (CO₂) emissions, more than 70% of which are emitted from cities and urban areas (1). Despite their critical importance, our ability to evaluate reported emissions and to monitor and inform on the effectiveness of emissions reduction policies over the coming decades is currently limited (2, 3). This limitation was recently highlighted by the 2020–2021 pandemic of the coronavirus disease 2019 (COVID-19). To mitigate the spread of the virus, many countries implemented social distancing measures at national or regional scales, resulting in sudden and severe temporary reductions in emissions of CO₂ from fossil fuels (ffCO₂) (3–6) and anthropogenic air pollutants (6–8). While numerous studies have successfully reported on air pollutant COVID-19 reductions as observed from atmospheric measurements (6–8), determining ffCO₂ COVID-19 reductions in the atmosphere has been substantially more challenging, owing to the large variations in atmospheric CO₂ caused by terrestrial biosphere fluxes (9).

The Paris Agreement invokes an increased imperative to report anthropogenic CO₂ emissions accurately at the country and sub-country scale with transparency and consistency (2) and to develop methods for independent evaluation (10, 11). Currently, anthropogenic CO₂ emissions are self-reported to the United Nations Framework Convention on Climate Change using an indirect “bottom-up” approach, based primarily on energy statistics and emission factors, and an agreed methodology (12); however, large inconsistencies in bottom-up approaches have been reported, arising from inaccuracies in energy statistics and/or emission factors (13–15).

The Global Carbon Budget 2020 (5) provided a detailed assessment of the impact of COVID-19 on ffCO₂ emissions during 2020 at global and regional scales, based on a range of bottom-up assessments, including those that also incorporate recently available activity and mobility tracking data, such as the Carbon Monitor product (4) and the Priestley Centre estimate (6). Relative changes in emissions during 2020 from this suite of bottom-up estimates reveal large inconsistencies in many regions, such as the EU27 (the 27 member states of the European Union), for which the reductions in ffCO₂ are 9.6% [University of East Anglia estimate, hereafter “UEA” (3)], 12.9% (Priestley), 7.1% (Carbon Monitor), and 17% (Global Carbon Budget) (5). Furthermore, since the suite of Global Carbon Budget bottom-up estimates is not available in real time, year-to-date projections were included, on the basis of forward extrapolation of emissions reductions to the end of the year 2020, instead of using emissions estimates based on actual lockdown measures (5). Nevertheless, this comparison of methods gives an indication of the uncertainty in regional estimates based on indirect proxies.

Attempts to detect and quantify COVID-19-associated ffCO₂ emissions reductions using more direct “top-down” methods, based on atmospheric measurements and modeling, have largely been unsuccessful so far, particularly at regional and country scales. In the United Kingdom, a study based on atmospheric CO₂ data from the Deriving Emissions linked to Climate Change network (16) found that COVID-19 signals will only be detectable in daily CO₂ mole fractions after at least 33 months of sustained emissions reductions (9). Another study, using atmospheric CO₂ data from the European Integrated Carbon Observing System network (<https://icos-cp.eu/observations/atmosphere/stations>) and the Stochastic Time-Inverted Lagrangian Transport model (STILT) (17), was unable to detect COVID-19-related reductions in ffCO₂ associated with the first wave of lockdown measures in Europe (<https://icos-cp.eu/sc2020/abstracts#152>). In East China, a study using satellite CO₂ retrievals was also unsuccessful (18). In all of these studies, COVID-19 ffCO₂ reductions were obscured by fluxes of CO₂ between the atmosphere and the terrestrial biosphere, which are typically much

¹Centre for Ocean and Atmospheric Sciences, School of Environmental Sciences, University of East Anglia, Norwich NR4 7TJ, UK. ²National Centre for Atmospheric Science, University of East Anglia, Norwich NR4 7TJ, UK. ³Department of Meteorology and Air Quality, Wageningen University and Research, 6700AA Wageningen, the Netherlands. ⁴Department of Biogeochemical Systems, Max Planck Institute for Biogeochemistry, Jena, Germany.

*Corresponding author. Email: p.pickers@uea.ac.uk

larger than ffCO₂ emissions. At Hateruma Island, Japan, atmospheric measurements of CO₂ and methane (CH₄) were used to infer COVID-19–related ffCO₂ reductions, mostly from wintertime data (when biospheric activity is suppressed), and by assuming that biospheric-related variability in CO₂:CH₄ ratios was not different in 2020 compared to previous years (19).

At the urban scale, where the contribution of ffCO₂ emissions relative to biospheric CO₂ emissions is usually larger, detection of COVID-19 signals has been possible in some locations using observations from satellites. A reduction in ffCO₂ emissions of 11.5% was detected in China during January to April 2020 (compared to the same months in 2019) using satellite-based nitrogen dioxide (NO₂) observations and CO₂-to-NO_x emission ratios from a bottom-up estimate to formulate proxy ffCO₂ observations (20); however, about half of the satellite grids were excluded from this analysis owing to the prevalence of natural emissions in the observations, which rendered the data unusable for ffCO₂ quantification. In another study, CO₂ measurements from a high-density low-cost sensor network in the San Francisco Bay Area were combined with satellite measurements of solar-induced fluorescence (a proxy for biospheric CO₂ emissions) and a high-resolution bottom-up ffCO₂ emissions prior within an atmospheric inversion framework. The authors found a 30% reduction in ffCO₂ emissions during a 6-week period of the city's "shelter-in-place" order, compared to the previous 6-week period (21). In both of these studies, the detection of a distinct COVID-19 signal was only made possible owing to (i) the use of proxy/"tracer" observations to separate the anthropogenic and natural contributions to the total atmospheric CO₂ signal, (ii) the availability of high-resolution emissions or emission ratio information from bottom-up inventories, and (iii) the selection of urban-based measurement locations, where ffCO₂ signals are comparatively larger [typically from 0 to 30 parts per million (ppm)] than signals at non-urban sites (usually less than 10 ppm).

In addition to satellite-based proxies/tracers, natural and anthropogenic signals in CO₂ can also be separated with ground-based atmospheric measurements, using radiocarbon data (¹⁴CO₂), carbon monoxide data (CO), or a combination of both (22–29). ¹⁴CO₂, the current "gold-standard" ground-based ffCO₂ tracer method, is a high-precision measurement that has been recently used to successfully provide a top-down assessment of ffCO₂ emissions in the United States (11). The main limitations of using ¹⁴CO₂ are twofold: First, it is currently only possible to measure atmospheric ¹⁴CO₂ with high accuracy from discrete samples (i.e., noncontinuously, with relatively low temporal resolution), which are moreover expensive and laborious to analyze (30); second, in some regions, such as the United Kingdom, ¹⁴CO₂ measurements can be severely influenced by CO₂ emissions from gas-cooled nuclear power plants, which obscure ffCO₂ signals in ¹⁴CO₂ data (31, 32). CO, a continuous high-frequency ffCO₂ tracer that is easier to measure and is unaffected by nuclear power plant emissions, can also be used, either as an alternative to ¹⁴CO₂ sampling or in conjunction with ¹⁴CO₂; however, CO-based ffCO₂ is limited by poor precision and accuracy, mostly arising from highly variable and inaccurate CO:ffCO₂ emission ratio information, which is required for CO-based ffCO₂ quantification (24, 25). Despite these limitations, ground-based measurements are more precise and accurate than satellite-based measurements; conversely, satellite-based measurements provide higher spatial coverage than ground-based measurements. To date, very few studies have been able to use ground-based atmospheric

measurements to provide a top-down assessment of COVID-19–related ffCO₂ emissions reductions.

The rate of COVID-19–related emissions reductions during 2020–2021 was similar to the rate of long-term emissions reductions required by the Paris Agreement to reach net zero emissions and limit global temperature rise in the range of 1.5° to 2°C. COVID-19 has demonstrated that, despite the critical importance of ffCO₂ emissions reductions for climate change policy, we do not currently have systems in place—either bottom-up, top-down, or a combined approach—to monitor and report ffCO₂ emissions at global, regional, or country scales in near real time (3).

Here, we present a new ground-based measurement approach for quantifying the regional ffCO₂ component of the atmospheric CO₂ mole fraction (in parts per million) using atmospheric potential oxygen (APO) data. We demonstrate the potential of APO as a ffCO₂ tracer by detecting and quantifying COVID-19 ffCO₂ reductions in the atmosphere associated with the first two waves of the pandemic in the United Kingdom, using continuous data from the Weybourne Atmospheric Observatory (WAO) in the United Kingdom and a machine learning algorithm. The APO-based assessment we present here separates biospheric and anthropogenic signals in atmospheric CO₂ with high frequency (e.g., daily or subdaily scales) and in near real time, which is an important first step toward robust quantification of absolute ffCO₂ emissions using atmospheric data. Our approach does not quantify absolute emissions, but, with the use of machine learning, we are able to quantify relative changes in emissions using APO data, which represents a major achievement in top-down observation-based ffCO₂ emissions quantification efforts. Using a combined APO and machine learning approach, we have detected a local 1.6-ppm reduction in daily-mean ffCO₂ observed at WAO during March to July 2020 compared to the non-pandemic "counterfactual scenario" (i.e., compared to the expected ffCO₂ during 2020 if the COVID-19 pandemic had not occurred), and a 1.3-ppm daily-mean reduction during November 2020 to January 2021. These two U.K. lockdown periods are separated by a period of recovery, from August to October 2020, characterized by little reduction in ffCO₂. Our APO-based estimate is in good agreement with the spread of ffCO₂ reductions determined from three independent bottom-up emissions estimates for the United Kingdom.

