

1 **Title:** Balancing Act: Groundwater microbiome's resilience and vulnerability to hydroclimatic
2 extremes

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25 **Abstract**

26 Groundwater health is increasingly threatened by climate change, which alters precipitation patterns,
27 leading to groundwater recharge shifts. These shifts impact subsurface microbial communities, crucial
28 for maintaining ecosystem functions. In this decade-long study of carbonate aquifers, we analyzed 815
29 bacterial 16S rRNA gene datasets, 226 dissolved organic matter (DOM) profiles, 387 metabolomic
30 datasets, and 174 seepage microbiome sequences. Our findings reveal distinct short- and long-term
31 temporal patterns of groundwater microbiomes driven by environmental fluctuations. Microbiomes of
32 hydrologically connected aquifers exhibit lower temporal stability due to stochastic processes and
33 greater susceptibility to surface disturbances, yet they demonstrate remarkable resilience. Conversely,
34 isolated aquifer microbiomes show resistance to short-term changes, governed by deterministic
35 processes, but exhibit reduced stability under prolonged stress. Variability in seepage-associated
36 microorganisms, DOM, and metabolic diversity further drive microbiome dynamics. These findings
37 highlight the dual vulnerability of groundwater systems to acute and chronic pressures, emphasizing
38 the critical need for sustainable management strategies to mitigate the impacts of hydroclimatic
39 extremes.

40 Introduction

41 Groundwater, the largest global reservoir of accessible freshwater, is increasingly jeopardized by the
42 dual pressures of anthropogenic activities and climate change, with widespread implications for water
43 availability, quality, and ecosystem health^{1,2,3}. Future climate scenarios predict profound shifts in
44 groundwater recharge, driven predominantly by changing precipitation patterns and the intensification
45 of hydroclimatic extremes^{2,3}. These shifts threaten groundwater sustainability, particularly in vulnerable
46 carbonate aquifers where water storage and flow dynamics are intricately linked to lithology and
47 hydrological connectivity^{4,5}. Understanding the resilience of groundwater ecosystems and their ability
48 to maintain ecological functions under such pressures is critical for sustainable management strategies^{3,6}.

49 Microbial communities are central to groundwater ecosystem health, playing pivotal roles in
50 biogeochemical cycling, contaminant degradation, and ecosystem resilience^{7,8,9}. Traditionally
51 considered static due to the relatively constant conditions of subsurface environments, this view has
52 persisted largely due to the lack of long-term observational studies. More recent studies challenge this
53 notion by demonstrating that groundwater microbiomes exhibit a dynamic nature, driven by both short-
54 term recharge events^{10,11,12} and longer-term environmental changes¹³. For instance, microbial
55 immigration during recharge can trigger compositional shifts^{12,13}, while prolonged droughts may alter
56 microbial activity through changes in water chemistry¹⁴. Additionally, rain infiltration can mobilize soil
57 microorganisms, transporting them to the groundwater via seepage^{15,16,17,18}. These findings underscore
58 the importance of hydrological connectivity as a key factor influencing microbial dispersal, assembly,
59 and ecosystem resilience^{15,19,20,21}.

60 Hydrological connectivity, shaped by aquifer permeability, flow dynamics, and recharge
61 patterns, governs the flux of surface-derived inputs into groundwater and mediates microbial
62 dispersal^{5,15,19,22}. Hydrologically connected aquifers, often characterized by karstification and higher
63 permeability, are particularly susceptible to surface-derived disturbances such as nutrient influxes,
64 organic contaminants, and pathogens^{4,22,23}. In contrast, hydrologically isolated systems may exhibit
65 greater stability¹¹ but face significant stress under conditions of prolonged environmental change, such

66 as reduced recharge during droughts¹⁴. These contrasting dynamics necessitate a framework for
67 quantifying temporal stability (consistency), resistance (insensitivity to disturbance), and resilience
68 (ability to recovery) of groundwater microbiomes across gradients of hydrological connectivity^{9,21}.