RESULTS

Calculation of ffCO₂ from APO

APO is a tracer that combines oxygen (O₂) and CO₂ observations (APO = O₂ + 1.1 × CO₂) (33), where the value of 1.1 denotes the mean –O₂:CO₂ molar ratio of terrestrial biosphere-atmosphere exchange (34). APO is therefore, by design, invariant to terrestrial biosphere exchange processes. We calculate ffCO₂ from APO, which we refer to as ffCO₂[APO], according to

$$\text{ffCO}_2[\text{APO}] = \frac{\text{APO} - \text{APO}_{\text{BL}}}{R_{\text{APO}}} \quad (1)$$

where APO_{BL} is the "baseline" APO value, which is determined statistically (see Fig. 1C and Materials and Methods), and R_{APO} is the molar ratio (R) of APO:CO₂ for fossil fuel emissions, derived from an emissions database product (35) (see Materials and Methods for details). Both APO and APO_{BL} have units of "per meg," rather than mole fraction units (such as parts per million) because O₂ is not a

trace gas and its mole fraction is therefore affected by small changes in other gases, such as CO₂ (36). While APO has typically been used in the past to remove the influence of land biospheric exchange to isolate air-sea O₂ and CO₂ fluxes (33, 37), here, we use APO to isolate fossil fuel emissions by subtracting a baseline that includes air-sea influences on APO (which mostly operate on long-term and seasonal time frames; see Materials and Methods). The numerator (APO-APO_{BL}) thus isolates short-term anomalies in APO, from which ffCO₂[APO] is determined. We convert the units of APO-APO_{BL} from per meg to parts per million equivalent by dividing by 4.77 (38). A full derivation of Eq. 1 is provided in appendix SA of the Supplementary Materials.

Figure 1 shows the atmospheric CO₂, O₂ and APO record from WAO in the United Kingdom, and the resulting ffCO₂ calculated from APO. Because of WAO's rural location on the north Norfolk coast, atmospheric transport variability exerts a substantial influence on the observed ffCO₂ signal (figs. S1 and S2). Air masses arriving at the site from the North Sea and the Arctic Ocean are generally associated with lower ffCO₂ compared to air masses arriving from the direction of southern England, the Midlands, or from the European continent. A cumulative plot of ffCO₂ for each year during 2011–2020, shown in Fig. 2, reveals that there is no apparent difference in the total ffCO₂ observed during 2020 compared to previous years because the dominating influences of wind

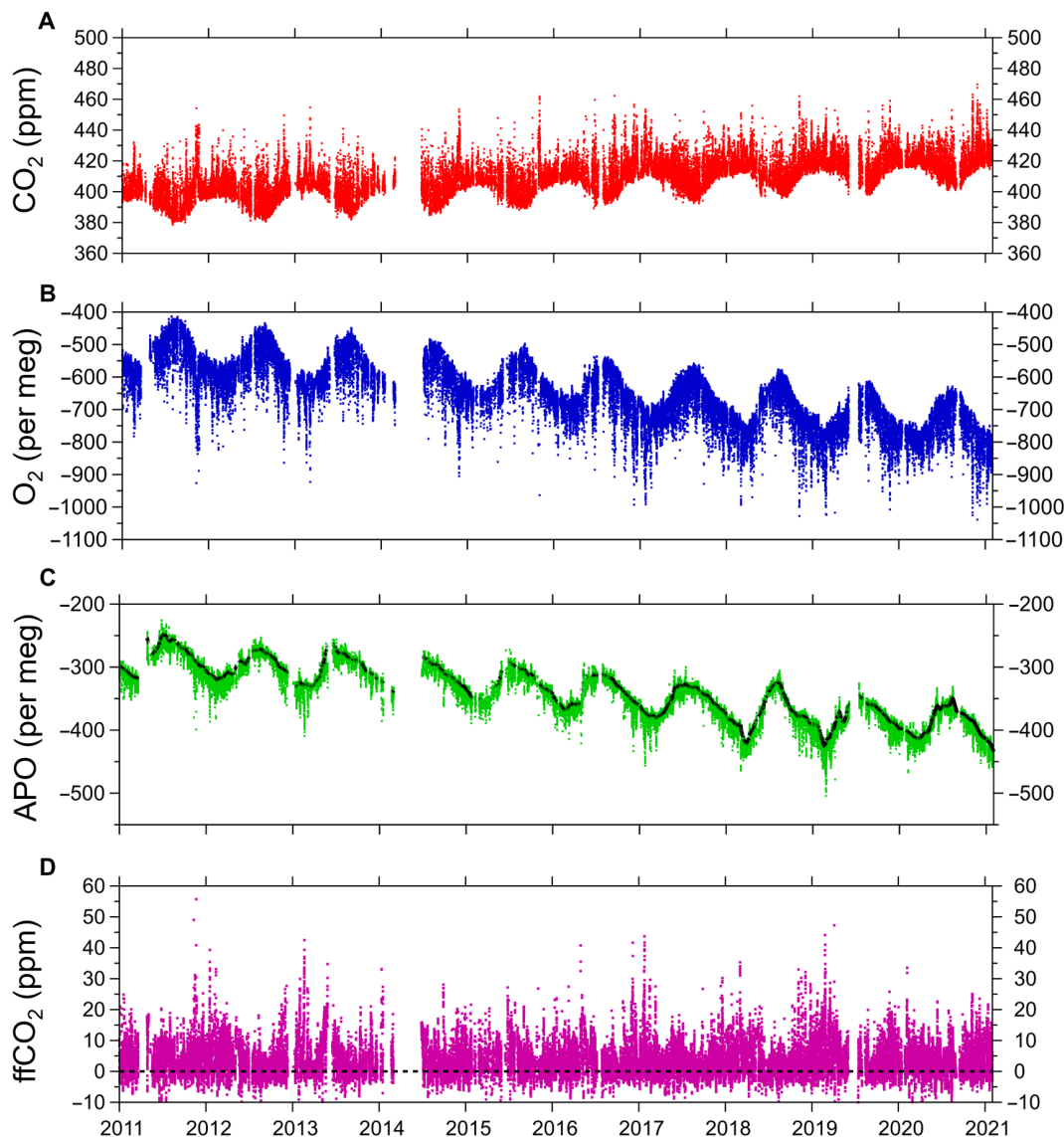


Fig. 1. Hourly atmospheric CO₂, O₂, and APO observations and calculated ffCO₂ from the WAO, 2011–2021. (A) Atmospheric CO₂ in parts per million. (B) Atmospheric O₂ in per meg units. A 1-ppm change in CO₂ is equivalent to a 4.77-per meg change in O₂ (38). (C) APO, also in per meg units. The black points in (C) are the statistically determined “baseline,” i.e., the APO_{BL} term in Eq. 1. (D) ffCO₂, calculated from APO by removing the baseline signal in (C) from the APO observations and dividing by R_{APO} as in Eq. 1. The black dashed line denotes “zero” ffCO₂, which is defined as the statistically determined baseline APO concentration. (A) to (C) show seasonality that is driven mostly by terrestrial biospheric processes (CO₂ and O₂) and oceanic processes (O₂ and APO). Shorter-term variability in all panels is driven by diurnal processes, changes in meteorological conditions, synoptic-scale variability, and ffCO₂ emissions. Gaps in the data are caused by instrument downtimes. x axis tick labels denote the beginning of the year shown.

direction and atmospheric transport on the ffCO_2 signals at WAO have not been accounted for.

COVID-19 ffCO_2 detection

We have applied the APO method to detect and quantify the reduction in $\text{ffCO}_2[\text{APO}]$ associated with the U.K. COVID-19 lockdown restrictions, using an 11-year continuous APO dataset from WAO and a random forest machine learning algorithm (39) to account for the effects of weather and atmospheric transport processes on APO (Fig. 3). A suite of 10 independent variables with hourly time resolution (see Materials and Methods) was used to train the machine learning algorithm to model ffCO_2 for the period 2010–2019. The algorithm was then used to predict the counterfactual case for February 2020 to January 2021, that is, the expected ffCO_2 that would have been observed at WAO during this period had there been no pandemic. Weekly differences are shown in Fig. 3A, which indicate reductions in ffCO_2 relative to the counterfactual (nonpandemic) prediction during periods when COVID-19 restrictions were in place.

The COVID-19 influence on ffCO_2 detected at WAO is highlighted using the cumulative signal, shown in Fig. 3B, which accumulates differences in the short-term variability of the daily values. We find a sustained decrease in daily-mean $\text{ffCO}_2[\text{APO}]$ relative to the counterfactual (nonpandemic) prediction of 1.6 ppm from 20 March to 31 July 2020, coinciding with the first period of U.K. lockdown, a recovery period during August to October 2020 during which U.K. lockdown restrictions were eased and $\text{ffCO}_2[\text{APO}]$ increased slightly by 0.2 ppm, and a second sustained decrease in daily-mean $\text{ffCO}_2[\text{APO}]$ of 1.3 ppm from early November 2020 until the end of January 2021, during which a national lockdown

was reintroduced. We deem these sustained signals to be caused by reductions in ffCO_2 emissions within the footprint of WAO, where the term “footprint” denotes the sensitivity of measurements at WAO to emissions located upwind of the site location (24). WAO is influenced most by southwesterly winds, so the site predominantly captures ffCO_2 signals from London and southern England over emissions from other wind sectors. WAO is therefore not representative of the United Kingdom as a whole (signals at WAO can also include emissions from continental Europe in addition to U.K. emissions under certain atmospheric conditions).

For the full year 2020, we find an overall mean reduction of 0.7 ppm. This estimate is higher than the expected global CO_2 change associated with COVID-19 emissions reductions [which is ~ 0.3 ppm for an 8% reduction in 2020 annual emissions according to another study (40)], most likely because the U.K. drop in COVID-19 emissions was substantially larger than the global average (41).

Comparison with independent estimates

Both the timing of the onset and the shape of the cumulative ffCO_2 signal from APO (Fig. 3B) agree with three bottom-up estimates of the emissions decrease from COVID-19 lockdown measures for the United Kingdom, based on indirect activity data (black lines in Fig. 3B). The magnitudes of these estimates are not directly comparable because the $\text{ffCO}_2[\text{APO}]$ top-down signal is in units of parts per million \times days and is not representative of the United Kingdom as a whole, while the bottom-up COVID-19 signals are in megatons of CO_2 and are U.K. totals. In addition, our $\text{ffCO}_2[\text{APO}]$ reduction is relative to the counterfactual prediction for 2020, whereas the bottom-up estimates are relative to emissions for 2019. Nevertheless, for the period 1 February 2020 to 31 January 2021, we find a 23%

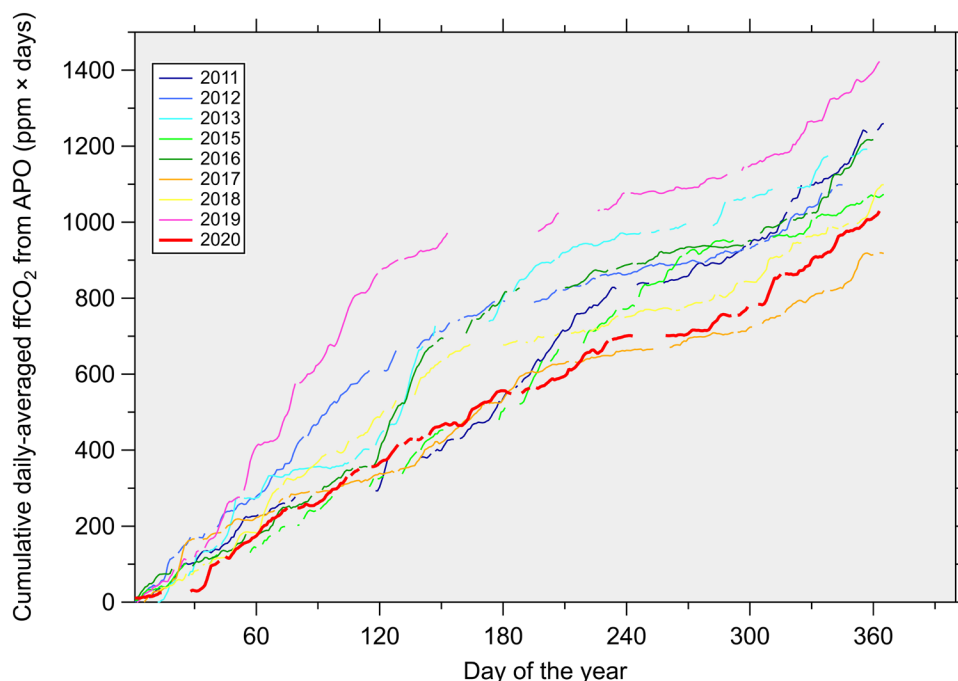


Fig. 2. Cumulative daily ffCO_2 from APO in parts per million \times days observed at WAO. Nonpandemic years (2011 to 2019) are shown by the thinner colored lines, except for the year 2014, which is omitted because of large gaps in the data. The year 2020, during which the COVID-19 pandemic started, is shown by the thicker red line. The influence of gaps on the cumulative signals have been accounted for by adjusting the ffCO_2 by the proportion of days that are missing data in each year. The 29 February has similarly been excluded where relevant, to allow a fair comparison between leap and nonleap years.

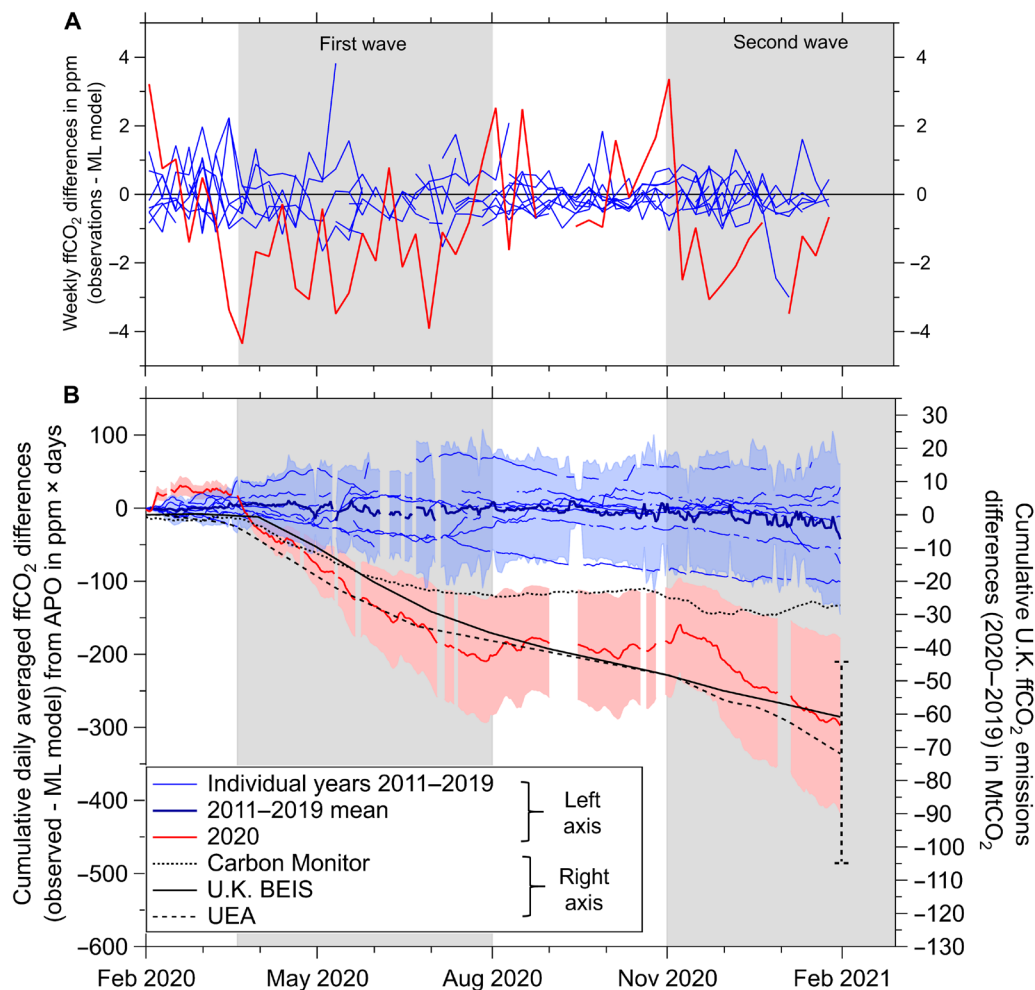


Fig. 3. Reduction in WAO ffCO_2 associated with COVID-19 lockdowns. (A) Differences in ffCO_2 [as ffCO_2 determined from APO minus modeled ffCO_2 determined from a random forest machine learning (ML) algorithm], shown as weekly differences. The first and second U.K. COVID-19 waves are indicated by the gray background shading. Differences for the individual years 2011–2019 are shown in blue. The period February 2020 to January 2021 is shown in red. All units are parts per million; x-axis major tick marks denote the first day of the month. Uncertainties are omitted from this panel for clarity. (B) Same as (A), but shown as cumulative daily-averaged ffCO_2 in units of parts per million \times days. The thick blue line indicates the 2011–2019 mean. Uncertainties are as follows: The blue shading is the $\pm 2\sigma$ (95%) SD of the 2011–2019 mean, and represents the uncertainty of the training model (i.e., if the model performance was perfect, then the blue lines would all be zero), which, in part, arises from the long-term decreasing trend in U.K. emissions over the period 2011–2019; ffCO_2 uncertainty for February 2020 onward is shown by the pale red shading and arises from the poorer performance of the predictive model relative to the training model (see the “Analysis of uncertainties” section for details). For comparison with our $\text{ffCO}_2[\text{APO}]$ detected COVID-19 signal, we also show 2020–2019 differences from three bottom-up U.K. emissions estimates (black lines) on the right-hand axis in units of MtCO_2 (see Materials and Methods). Only the UEA value (black dashed line) includes an estimate of uncertainty, shown by the vertical error bar.

reduction (range of 14 to 32%) from $\text{ffCO}_2[\text{APO}]$, compared to a 17% reduction from the U.K. BEIS (Department for Business, Energy, and Industrial Strategy) national inventory. Estimates based on proxy data give an 8% reduction from Carbon Monitor (4) and a 21% reduction (range of 13 to 30%) from the updated UEA estimate (3). Uncertainty ranges are not available for Carbon Monitor and U.K. BEIS estimates.

The spread in bottom-up estimates shown in Fig. 3B can in part be accounted for by differences in international aviation and shipping (IAS) emissions, which are included in the UEA estimate but are not included in the U.K. BEIS estimate. The Carbon Monitor estimate includes international aviation but not shipping. U.K. IAS emissions reductions resulting from the COVID-19 pandemic are estimated to be 17.3 MtCO_2 in 2020 (41), which would bring the

Carbon Monitor estimate closer to the UEA estimate but does not account for all of the offset. A similar adjustment to the U.K. BEIS estimate would shift it lower than the UEA value, so that even if IAS is included in all three bottom-up estimates, a range of $\sim 30 \text{ MtCO}_2$ would still persist. In addition, since only the UEA estimate includes an estimate of uncertainty, it is likely that the spread between bottom-up estimates would be substantially larger than this if uncertainties were available for the U.K. BEIS and Carbon Monitor estimates.

Analysis of uncertainties

We will show in this section that it is not required to quantify uncertainty in $\text{ffCO}_2[\text{APO}]$ to calculate the uncertainty in relative ffCO_2 emissions reductions from APO, but we nevertheless include

detailed information in Materials and Methods on $\text{ffCO}_2[\text{APO}]$ uncertainties for reasons of transparency and completeness. In summary, $\text{ffCO}_2[\text{APO}]$ uncertainties are calculated for each of the terms in Eq. 1: the R_{APO} value for WAO, the statistically derived baseline uncertainty APO_{BL} , and the uncertainty of the APO data themselves. Total hourly $\text{ffCO}_2[\text{APO}]$ uncertainty can thus be calculated by subtracting and dividing the absolute and relative errors according to the rules of error propagation.

For our COVID-19 analysis, while the $\text{ffCO}_2[\text{APO}]$ values could contain bias, either from inaccuracies in APO_{BL} or from an inaccurate R_{APO} value, we do not expect such bias, if it exists, to translate into error in our COVID-19-related relative ffCO_2 emissions reduction estimate because the random forest algorithm is trained on similarly biased $\text{ffCO}_2[\text{APO}]$ data (and since we calculate observation-model differences, most bias should cancel). This assumption relies on our $\text{ffCO}_2[\text{APO}]$ error remaining the same during both the training and predictive steps. For the APO measurement data and APO_{BL} , there is no indication in the diagnostic data that this is not the case: Measurement performance during 2020–2021 was similar to previous years, and there is no reason why the statistical baseline fitting routine should be less accurate during this time period than previously nor any evidence in the data to suggest that this is the case.