69 While previous studies have focused on microbiome compositional stability across ecosystems,
70 functional stability is also a key indicator of ecosystem health^{13,24,25,26}. To this end, we applied
71 metabolomics techniques to probe system activity in our long-term study²⁷. Untargeted liquid
72 chromatography–high-resolution mass spectrometry (LC-HRMS) can profile environmental meta-
73 metabolomes, comprising metabolites released by microorganisms and anthropogenic activity. This
74 approach is complemented by direct infusion–high-resolution mass spectrometry (DI-HRMS) resolving
75 the full complexity of DOM without chromatographic bias^{28,29,30,31}. This integrated information from
76 both approaches together with microbiome function analysis using PICRUSt2 resolves quantitatively
77 the impact of the organic chemical landscape on microbiome diversity^{32,33}. This allowed us to explore
78 groundwater microbiome stability, variability, and overall ecosystem health.

79 Our work expands directly on discussions highlighting not only the impacts of climate change
80 on groundwater recharge², but also on groundwater quality¹⁴. Building on these foundations, we
81 investigate the critical links between hydrological connectivity and groundwater microbiome stability
82 in carbonate aquifers. Here, we conducted a groundwater study along a hillslope well transect in the
83 Hainich Critical Zone Exploratory (CZE)^{34,35}, a geological setting characterized by alternating
84 limestone and mudstone strata, resulting in varying hydrological connectivity within the carbonate-rock
85 aquifers. Two footslope wells, located in thinner and less connected fractured mudstone-dominated
86 aquifers, are more hydrologically isolated and contain old carbon sources (*e.g.*, > 4,500 years), while
87 other wells in limestone-dominated aquifers having wider fractures exhibit greater hydrological
88 connectivity due to minor karstification and higher permeabilities^{5,36,37}. Previous studies have indicated
89 the importance of connectivity through seepage transport of soil microorganisms into
90 groundwater^{15,16,17,18}. Consequently, the Hainich CZE well network enables comparisons of microbiome
91 variability and stability across aquifers of different levels of hydrological connectivity.

92 Capitalizing on a decade-long groundwater survey comprising 815 bacterial 16S rRNA gene-
93 sequence datasets and accompanying hydrochemical analyses, 226 DOM analyses, and 387
94 metabolomic analyses combined with 174 seepage microbiome sequencing datasets, this study provides
95 the most comprehensive assessment of groundwater microbiome variability and stability, and their
96 potential drivers to date. To better assess the responses of precious groundwater resources to
97 disturbances caused by increasing hydroclimatic extremes, this investigation addresses four key
98 unknowns: (i) the extent to which microbiomes of hydrologically connected vs. more isolated
99 groundwater exhibit similar temporal dynamics and stability, (ii) the means by which assembly
100 mechanisms explain microbiome variability, (iii) the extent to which seepage-associated
101 microorganisms contribute to microbiome variability, and (iv) the means by which alterations in
102 groundwater metabolome/DOM explain microbiome variability. By integrating microbial,
103 hydrochemical, and metabolomic datasets, this study reveals critical insights into the responses of
104 groundwater ecosystems to environmental disturbances and hydroclimatic extremes. These findings
105 have broad implications for predicting groundwater resilience and informing sustainable management
106 strategies under future climate scenarios.

107 **Results**

108 **Hydrological seasonality and long-term variability of groundwater microbiomes**

109 Our 10-year time series revealed distinct and persistent regular patterns of variation with alternating
110 greater and lesser similarity over a 12-month period. While these variations are on a 12-month cycle,
111 we use the descriptive term ‘sinusoidal’ because groundwater systems can integrate a number of
112 potentially seasonally varying signals and lag times. These sinusoidal patterns in the similarity of
113 shallow groundwater (< 100 m depth) microbial communities over time across all wells, despite
114 differences in groundwater hydrochemical parameters and microbiome composition among wells (Fig.
115 1, Supplementary Fig. 1a, b). These patterns demonstrated consistent periodicity, with distinct peaks in
116 similarities approximately every 12 months and minima roughly 6 months after each peak. This
117 behavior was not confined to a single bacterial phylum, as approximately half of the bacterial phyla

118 detected displayed sinusoidal patterns (Supplementary Tab. 1). Similar sinusoidal patterns were
119 observed for hydrochemical properties of the groundwater (Supplementary Fig. 1c).