For R_{APO} , a shift in this value associated with COVID-19 lockdowns would potentially bias our results because we used a mean R_{APO} value of 0.37 throughout the whole 2010–2021 period; however, it is unlikely that any COVID-19-related R_{APO} shift occurred, since bottom-up estimates have shown that the reduction in ffCO_2 emissions is mostly in aviation and surface transport (3, 4). The former should not have a large impact on our APO-based analysis because WAO is not situated near any airports. We do not expect the latter to substantially bias R_{APO} because surface transport emissions are predominantly from liquid-based fuels, with R_{APO} values of ~ 0.34 , which sit approximately in the middle of the R_{APO} range between solid- and gas-based fuel types (35). A 36% reduction in surface transport, as found by Le Quéré *et al.* (3), would correspond to a bias in the R_{APO} value of +0.02, which would have a very small impact on $\text{ffCO}_2[\text{APO}]$. In addition, we have calculated R_{APO} using the U.K. BEIS inventory (fig. S3), which shows a change in R_{APO} during the COVID-19 lockdown periods of only +0.01. During 2011–2019, R_{APO} does not change substantially except during 2012–2015, when the value changes by -0.05 because of reduced coal emissions. R_{APO} returns to its original 2011 value in 2016 due to increases in gas usage, which counteracts the previous influence of coal on R_{APO} . Thus, there is no requirement to account for $\text{ffCO}_2[\text{APO}]$ uncertainties in our COVID-19 analysis, although these do exist (see Materials and Methods) and may need to be considered carefully in potential follow-up studies.

It is, however, necessary to consider the performance of the machine learning algorithm and its associated uncertainty, which we assess as follows. First, we evaluate the performance of the model, as shown in Fig. 4. The model underestimates the true range of variability of the APO-based ffCO_2 but generally performs well with a relatively small bias (-0.05 ± 2.34 ppm; see Fig. 4A and fig. S4). The impact of the imperfectly trained model performance on our ability to robustly detect COVID-19 reductions in ffCO_2 is shown by the blue lines in Fig. 3B, which would all be zero if the trained model was perfect. The spread in these lines also includes the decreasing trend in U.K. emissions (-22% during 2011–2019, from BEIS data), which is not captured by the machine learning algorithm because

bottom-up emissions are deliberately not used in the model training as a variable. By examining the partial dependencies of the independent variables (Fig. 4B), which shows the relationship between each variable and ffCO_2 for the training model, we can also see whether the model performs as we would expect. We find that high ffCO_2 at WAO is associated with lower air temperatures (during winter), higher atmospheric pressure and lower wind speeds (during more stable conditions), higher radon-222 activity (more influence from ground-based sources and/or lower planetary boundary layer height), and when clustered trajectories from the Hybrid Single Particle Lagrangian Integrated Trajectory (HYSPPLIT) model (fig. S5) originate from the south, especially the southeast (i.e., from the European continent). Low ffCO_2 at WAO is observed during opposite conditions, such as during the summer months and when the wind direction is northerly (from the ocean). A ranking of the importance of each independent variable for the model training is shown in fig. S6. These results collectively indicate that the trained machine learning model provides a physically realistic representation of ffCO_2 at WAO.

Second, we recognize that the model prediction is less robust than the trained model, and an additional uncertainty ($\pm 40\%$; pale red shading in Fig. 3B) is assigned to the 2020–2021 counterfactual case to account for this. This $\pm 40\%$ model prediction uncertainty was estimated by quantifying the difference between the predicted counterfactual from a separate model that was trained on 2010–2018 data and used to predict the period 1 January to 31 December 2019 ffCO_2 , to the trained data from the original model for the period 1 January to 31 December 2019. Using the model to predict ffCO_2 during previous nonpandemic years does not result in erroneous COVID-19-type signals, which should occur if the model prediction is consistently prone to overestimation.

Third, to ensure that the random forest prediction is not overly sensitive to the data at the beginning of the period (i.e., to so-called “end effects”), we examined ffCO_2 reductions from a variety of predictions run with differing start dates (fig. S7). Although the choice of start date does have a small impact on the magnitude of the pandemic ffCO_2 signals, all predictions still show similar patterns associated with the two COVID-19 lockdown periods with a period of lockdown easing during summer 2020. None of the differences are outside of the uncertainties, except for during the first few months, when the uncertainties are small. The sensitivity of the random forest prediction to other parameters, such as the R_{APO} value used to calculate ffCO_2 from APO, and the stiffness of the APO baseline fit were also tested, but no notable differences were found.

COVID assessment from WAO atmospheric CO_2 data

We also applied the random forest machine learning algorithm to WAO atmospheric CO_2 data instead of APO data (Fig. 5) but found that a signal that could potentially be ascribed to ffCO_2 reductions only emerges from about mid-September 2020 onward, 6 months after lockdown restrictions were first introduced. The timing of this potential signal is not consistent with the timing of the United Kingdom’s second lockdown, and it is not possible to separate the observed CO_2 reduction into ffCO_2 emissions reductions versus changes (either reduction or enhancement) in biospheric CO_2 emissions. It is more likely that the initial reduction in CO_2 during the late summer of 2020 is caused by an anomaly in biospheric CO_2 fluxes, perhaps caused by the heatwave in the late summer; hence, no quantitative assessment of the impacts of COVID-19 on ffCO_2 can be made using the atmospheric CO_2 data.

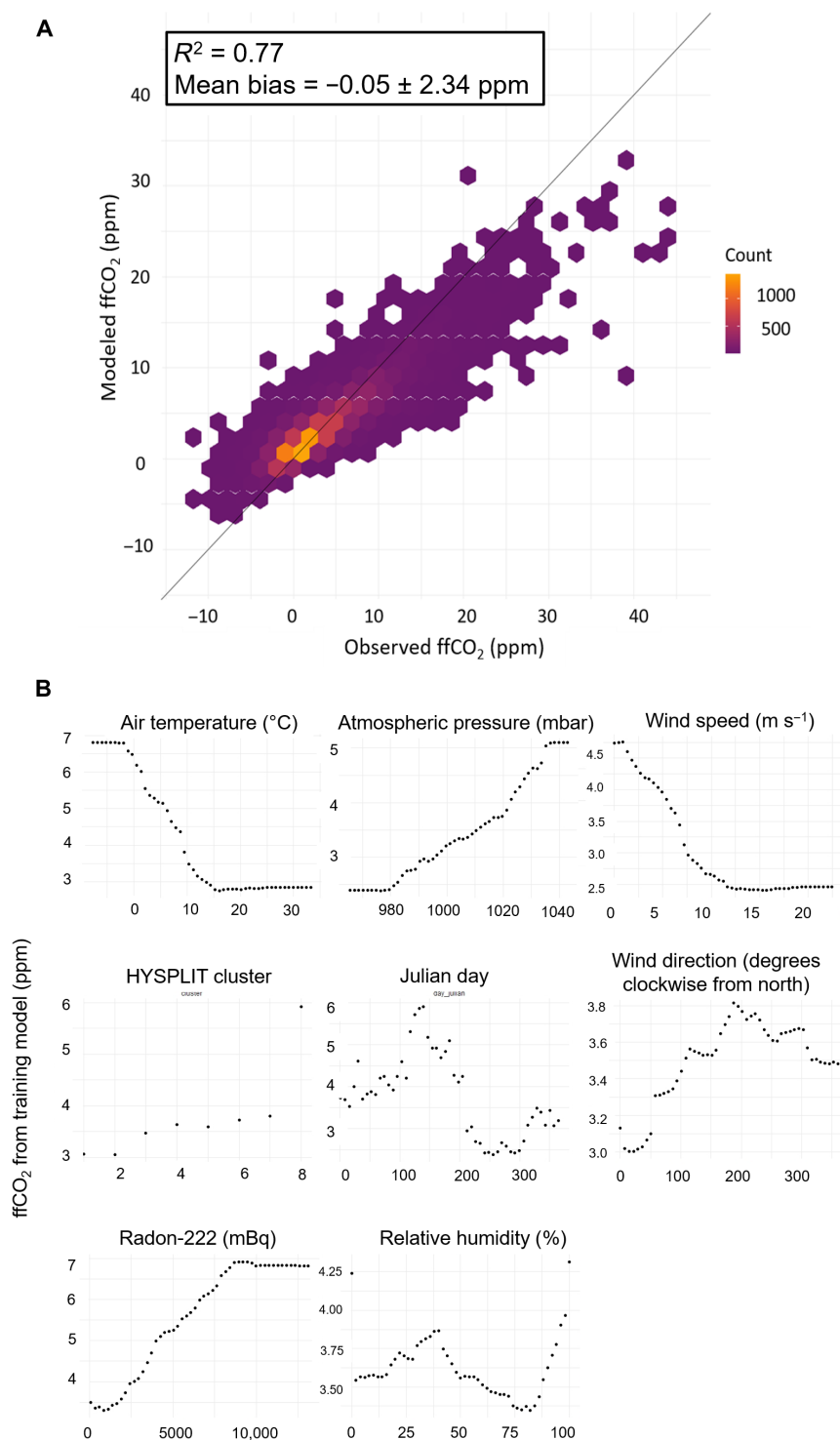


Fig. 4. Evaluation of random forest machine learning model. (A) Scatter plot of hourly observed versus modeled ffCO₂ from the random forest model (2010–2019 only), showing the mean of the differences $\pm 1\sigma$ SD. The plot is created using data from the model test set only, which are withheld from model training. The black line represents a 1:1 relationship. Observed ffCO₂ is calculated using the APO approach (see Materials and Methods). The model underestimates the true range of variability of the APO-based ffCO₂ but generally performs well. A histogram of the differences is shown in fig. S4. (B) Partial dependence plot of the key independent variables of the trained random forest model. The plots show the relationship between each independent variable and modeled ffCO₂ (from the trained model) and therefore provide insight into how variables are being used in the predictive model (39). See fig. S5 for the HYSPLIT cluster key.

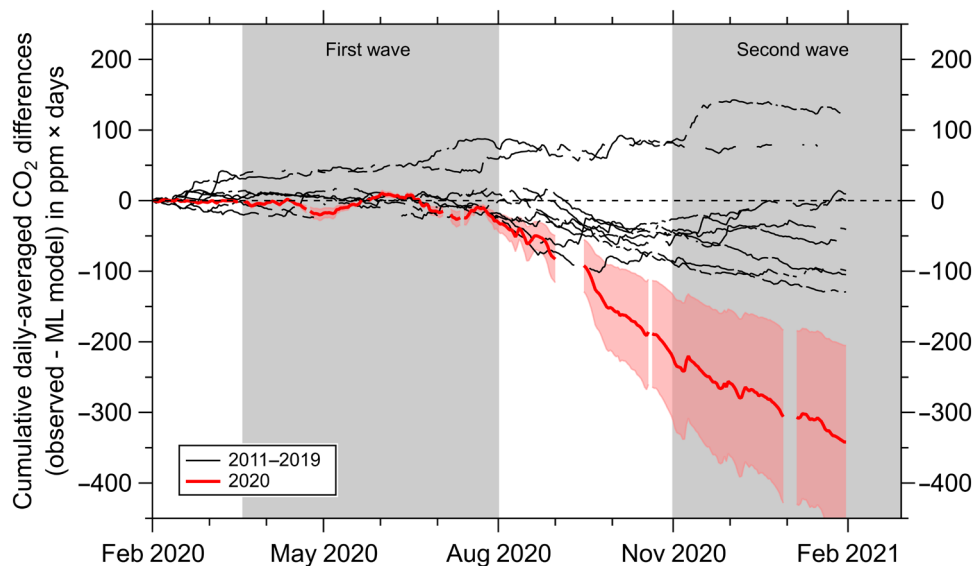


Fig. 5. Results of the random forest prediction using atmospheric CO₂ data. The years 2011–2019 are shown by the black lines. The year 2020 is shown by the red line, with $\pm 40\%$ uncertainty of the machine learning prediction indicated by the red shading. The uncertainty of the daily CO₂ observations themselves is not shown, since this is extremely small (typical hourly uncertainties are less than ± 0.1 ppm).

DISCUSSION

We have used APO, a derived tracer that can separate natural and anthropogenic contributions to regional atmospheric CO₂ variations, to quantify the reduction in ffCO₂ associated with the COVID-19 lockdowns in the United Kingdom during 2020–2021. Detection of COVID-19 signals is only possible owing to the high frequency of the APO measurements (which produce continuous, hourly data) and when combined with a machine learning algorithm that accounts for the influence of atmospheric transport variability on the observed ffCO₂. It should be noted that although the machine learning algorithm allows the comparison of ffCO₂ in the atmosphere between different years, it does not account for how atmospheric mixing translates CO₂ fluxes into atmospheric mole fractions. We also acknowledge that training a statistical model for predictive purposes using trending data is likely to be inappropriate unless the trend is first removed. For our study, while there has been a decreasing trend in U.K. ffCO₂ emissions since 2010, this trend is not visible in the APO-based ffCO₂ at WAO (as shown in Fig. 2), due to the dominating influence of atmospheric transport variability on the ffCO₂ signal compared to a relatively smaller decreasing trend; thus, in our case, there is no requirement to account for any ffCO₂ emissions trend before training.

It might be feasible to replicate our machine learning–based analysis using a discrete, low-frequency ffCO₂ tracer, but only with much larger uncertainty, since the random forest algorithm relies on having a considerable number of high-frequency (e.g., hourly) ffCO₂ values. In this case, the timely detection of COVID-19 signals at WAO is only made possible by the availability of a continuous ffCO₂ tracer combined with a method to remove the effects of atmospheric transport on the ffCO₂ signal, such as the machine learning algorithm we have applied here.

Our COVID-19 signal, detected directly from atmospheric measurements as a mean decrease of 0.7 ppm in daily observed ffCO₂ from March to December 2020, is in broad agreement with three U.K. bottom-up emissions estimates, based on indirect energy and

activity data. We refer to atmospheric measurement-based top-down estimates as direct and bottom-up estimates as indirect, because what matters from a climate change perspective is the change in radiative forcing in the atmosphere, caused by changes in atmospheric greenhouse gas concentrations, which top-down methods are able to measure directly. Our APO-based analysis is able to resolve the two U.K. lockdown periods individually, which are separated by a period of recovery during the summer of 2020 when lockdown restrictions were eased. This APO analysis, using data from a single U.K. measurement station, indicates that a network of continuous APO measurement sites would have strong potential for providing top-down estimates of ffCO₂ emissions at regional scales, which corroborates the results of a recent modeling analysis (42). Furthermore, since the WAO APO data are measured and calibrated in situ and in real time, APO data could be highly beneficial in providing timely top-down ffCO₂ estimates in the future.

The APO network of stations is currently sparse with few measurement sites ideally situated to capture anthropogenic emissions signals. Thus, using APO as a tool for top-down ffCO₂ emissions quantification efforts at scale will require investment in precise and accurate atmospheric O₂ and CO₂ measurements, which are technically challenging, and improved knowledge of R_{APO} from emissions inventories.

MATERIALS AND METHODS

Measurements and calculations of CO₂, O₂, APO, and ffCO₂[APO] at the WAO, United Kingdom

WAO is situated on the north Norfolk coast (53°N) in a rural part of the United Kingdom. Atmospheric O₂ is measured every two minutes with a Sable Systems International Inc. “Oxzilla II” electrochemical fuel cell analyzer, and CO₂ is measured with a Siemens Corporation “Ultramat 6E” nondispersive infrared analyzer (43). O₂ measurements are reported on the Scripps Institution of Oceanography, USA O₂ scale (44) and CO₂ measurements are reported

on the World Meteorological Organization CO₂ X2007 scale (45). Atmospheric O₂ measurements are reported as δ(O₂/N₂) ratios in per meg units, rather than mole fractions, because O₂ is not a trace gas, and its mole fraction is therefore affected by small changes in other gases, such as CO₂ (36).

APO is calculated from measurements of atmospheric O₂ and CO₂ (33) according to

$$\text{APO} \approx \delta(\text{O}_2/\text{N}_2) + 1.1/0.2094 \times (350 - \text{CO}_2) \quad (2)$$

where the value of 1.1 denotes the mean –O₂:CO₂ molar ratio of terrestrial biosphere-atmosphere exchange (34), 0.2094 is the standard mole fraction of O₂ in dry air, and 350 is an arbitrary reference value for CO₂ in parts per million, based on the CO₂ mole fraction of cylinders that define zero on the Scripps Institution of Oceanography calibration scale. By summing O₂ and CO₂ observations in this way, APO becomes invariant to terrestrial biosphere exchange by definition. Variability in APO therefore reflects mostly ocean-atmosphere exchange (on seasonal and long-term time scales) and fossil fuel combustion (on short- and long-term time scales). Our calculation of APO is approximate because it does not account for the influences of CH₄ and CO (33), which have a negligible effect on APO at WAO.

The stoichiometry of O₂ and CO₂ exchange between the atmosphere and terrestrial biosphere has been shown to be robust on a global scale. Atmospheric O₂ and CO₂ observations from a range of locations sampling well-mixed tropospheric air have consistently found –O₂:CO₂ ratios to be within 1.10 ± 0.05, with very little temporal or spatial variability observed. An extensive range of surveys have found that all major organic pools on land have –O₂:CO₂ values ranging from 1.0 to 1.2 (46), while the –O₂:CO₂ exchange ratio of unpolluted air at WAO has previously been shown to lie within 1.10 ± 0.05 (47).

We calculate the regional fossil fuel component of atmospheric CO₂ mole fractions (ffCO₂) in units of parts per million at WAO using hourly averaged APO data as shown in Eq. 1. APO_{BL}, the hourly APO “baseline” values (i.e., values that are representative of the well-mixed troposphere of the wider region), was determined using a statistical baseline fitting method (48), because there is not presently a suitable up-wind station with APO data from which we can obtain a measured baseline; thus, we sometimes obtain negative ffCO₂ values. R_{APO} is the hourly APO:CO₂ combustion ratio for fossil fuel emissions, calculated by converting –O₂:CO₂ molar ratios for fossil fuel combustion (α_F) from the “COFFEE” (CO₂ release and oxygen uptake from fossil fuel emission estimate) database (35) into –APO:CO₂ molar ratios according to

$$R_{\text{APO}} = \alpha_{\text{F}} - 1.1 \quad (3)$$

where 1.1 is the –O₂:CO₂ molar ratio of terrestrial biosphere-atmosphere exchange as mentioned above. For our COVID analysis, we use a mean R_{APO} value at WAO of 0.37, obtained from STILT (17), which is run using gridded α_F values from the COFFEE database.

APO has historically been used to attribute variations in atmospheric O₂ due to oceanic processes [e.g., (33, 37, 49, 50)]; however, we estimate the effects of oceanic influences on our APO ffCO₂ estimates to be minimal, because oceanic influences on APO mostly occur on seasonal or longer time scales, with short-term variations in APO dominated by fossil fuel emissions, and because the influence of short-term ocean-related variability on O₂ at WAO has been

estimated to be only ~6% (51); thus, the majority of oceanic influences in APO are incorporated into APO_{BL} and are excluded from ffCO₂[APO].

Calculation of hourly ffCO₂[APO] uncertainties

We include here information about how to account for the sources of hourly uncertainty associated with each term in Eq. 1 as follows. Uncertainty in the APO data is calculated from typical ±1σ SDs in the hourly CO₂ and O₂ measurements at WAO during stable atmospheric conditions with low natural variability. In this manner, we incorporate uncertainty in analyzer performance and routine calibrations, as well as any uncertainty introduced by the measurement system, in pumping outside air from the tower, drying it, and passing it to the O₂ and CO₂ analyzers. Our APO data uncertainty also incorporates an estimate of the uncertainty associated with the –O₂:CO₂ molar ratio of terrestrial biosphere-atmosphere exchange, for which we use ±0.05 (34). We obtain a typical hourly APO uncertainty of approximately ±2 per meg or ±0.4 ppm equivalent units. Using a mean short-term range of observed APO variability of 49 per meg, the relative APO measurement uncertainty is therefore ±4.1%. The APO range is used rather than the mean APO value because APO is a quantity defined relative to an arbitrary reference.