120 Sinusoidal amplitudes of microbiome community similarities varied among wells. Distance-
121 based redundancy analysis (dbRDA) indicated that hydrological seasons (*i.e.*, early summer, later
122 summer, early winter, and later winter) accounted for one to six percent of the microbiome variation (P
123 < 0.05 ; Supplementary Fig. 2), except in well H53. In contrast, the phylogenetic structure of
124 groundwater microbiomes exhibited weaker or no sinusoidal patterns (Supplementary Fig. 3). Only one
125 to three percent of the variation in phylogenetic structure at four wells (H32, H41, H43, and H52) was
126 influenced by hydrological seasons (dbRDA, $P < 0.05$). These findings suggest that various responses
127 to changing hydrological seasons elicited by phylogenetically closely related microorganisms may
128 offset seasonal effects.

129 Beyond sinusoidal patterns, all groundwater microbiomes showed long-term variability, as
130 indicated by Mantel tests ($P < 0.05$; Fig. 1a). Microbial community similarity declined over time across
131 all wells, with temporal turnover rates (*i.e.*, the rate of community similarity change) ranging from 1.8%
132 to 5.6% per year (Fig. 1a). While the phylogenetic structures of the microbiome compositions also
133 exhibited long-term variability (Mantel tests, $P < 0.05$), their temporal turnover rates were lower (0.8-
134 2.6% per year, Supplementary Fig. 3). These results suggest that roughly half of the turnover in
135 groundwater microbiomes may be directed towards phylogenetically closely related microorganisms.

136 **Contrasting temporal patterns of groundwater microbiomes**

137 Groundwater microbiomes exhibited contrasting short-term variations, with mean Bray-Curtis
138 distances between samples at one-month intervals ranging from 68% at well H14 to 25% at well H52,
139 reflecting the strength of community composition fluctuations over time (Fig. 1b and 2a). In contrast,
140 groundwater hydrochemical parameters, indicators of environmental changes, showed more consistent
141 short-term variations, with mean Euclidean distances ranging from 14% to 26% (Supplementary Fig.
142 4). This disparity in short-term variations between hydrochemical parameters and microbiome

143 compositions suggests varying levels of microbiome resistance to environmental changes, with greater
144 short-term microbiome variation corresponding to lower resistance.

145 Pronounced short-term (1 month) variability was observed in microbiomes at wells in
146 hydrologically connected aquifers (H14, H43, H41, and H51), while pronounced long-term (10 years)
147 variability was detected in wells representing more hydrologically isolated groundwater (H52 and H53;
148 Fig. 2b). Microbiomes with pronounced short-term variability yielded greater mean Shannon index
149 values (5.2-5.9) and smaller fractions of core microorganisms (9-29%), compared to those with
150 pronounced long-term variability, which exhibited lower Shannon indices (4.5-4.7) and larger core
151 bacterial fractions (55-65%; Fig. 2c, d). Microbiomes at well H32 exhibited equally elevated short-term
152 and long-term variability, with intermediate characteristics between hydrologically connected and
153 isolated groundwaters, which may reflect fluctuating hydrological connectivity at this site (Fig. 2).

154 Microbiomes in hydrologically connected groundwater exhibited lower resistance but higher
155 resilience, as indicated by their lower temporal turnover rates (1.8-3.9% year⁻¹) compared to
156 hydrologically isolated groundwater (5.2-5.6% year⁻¹; Fig. 1a). Here, resilience refers to the ability of
157 groundwater microbiomes to maintain community composition over time despite ongoing
158 environmental change (as suggested by temporal variations in groundwater hydrochemical parameters
159 across all wells; supplementary Fig. 1c). A low resilience is indicated by a high microbial community
160 turnover rate. Temporal stability, measured by mean pairwise Bray-Curtis similarities, was significantly
161 lower in microbiomes of more hydrologically connected groundwater (16-31%) than their
162 hydrologically isolated counterparts (44-48%; Fig. 3a).

163 **Impact of seepage-associated microorganisms on groundwater microbiome variations**

164 To appraise the extent of variability in groundwater microbiomes driven by soil-borne microorganisms
165 transported with the seepage into the groundwater, we compared 16S rRNA gene datasets¹⁶ from soil
166 seepage (23-60 cm depth) in local recharge areas with those from groundwater microbiomes. Seepage
167 microbial diversity, dominated by Gammaproteobacteria (32%), Alphaproteobacteria (24%), and

168 Parcubacteria (10%), differed significantly from that of groundwaters (Permutation tests, $P = 0.001$;
169 Supplementary Fig. 5).