Determination of a “true” atmospheric baseline concentration is complicated and poorly defined for almost all atmospheric greenhouse gases and pollutants. Ideally, one would use other stations to directly measure an appropriate baseline from clean-air/up-wind sites; however, it is rare that atmospheric measurement networks are sufficiently dense or optimized to allow this method of baseline determination. Since there is currently no upwind O₂ measurement site for WAO, we use “rfbaseline,” a statistical baseline estimation technique with a smoothing window of approximately 1 week, which is fitted to the WAO APO data themselves and which tracks values deemed to be unaffected by local influences (48). Uncertainty of the baseline is estimated by recalculating ffCO₂[APO] with varying “stiffness” in the fitting routine, using span values of 0.03, 0.06, and 0.12, representing smoothing windows of approximately 1, 2, and 4 weeks, respectively. The choice of baseline stiffness mostly affects the magnitude of the ffCO₂ obtained, not the variability, which is determined by variability in the APO data. We estimate an APO_{BL} uncertainty of ±2.81 per meg, or ±0.6 ppm equivalent units, calculated from the mean differences between the 1- and 4-week smoothed baselines compared to the 2-week smoothed baseline. As a relative uncertainty, we find APO_{BL} to be ±28%, based on a mean short-term range of APO_{BL} variability of 10 per meg.

Uncertainty in the gridded fossil fuel emission ratio estimates of R_{APO} are not provided in COFFEE; instead, we use the ±1σ SD (i.e., the variability) around the mean STILT-COFFEE R_{APO} value of 0.37 as a proxy for uncertainty in R_{APO} at WAO, giving an absolute value of ±0.11. The relative uncertainty in R_{APO} is therefore ±31% (±1σ divided by the mean, multiplied by 100). The uncertainties of both α_F and the O₂:CO₂ molar ratio of terrestrial biosphere-atmosphere exchange are included in our total R_{APO} uncertainty as per Eq. 3.

Converting the absolute uncertainty of each term in Eq. 1 into relative uncertainties allows the total hourly ffCO₂[APO] uncertainty to be calculated using the rules of error propagation. We estimate a mean hourly ffCO₂[APO] uncertainty of ±42%, which is dominated by our estimate of uncertainty in R_{APO}, followed by our

estimate of uncertainty in APO_{BL} . At WAO, during the period 2010–2021, the mean $\text{ffCO}_2[\text{APO}]$ detected is 3.4 ± 1.4 ppm. We reiterate that our estimate of R_{APO} uncertainty, while being the largest term, is only a proxy for the actual uncertainty owing to a lack of available information and that future work on using APO to quantify ffCO_2 should focus on refining R_{APO} values and their uncertainties. Uncertainty in APO_{BL} can likely be reduced with denser networks of APO data, which may allow APO_{BL} to be calculated from other sites, instead of having to rely on statistical fitting methods that likely have a higher uncertainty.

Machine learning analysis and COVID-19 signal detection

To account for the influence of atmospheric transport on our $\text{ffCO}_2[\text{APO}]$ dataset, we use the “rmweather” R package (version 0.1.51) (52, 53), which has previously been used with air pollution datasets (39, 53, 54). The rmweather package uses random forest, an ensemble decision tree machine learning method (55) that splits observations using a binary algorithm into two homologous groups, known as branches, repeating the process until the “tree” is fully grown [“node purity” is achieved (54)]. Decision trees are prone to overfitting (56), but random forest mitigates this by growing many individual decision trees from a training set in a process called bagging (bootstrap aggregation), which creates a forest of decorrelated trees, since each has been grown on different subsets of the training set.

Using rmweather, we train a random forest model of 300 trees at WAO for the period 2010–2019 using 10 independent variables: hourly meteorological observations (wind speed, wind direction, air temperature, relative humidity, and atmospheric pressure), which are measured in situ at WAO; temporal factors (day of the year, day of the week, and hour of the day); hourly atmospheric radon-222 activity, a tracer for atmospheric mixing that has been measured at WAO since April 2018 using an Australian Nuclear Science and Technology Organisation monitor (57); and hourly 24-hour-long HYSPLIT (<https://ready.noaa.gov/HYSPLIT.php>) model (58) backward-run trajectories, clustered into eight groups (see fig. S5) using k -means clustering and the openair package in R (59) (<https://davidcarslaw.github.io/openair/>). For the dependent variable, we use hourly $\text{ffCO}_2[\text{APO}]$, calculated using Eq. 1. Meteorological data are cross-checked against an independent (but colocated) dataset operated by the U.K. Met Office as a quality control measure. The training set consists of 80% of the data, with 20% reserved for testing.

Performance of the trained model was assessed for bias and goodness of fit, as shown in Fig. 4. We use the model, trained on the 2010–2019 data to predict the counterfactual ffCO_2 that would have been observed at WAO during the time period 1 February 2020 to 31 January 2021, if the COVID-19 pandemic had not occurred. This counterfactual prediction is then compared to the observed $\text{ffCO}_2[\text{APO}]$ values over the same time period to estimate the impact of COVID-19 lockdown measures on ffCO_2 at WAO. We calculate the relative percent change in emissions, E , by taking the ratio of the cumulative $\text{ffCO}_2[\text{APO}]$ and cumulative ffCO_2 counterfactual signals shown in Fig. 3B according to

$$E = \left(1 - \frac{\text{ffCO}_2[\text{APO}]_{\text{cumulative}}}{\text{ffCO}_2[\text{counterfactual}]_{\text{cumulative}}} \right) \times 100\% \quad (4)$$

As mentioned in the “Analysis of uncertainties” section (see Results), the uncertainty range we report on the COVID-19 ffCO_2 relative emissions reductions shown in Fig. 3B is based solely on the

uncertainty associated with the machine learning algorithm (calculated from the difference in model performance between training and predictive results), since the APO-based uncertainties cancel.

Comparison with bottom-up inventory estimates

We compare our COVID-19 signal detected from $\text{ffCO}_2[\text{APO}]$ data to three bottom-up inventory estimates for the United Kingdom, which quantify COVID lockdown emissions reductions by comparing 2020 emissions to those from 2019. These are the following: inland energy consumption statistics from the U.K. BEIS, which we convert to CO_2 emissions estimates in units of megatons using coal, gas, and oil conversion factors and by tuning to annual emissions from previous years; an updated version (March 2021) of the UEA estimate (3), based on a combination of energy, activity, and policy data; and an estimate from Carbon Monitor, based on fuel consumption and activity data (4). Only the Carbon Monitor emissions have daily resolution; the UEA and U.K. BEIS estimates are monthly. We exclude the 29 February from all estimates (bottom-up and top-down) to enable a comparison of leap years to nonleap years. At the time of publication, only the UEA estimate fully includes emissions from IAS.

SUPPLEMENTARY MATERIALS

Supplementary material for this article is available at <https://science.org/doi/10.1126/sciadv.abl9250>

REFERENCES AND NOTES

1. IEA, International Energy Agency World Energy Outlook 2008 (2008), chap. 8, pp. 179–193.
2. G. P. Peters, C. Le Quéré, R. M. Andrew, J. G. Canadell, P. Friedlingstein, T. Ilyina, R. B. Jackson, F. Joos, J. I. Korsbakken, G. A. McKinley, S. Sitch, P. Tans, Towards real-time verification of CO_2 emissions. *Nat. Clim. Chang.* **7**, 848–850 (2017).
3. C. Le Quéré, R. B. Jackson, M. W. Jones, A. J. P. Smith, S. Abernethy, R. M. Andrew, A. J. De-Gol, D. R. Willis, Y. Shan, J. G. Canadell, P. Friedlingstein, F. Creutzig, G. P. Peters, Temporary reduction in daily global CO_2 emissions during the COVID-19 forced confinement. *Nat. Clim. Chang.* **10**, 647–653 (2020).
4. Z. Liu, P. Ciais, Z. Deng, R. Lei, S. J. Davis, S. Feng, B. Zheng, D. Cui, X. Dou, B. Zhu, R. Guo, P. Ke, T. Sun, C. Lu, P. He, Y. Wang, X. Yue, Y. Wang, Y. Lei, H. Zhou, Z. Cai, Y. Wu, R. Guo, T. Han, J. Xue, O. Boucher, E. Boucher, F. Chevallier, K. Tanaka, Y. Wei, H. Zhong, C. Kang, N. Zhang, B. Chen, F. Xi, M. Liu, F. M. Bréon, Y. Lu, Q. Zhang, D. Guan, P. Gong, D. M. Kammen, K. He, H. J. Schellnhuber, Near-real-time monitoring of global CO_2 emissions reveals the effects of the COVID-19 pandemic. *Nat. Commun.* **11**, 5172 (2020).
5. P. Friedlingstein, M. O’Sullivan, M. W. Jones, R. M. Andrew, J. Hauck, A. Olsen, G. P. Peters, W. Peters, J. Pongratz, S. Sitch, C. Le Quéré, J. G. Canadell, P. Ciais, R. B. Jackson, S. Alin, L. E. O. C. Aragão, A. Arneeth, V. Arora, N. R. Bates, M. Becker, A. Benoit-Cattin, H. C. Bittig, L. Bopp, S. Bultan, N. Chandra, F. Chevallier, L. P. Chini, W. Evans, L. Florentie, P. M. Forster, T. Gasser, M. Gehlen, D. Gilfillan, T. Gkritzalis, L. Gregor, N. Gruber, I. Harris, K. Hartung, V. Haverd, R. A. Houghton, T. Ilyina, A. K. Jain, E. Joetzier, K. Kadono, E. Kato, V. Kitidis, J. I. Korsbakken, P. Landschützer, N. Lefèvre, A. Lenton, S. Lienert, Z. Liu, D. Lombardozzi, G. Marland, N. Metz, D. R. Munro, J. E. M. S. Nabel, S. I. Nakaoka, Y. Niwa, K. O’Brien, T. Ono, P. I. Palmer, D. Pierrot, B. Poulter, L. Resplandy, E. Robertson, C. Rödenbeck, J. Schwinger, R. Séférian, I. Skjelvan, A. J. P. Smith, A. J. Sutton, T. Tanhua, P. P. Tans, H. Tian, B. Tilbrook, G. van der Werf, N. Vuichard, A. P. Walker, R. Wanninkhof, A. J. Watson, D. Willis, A. J. Wiltshire, W. Yuan, X. Yue, S. Zaehle, Global carbon budget 2020. *Earth Syst. Sci. Data* **12**, 3269–3340 (2020).
6. P. M. Forster, H. I. Forster, M. J. Evans, M. J. Gidden, C. D. Jones, C. A. Keller, R. D. Lamboll, C. L. Quéré, J. Rogelj, D. Rosen, C.-F. Schleussner, T. B. Richardson, C. J. Smith, S. T. Turnock, Current and future global climate impacts resulting from COVID-19. *Nat. Clim. Chang.* **10**, 913–919 (2020).
7. Y. Rybarczyk, R. Zalakeviciute, Assessing the COVID-19 impact on air quality: A machine learning approach. *Geophys. Res. Lett.* **48**, e2020GL091202 (2021).
8. H. Petetin, D. Bowdalo, A. Soret, M. Guevara, O. Jorba, K. Serradell, C. Pérez García-Pando, Meteorology-normalized impact of the COVID-19 lockdown upon NO_2 pollution in Spain. *Atmos. Chem. Phys.* **20**, 11119–11141 (2020).
9. H. F. Dacre, L. M. Western, D. Say, S. O’Doherty, T. Arnold, C. Rennick, E. Hawkins, Detectability of COVID-19 global emissions reductions in local CO_2 concentration measurements. *Environ. Res. Lett.* **16**, 094043 (2021).