170 SourceTracker analyses revealed that hydrologically connected groundwater harbored
171 significantly higher fractions of seepage-associated microorganisms (2.4-8.9% of the total community)
172 than did hydrologically isolated groundwaters (0.2-1.7%; Wilcoxon tests, $P < 0.001$; Fig. 4b). Despite
173 these variations across wells, however, seepage-associated microorganisms consistently accounted for
174 a mere four percent of groundwater microbiome variation (dbRDA; supplementary Fig. 2). These results
175 suggest that low survival rates of seepage-associated microorganisms limit their contribution in shaping
176 groundwater microbiome composition.

177 A significant correlation was observed between the proportion of seepage-associated
178 microorganisms and groundwater level fluctuations at wells H14 and H41 (Fig. 4), while a significant
179 correlation between precipitation and groundwater levels was observed only at the shallowest well H14
180 ($P_{\text{Pearson}} = 0.03$; $r_{\text{Pearson}} = 0.04$; Supplementary Fig. 6). The fractions of seepage-associated
181 microorganisms at wells H41 and H43 were significantly greater during groundwater recharge (rising
182 water levels; 3.9% and 3.1%, respectively) than groundwater recession (falling water levels; 1.4% and
183 2%, respectively; Wilcoxon test, $P < 0.01$). Seepage-associated microbial diversity in groundwater was
184 dominated by Alphaproteobacteria (mainly *Caulobacter*, *Sphingobium*), Gammaproteobacteria (mainly
185 *Polaromonas*, *Rhodoferrax*), and Parcubacteria (mainly *Candidatus* Adlerbacteria, *Candidatus*
186 *Nomurabacteria*).

187 **Contrasting assembly processes in groundwater microbiomes**

188 Microbiome composition was driven primarily by stochastic processes (stochasticity: 67-82%) in
189 hydrologically connected groundwaters and deterministic processes in more hydrologically isolated
190 groundwaters (stochasticity: 38-42%; Fig. 2e). Temporal dispersal limitation (dispersal limitation over
191 time at one site) ranged from 18% to 32% in hydrologically connected groundwaters and 14% to 15%
192 in more hydrologically isolated groundwaters (Fig. 2e), indicative of more dynamic population
193 structures in the former. This was supported by higher Shannon indices, smaller fractions of core

194 community members, and increased seepage-associated microbial inputs in hydrologically connected
195 groundwaters (Figs. 2c, d; 4b). Correlations between elevated temporal dispersal limitations and lower
196 temporal stabilities in groundwater microbiomes were confirmed by a linear regression model (Fig. 2f).

197 Results of dbRDA analyses highlighted the importance of environmental selection, particularly
198 homogeneous selection, in driving microbiome assembly in hydrologically isolated groundwaters
199 (Supplementary Tab. 2). Hydrochemical parameters accounted for 45-46% and 12-30% of microbiome
200 variation in hydrologically isolated groundwaters and connected groundwaters, respectively.
201 Groundwater level fluctuations yielded the greatest impacts, accounting for 14-21% and 25% of
202 microbiome variation in hydrologically isolated and connected groundwaters, respectively (dbRDA,
203 Supplementary Fig. 2). Regression analyses yielded significant long-term groundwater table declines
204 across all wells, despite periodic recharge, with rates ranging from 7 cm year⁻¹ at well H14 to 126 cm
205 year⁻¹ at well H41 (Supplementary Fig. 6). Other hydrochemical parameters (*e.g.*, temperature) also
206 varied significantly with groundwater levels, leading to significant changes in groundwater
207 environments over time (Mantel test, $P < 0.05$; Supplementary Fig. 1, 7).

208 Collectively, hydrochemical parameters, hydrological seasons and seepage-associated
209 microbial input accounted for 15-33% and 47-50% of microbiome variation in hydrologically connected
210 and isolated groundwaters, respectively (Supplementary Tab. 2).

211 **Functional potentials of groundwater microbiomes**

212 PICRUSt2 identified 390 distinct MetaCyc metabolic pathways from 16S rRNA gene-sequence datasets.
213 Much like the case for groundwater microbiome diversity, the potential metabolic pathway
214 compositions exhibited both sinusoidal patterns and long-term variability (except well H41 which
215 exhibited only sinusoidal patterns; Fig. 5a). Compared to microbiome diversity metrics, the potential
216 metabolic pathway compositions exhibited higher temporal stability (mean Bray-Curtis similarity = 90-
217 95%) and resilience (temporal turnover rates = 0.08-0.9% year⁻¹; Fig. 2b and 5a).

218 **Temporal patterns of groundwater metabolome compositions**

219 LC-HRMS-derived metabolomics revealed both long-term variability and, albeit weak, sinusoidal
220 patterns in the organic landscape (Fig. 5b). Significant correlations were observed between variations
221 in microbiome and metabolic compositions across all wells (Mantel test; Fig. 5b). Variations in
222 metabolomes, indicated by Shannon indices and primary axes of Principal Coordinates Analyses
223 (PCoA), accounted for 11-16% and 39% of microbiome variation in hydrologically connected and
224 isolated groundwaters, respectively (dbRDA; Supplementary Fig. 2).

225 Metabolome variation, together with hydrochemical parameters, hydrological seasons, and
226 seepage-associated microbial input accounted for 31-59% of microbiome variation, roughly 10% more
227 than when metabolome variation was excluded (dbRDA; Supplementary Tab. 2). In addition, the
228 temporal stability of metabolomes (34-45%) was more consistent across wells than that of groundwater
229 microbiomes (per mean Bray-Curtis similarity; Fig. 3).

230 **Temporal patterns of groundwater DOM concentration and compositions**

231 Dissolved organic carbon (DOC) concentrations ranged from below detection limit ($< 0.5 \text{ mg L}^{-1}$) to
232 2.8 mg L^{-1} , with higher concentrations detected in shallow wells H14, H43, and H32 (Supplementary
233 Fig. 8). Variations in DOC accounted for one to six percent of the microbiome variation in wells H41
234 and H43, which were most affected by hydrological seasons (dbRDA; Supplementary Fig. 2). While
235 bulk DOM compositions elucidated via DI-HRMS lacked significant correlation with microbiome
236 variations (Mantel test; Fig. 5c), they exhibited greater resilience (lower turnover rates) and temporal
237 stability (greater mean Bray-Curtis similarity: 79-84%; Fig. 3, 5c) than microbiomes.

238 To link DOM composition to microbial functions, we compared the compound classes inferred
239 from DI-HRMS data to the potential degradation pathways inferred from 16S rRNA gene-sequences
240 (Supplementary Tab. 3). The relative abundances of most degradation pathways were significantly
241 greater in microbiomes in hydrologically connected than their hydrologically isolated counterparts
242 (Wilcoxon test, $P < 0.001$), except for C1 compound utilization and polymeric compound degradation.

243 The relative abundances of two of the six DOM compound classes identified differed
244 significantly between hydrologically connected and isolated groundwaters (Fig. 6; Supplementary Fig.

245 9). Condensed aromatic structures, likely derived from Muschelkalk sediments and underlying
246 Permian/Jurassic formations, terrestrial plant material, or microbial necromass^{38,39,40}, were more
247 abundant in hydrologically connected groundwaters (0.15-0.4% vs. 0.04-0.08%; Wilcoxon test, $P < 0.01$;
248 Fig. 5). This suggests that higher levels of condensed aromatics corresponded to an increased
249 degradation potential for these compounds (Fig. 5). Conversely, relative abundances of peptide-like
250 DOM, potentially sourced from extracellular enzymes and microbial necromass³⁹, were significantly
251 lower in hydrologically connected groundwaters (0.12-0.2% vs. 0.22-0.36%; Fig. 6). Yet, microbiomes
252 in hydrologically connected groundwaters exhibited greater degradation potential for amino acids, even
253 though their potential ability to synthesize them was not consistently lower (Fig. 6; Supplementary Fig.
254 10). This suggests that these communities recycle environmental proteins to a greater extent than their
255 counterparts in more hydrologically isolated groundwater.