10. European Commission, Joint Research Centre, Ciais, P., Towards a European operational observing system to monitor fossil :CO₂ emissions : final report from the expert group (Publications Office, 2016); <https://data.europa.eu/doi/10.2788/350433>.
11. S. Basu, S. J. Lehman, J. B. Miller, A. E. Andrews, C. Sweeney, K. R. Gurney, X. Xu, J. Southon, P. P. Tans, Estimating US fossil fuel CO₂ emissions from measurements of ¹⁴C in atmospheric CO₂. *Proc. Natl. Acad. Sci. U.S.A.* **117**, 13300–13307 (2020).
12. E. Calvo Buendía, S. Guendehou, B. Limmeechokchai, R. Pipatti, Y. Rojas, R. Sturgiss, K. Tanabe, T. Wirth, 2019 refinement to the 2006 IPCC guidelines for National Greenhouse Gas Inventories, in *IPCC 2019* (IPCC, 2019).
13. E. Nisbet, R. Weiss, Top-down versus bottom-up. *Science* **328**, 1241–1243 (2010).
14. K. R. Gurney, J. Huang, K. Coltin, Bias present in US federal agency power plant CO₂ emissions data and implications for the US clean power plan. *Environ. Res. Lett.* **11**, 064005 (2016).
15. P. Ciais, A. J. Dolman, A. Bombelli, R. Duren, A. Peregón, P. J. Rayner, C. Miller, N. Gobron, G. Kinderman, G. Marland, N. Gruber, F. Chevallier, R. J. Andres, G. Balsamo, L. Bopp, F. M. Bréon, G. Broquet, R. Dargaville, T. J. Battin, A. Borges, H. Bovensmann, M. Buchwitz, J. Butler, J. G. Canadell, R. B. Cook, R. DeFries, R. Engelen, K. R. Gurney, C. Heinze, M. Heimann, A. Held, M. Henry, B. Law, S. Luysaert, J. Miller, T. Moriyama, C. Moulin, R. B. Myneni, C. Nussli, M. Obersteiner, D. Ojima, Y. Pan, J. D. Paris, S. L. Piao, B. Pouler, S. Plummer, S. Quegan, P. Raymond, M. Reichstein, L. Rivier, C. Sabine, D. Schimel, O. Tarasova, R. Valentini, R. Wang, G. van der Werf, D. Wickland, M. Williams, C. Zehner, Current systematic carbon-cycle observations and the need for implementing a policy-relevant carbon observing system. *Biogeosciences* **11**, 3547–3602 (2014).
16. K. M. Stanley, A. Grant, S. O'Doherty, D. Young, A. J. Manning, A. R. Stavert, T. G. Spain, P. K. Salameh, C. M. Harth, P. G. Simmonds, W. T. Sturges, D. E. Oram, R. G. Derwent, Greenhouse gas measurements from a UK network of tall towers: Technical description and first results. *Atmos. Meas. Tech.* **11**, 1437–1458 (2018).
17. J. C. Lin, C. Gerbig, S. C. Wofsy, A. E. Andrews, B. C. Daube, K. J. Davis, C. A. Grainger, A near-field tool for simulating the upstream influence of atmospheric observations: The stochastic time-inverted Lagrangian transport (STILT) model. *J. Geophys. Res. Atmos.* **108**, 4493 (2003).
18. M. Buchwitz, M. Reuter, S. Noël, K. Bramstedt, O. Schneising, M. Hilker, B. Fuentes Andrade, H. Bovensmann, J. P. Burrows, A. di Noia, H. Boesch, L. Wu, J. Landgraf, I. Aben, C. Retscher, C. W. O'Dell, D. Crisp, Can a regional-scale reduction of atmospheric CO₂ during the COVID-19 pandemic be detected from space? A case study for East China using satellite XCO₂ retrievals. *Atmos. Meas. Tech.* **14**, 2141–2166 (2021).
19. Y. Tohjima, P. K. Patra, Y. Niwa, H. Mukai, M. Sasaki, T. Machida, Detection of fossil-fuel CO₂ plummets in China due to COVID-19 by observation at Hateruma. *Sci. Rep.* **10**, 18688 (2020).
20. B. Zheng, G. Geng, P. Ciais, S. J. Davis, R. V. Martin, J. Meng, N. Wu, F. Chevallier, G. Broquet, F. Boersma, R. van der A, J. Lin, D. Guan, Y. Lei, K. He, Q. Zhang, Satellite-based estimates of decline and rebound in China's CO₂ emissions during COVID-19 pandemic. *Sci. Adv.* **6**, eabd4998 (2020).
21. A. J. Turner, J. Kim, H. Fitzmaurice, C. Newman, K. Worthington, K. Chan, P. J. Woolridge, P. Köehler, C. Frankenberg, R. C. Cohen, Observed impacts of COVID-19 on urban CO₂ emissions. *Geophys. Res. Lett.* **47**, e2020GL090037 (2020).
22. I. Levin, S. Hammer, E. Eichelmann, F. R. Vogel, Verification of greenhouse gas emission reductions: The prospect of atmospheric monitoring in polluted areas. *Philos. Trans. R. Soc. A* **369**, 1906–1924 (2011).
23. I. Levin, B. Kromer, M. Schmidt, H. Sartorius, A novel approach for independent budgeting of fossil fuel CO₂ over Europe by ¹⁴CO₂ observations. *Geophys. Res. Lett.* **30**, 2194 (2003).
24. S. N. Vardag, C. Gerbig, G. Janssens-Maenhout, I. Levin, Estimation of continuous anthropogenic CO₂: Model-based evaluation of CO₂, CO, δ¹³C(CO₂) and Δ¹⁴C(CO₂) tracer methods. *Atmospheric Chem. Phys.* **15**, 12705–12729 (2015).
25. F. R. Vogel, S. Hammer, A. Steinhof, B. Kromer, I. Levin, Implication of weekly and diurnal ¹⁴C calibration on hourly estimates of CO-based fossil fuel CO₂ at a moderately polluted site in southwestern Germany. *Tellus B*, 512–520 (2010).
26. J. Turnbull, P. Rayner, J. Miller, T. Naegler, P. Ciais, A. Cozic, On the use of ¹⁴CO₂ as a tracer for fossil fuel CO₂: Quantifying uncertainties using an atmospheric transport model. *J. Geophys. Res. Atmos.* **114**, 1–14 (2009).
27. J. C. Turnbull, J. B. Miller, S. J. Lehman, P. P. Tans, R. J. Sparks, J. Southon, Comparison of ¹⁴CO₂, CO, and SF₆ as tracers for recently added fossil fuel CO₂ in the atmosphere and implications for biological CO₂ exchange. *Geophys. Res. Lett.* **33**, 1–14 (2006).
28. J. C. Turnbull, C. Sweeney, A. Karion, T. Newberger, S. J. Lehman, P. P. Tans, K. J. Davis, T. Lauvaux, N. L. Miles, S. J. Richardson, M. O. Cambaliza, P. B. Shepson, K. Gurney, R. Patarasuk, I. Razlivanov, Toward quantification and source sector identification of fossil fuel CO₂ emissions from an urban area: Results from the INFLUX experiment. *J. Geophys. Res. Atmos.* **120**, 292–312 (2015).
29. J. B. Miller, S. J. Lehman, K. R. Verhulst, C. E. Miller, R. M. Duren, V. Yadav, S. Newman, C. D. Sloop, Large and seasonally varying biospheric CO₂ fluxes in the Los Angeles megacity revealed by atmospheric radiocarbon. *Proc. Natl. Acad. Sci. U.S.A.* **117**, 26681–26687 (2020).
30. I. Levin, U. Karstens, M. Eritt, F. Maier, S. Arnold, D. Rzesanke, S. Hammer, M. Ramonet, G. Vitková, S. Conil, M. Helliasz, D. Kubistin, M. Lindauer, A dedicated flask sampling strategy developed for integrated carbon observation system (ICOS) stations based on CO₂ and CO measurements and stochastic time-inverted Lagrangian transport (STILT) footprint modelling. *Atmospheric Chem. Phys.* **20**, 11161–11180 (2020).
31. H. D. Graven, N. Gruber, Continental-scale enrichment of atmospheric ¹⁴CO₂ from the nuclear power industry: Potential impact on the estimation of fossil fuel-derived CO₂. *Atmospheric Chem. Phys.* **11**, 12339–12349 (2011).
32. A. Wenger, K. Pugsley, S. O'Doherty, M. Rigby, A. J. Manning, M. F. Lunt, E. D. White, Atmospheric radiocarbon measurements to quantify CO₂ emissions in the UK from 2014 to 2015. *Atmospheric Chem. Phys.* **19**, 14057–14070 (2019).
33. B. B. Stephens, R. F. Keeling, M. Heimann, K. D. Six, R. Murnane, K. Caldeira, Testing global ocean carbon cycle models using measurements of atmospheric O₂ and CO₂ concentration. *Global Biogeochem. Cycles* **12**, 213–230 (1998).
34. J. P. Severinghaus, "Studies of the terrestrial O₂ and carbon cycles in sand dunes gases and in Biosphere 2," thesis, Columbia University (1995).
35. J. Steinbach, C. Gerbig, C. Rödenbeck, U. Karstens, C. Minejima, H. Mukai, The CO₂ release and oxygen uptake from fossil fuel emission estimate (COFFEE) dataset: Effects from varying oxidative ratios. *Atmospheric Chem. Phys.* **11**, 6855–6870 (2011).
36. R. F. Keeling, A. C. Manning, E. M. McEvoy, S. R. Shertz, Methods for measuring changes in atmospheric O₂ concentration and their application in southern hemisphere air. *J. Geophys. Res. Atmos.* **103**, 3381–3397 (1998).
37. P. A. Pickers, A. C. Manning, W. T. Sturges, C. le Quéré, S. E. Mikaloff Fletcher, P. A. Wilson, A. J. Etchells, In situ measurements of atmospheric O₂ and CO₂ reveal an unexpected O₂ signal over the tropical Atlantic Ocean. *Global Biogeochem. Cycles* **31**, 1289–1305 (2017).
38. E. A. Kozlova, A. C. Manning, Y. Kisilyakhov, T. Seifert, M. Heimann, Seasonal, synoptic, and diurnal-scale variability of biogeochemical trace gases and O₂ from a 300m tall tower in central Siberia. *Global Biogeochem. Cycles* **22**, GB4020 (2008).
39. S. K. Grange, D. C. Carslaw, Using meteorological normalisation to detect interventions in air quality time series. *Sci. Total Environ.* **653**, 578–588 (2019).
40. R. A. Betts, "Analysis: What impact will the coronavirus pandemic have on atmospheric CO₂?" www.carbonbrief.org/analysis-what-impact-will-the-coronavirus-pandemic-have-on-atmospheric-co2 (last accessed 26th October 2020).
41. C. L. Quéré, G. P. Peters, P. Friedlingstein, R. M. Andrew, J. G. Canadell, S. J. Davis, R. B. Jackson, M. W. Jones, Fossil CO₂ emissions in the post-COVID-19 era. *Nat. Clim. Chang.* **11**, 197–199 (2021).
42. F. Chevallier, C. W. partners, "D4.4 sampling strategy for additional tracers" (CHE Consortium, 2021); www.che-project.eu/node/243.
43. P. A. Wilson, University of East Anglia, Norwich, UK (2013); www.cramlab.uea.ac.uk/Documents/Wilson_Phil_PhD_Thesis_2013.pdf.
44. R. F. Keeling, A. C. Manning, W. J. Paplawsky, A. C. Cox, On the long-term stability of reference gases for atmospheric O₂/N₂ and CO₂ measurements. *Tellus Series B-Chemical and Physical Meteorology* **59**, 3–14 (2007).
45. C. L. Zhao, P. P. Tans, Estimating uncertainty of the WMO mole fraction scale for carbon dioxide in air. *Journal of Geophysical Research-Atmospheres* **111**, 10 (2006).
46. R. F. Keeling, A. C. Manning, in *Treatise on Geochemistry (Second Edition)*, H. D. Holland, K. K. Turekian, Eds. (Elsevier, 2014), pp. 385–404.
47. P. A. Pickers, School of Environmental Sciences, University of East Anglia (2016); www.cramlab.uea.ac.uk/Documents/Pickers_Penelope_PhD_Thesis_2016.pdf.
48. A. F. Ruckstuhl, S. Henne, S. Reimann, M. Steinbacher, M. K. Vollmer, S. O'Doherty, B. Buchmann, C. Hueglin, Robust extraction of baseline signal of atmospheric trace species using local regression. *Atmos. Meas. Tech.* **5**, 2613–2624 (2012).
49. Y. A. Eddebbbar, M. C. Long, L. Resplandy, C. Rödenbeck, K. B. Rodgers, M. Manizza, R. F. Keeling, Impacts of ENSO on air-sea oxygen exchange: Observations and mechanisms. *Global Biogeochem. Cycles* **31**, 901–921 (2017).
50. Y. Tohjima, Y. Terao, H. Mukai, T. MacHida, Y. Nojiri, S. Maksyutov, ENSO-related variability in latitudinal distribution of annual mean atmospheric potential oxygen (APO) in the equatorial western Pacific. *Tellus B* **67**, 25869 (2015).
51. B. Kuijpers, Wageningen University (2018); <http://edepot.wur.nl/455747>.
52. R Core Team, (R Foundation for Statistical Computing, 2018).
53. S. K. Grange, *rmweather*: Tools to conduct meteorological normalisation on air quality data. R package version 0.1.51. (2018).
54. S. K. Grange, D. C. Carslaw, A. C. Lewis, E. Boleti, C. Hueglin, Random forest meteorological normalisation models for Swiss PM10 trend analysis. *Atmospheric Chem. Phys.* **18**, 6223–6239 (2018).
55. L. Breiman, Random forests. *Mach. Learn.* **45**, 5–32 (2001).
56. S. B. Kotsiantis, Decision trees: A recent overview. *Artif. Intell. Rev.* **39**, 261–283 (2013).
57. D. Schmidthusen, S. Chambers, B. Fischer, S. Gilge, J. Hattaka, V. Kazan, R. Neubert, J. Paatero, M. Ramonet, C. Schlosser, S. Schmid, A. Vermeulen, I. Levin, A European-wide ²²²Rn and ²²²Rn progeny comparison study. *Atmos. Meas. Tech.* **10**, 1299–1312 (2017).

58. A. F. Stein, R. R. Draxler, G. D. Rolph, B. J. B. Stunder, M. D. Cohen, F. Ngan, NOAA'S HYSPLIT atmospheric transport and dispersion modeling system. *Bull. Am. Meteorol. Soc.* **96**, 2059–2077 (2015).
59. D. C. Carslaw, K. Ropkins, Openair — An R package for air quality data analysis. *Environ. Model. Software* **27–28**, 52–61 (2012).

Acknowledgments: We thank M. Patecki, P. Wilson, and T. Barningham for assisting with maintaining the WAO O₂ and CO₂ measurement system during the period 2010–2017.

Funding: Atmospheric O₂ and CO₂ measurements at WAO were funded by the U.K. Natural Environment Research Council (NERC) grants NE/F005733/1, NE/I013342/1, NE/I02934X/1, QUEST010005, NE/N016238/1, NE/S004521/1, and NE/R011532/1. The WAO atmospheric O₂ and CO₂ measurements have also been supported by the U.K. National Centre for Atmospheric Science (NCAS) from first December 2013 onward. P.A.P. was supported by a U.K. NERC PhD studentship (NE/K500896/1) from 2012 to 2016. A.C.M. and P.A.P. have received support from the CHE project, funded by the European Union's Horizon 2020 Research and Innovation Programme under grant agreement no. 776186. C.L.Q. received funding from the Royal Society (grant no. RP\R1\191063). I.T.L. received funding from Netherlands Organisation for Scientific Research (016.Veni.171.095 and SH-312-15). P.A.P., A.C.M., and G.L.F. received funding from the NERC project DARE-UK (NE/S004211/1). L.S.F. is

supported by an NERC PhD studentship (NE/L002582/1). **Author contributions:** Conceptualization: P.A.P., A.C.M., and C.L.Q. Methodology: P.A.P., A.C.M., G.L.F., I.T.L., C.G., and L.S.F. Investigation: P.A.P. Visualization: P.A.P. Supervision: P.A.P., A.C.M., C.L.Q., and W.T.S. Writing—original draft: P.A.P. Writing—review and editing: P.A.P., A.C.M., C.L.Q., G.L.F., I.T.L., C.G., L.S.F., and W.T.S. **Competing interests:** The authors declare that they have no competing interests. **Data and materials availability:** Weybourne atmospheric data (O₂, CO₂, radon, and meteorological parameters) are available online at the U.K. Centre for Environmental Data Analysis: <http://data.ceda.ac.uk/badc/weybourne/data/o2>, <http://data.ceda.ac.uk/badc/weybourne/data/co2>, <http://data.ceda.ac.uk/badc/weybourne/data/radon>, and <http://data.ceda.ac.uk/badc/weybourne/data/met>. The U.K. BEIS inland energy consumption data are available at <https://gov.uk/government/statistics/total-energy-section-1-energy-trends>, the emissions of the UEA bottom-up estimate are available at <https://icos-cp.eu/gcp-covid19>, and Carbon Monitor emissions are available at <https://carbonmonitor.org/>. All data needed to evaluate the conclusions in the paper are available from the links provided.

Submitted 13 August 2021

Accepted 7 March 2022

Published 22 April 2022

10.1126/sciadv.abl9250

Novel quantification of regional fossil fuel CO reductions during COVID-19 lockdowns using atmospheric oxygen measurements

Penelope A. PickersAndrew C. ManningCorinne Le QuéréGrant L. ForsterIngrid T. LujikxChristoph GerbigLeigh S. FlemingWilliam T. Sturges

Sci. Adv., 8 (16), eabl9250. • DOI: 10.1126/sciadv.abl9250

View the article online

<https://www.science.org/doi/10.1126/sciadv.abl9250>

Permissions

<https://www.science.org/help/reprints-and-permissions>

Use of this article is subject to the [Terms of service](#)

Science Advances (ISSN) is published by the American Association for the Advancement of Science, 1200 New York Avenue NW, Washington, DC 20005. The title *Science Advances* is a registered trademark of AAAS. Copyright © 2022 The Authors, some rights reserved; exclusive licensee American Association for the Advancement of Science. No claim to original U.S. Government Works. Distributed under a Creative Commons Attribution License 4.0 (CC BY).