256 Between May 2020 and October 2021, condensed aromatic structures and unsaturated
257 hydrocarbons significantly increased at shallow wells H14 and H32 (and others). Mean relative
258 abundances of condensed aromatics rose from 0.27% to 0.44% and from 0.19% to 0.39% (Wilcoxon
259 test, $P < 0.01$), while unsaturated hydrocarbons increased from 2.4% to 3.7% and from 2% to 3.8% at
260 wells H14 and H32, respectively (Wilcoxon test, $P < 0.01$). Following these increases, the degradation
261 potential for aromatics and carboxylic acids at H14 and H32 rose significantly, beginning in March and
262 April 2021, respectively. Aromatic degradation potential increased from 0.4% to 1.4% and from 0.06%
263 to 1% (Wilcoxon test, $P < 0.001$; Supplementary Fig. 11), while carboxylic acid degradation potential
264 increased from 0.3% to 0.6% and from 0.07% to 0.3% at wells H14 and H32, respectively (Wilcoxon
265 test, $P < 0.001$; Supplementary Fig. 11). The relative abundance of Gammaproteobacteria rose
266 significantly at H14 (30.4% to 69.6%) and H32 (15.7% to 34.7%; Wilcoxon test, $P < 0.001$; Fig. 1b).
267 These results suggest delayed changes in microbial composition and function in response to changes in
268 DOM in shallow groundwater wells.

269

270

271 Discussion

272 Our decade-long study of groundwater microbiomes unveiled hydrological seasonality and long-term
273 variability, challenging the traditional view of groundwater microbiomes as static entities^{7,36,41}. Trends
274 of continuous change in groundwater microbiomes were previously reported by Yan et al.¹³ from the
275 same study sites, but sinusoidal patterns emerged only with this 10-year analysis. These patterns follow
276 local seasonal hydrological changes, with hydrochemical parameters exhibiting annual sinusoidal
277 patterns corresponding to groundwater level variations driven by meteoric recharge⁵ (Supplementary
278 Fig. 1, 6, 7). Similar periodic groundwater level fluctuations have been observed in other shallow
279 aquifers (< 100 m depth), especially in karst systems^{10,11,12}. Sinusoidal patterns in microbiomes were
280 never reported, however, even though shifts resulting from environmental changes were reported^{12,22,42},
281 perhaps owing to shorter study periods and/or lower observation frequencies. Annual sinusoidal patterns,
282 driven by seasonal factors (*e.g.*, day length, temperature, nutrient availability), are common in long-
283 term studies of marine and lake environments, particularly marine surface layers^{25,26,43,44}.

284 Sinusoidal patterns in subsurface microbiome diversity may arise from specific microorganisms
285 being periodically favored by fluctuating environmental conditions, such as redox potential and/or
286 groundwater levels^{11,13}. Microorganisms benefitting from high groundwater levels include those
287 remobilized from rock surfaces in the vadose zone, as local carbonate rocks and planktonic groundwater
288 communities share up to 40% species diversity⁴⁵. This fraction might be even higher in porous aquifers⁴⁶,
289 rendering them important seeding banks for groundwater microbiomes²¹. In addition to enhancing
290 microbiome stability by increasing homogeneous selection (*e.g.*, well H41; Fig. 2e), the sinusoidal
291 patterns characterized in this study drastically improve predictions of microbiome change^{47,48}.

292 The Hainich CZE provides an ideal setting to study microbiome stability in carbonate-rock
293 aquifers of varying hydrological connectivity. Our extended in-depth characterization of seven selected
294 wells considered several distinct microbiomes along the groundwater monitoring transect⁴⁹, revealing
295 contrasting temporal patterns and stability in groundwater microbiomes over a mere 6 km. The
296 microbiomes of more isolated groundwaters with high temporal stability (*e.g.*, wells H52 and H53) are

297 influenced largely by deterministic processes, while those in hydrologically connected groundwaters
298 are shaped by stochastic processes (and as such exhibit lower temporal stability). In well H32, an
299 intermediate aquifer system, groundwater microbiomes exhibited both elevated short- and long-term
300 variability, reflecting sporadic hydrological connectivity likely resulting from complex flows in
301 fractured sedimentary bedrock aquifers⁵.

302 Our results corroborate the findings of recent groundwater studies in other geological
303 settings^{11,50,51}, suggesting that hydrological connectivity reduces the temporal stability of groundwater
304 microbiomes by increasing the role of temporal dispersal limitation via microbial immigration. Thus,
305 hydrological connectivity lowers the resistance of the groundwater microbiome while increasing its
306 resilience. Furthermore, microbial immigration during recurrent groundwater recharge events
307 represents intermittent disturbances that promote resilience and create opportunities for coexistence by
308 preventing competitive exclusion, thereby promoting diversity⁵².

309 In near-surface aquifers, seepage-associated microbial input is a key contributor of microbial
310 immigration, increasing the importance of temporal dispersal limitation in microbiome assembly^{20,22}.
311 Our findings suggest that 0.2-8.9% of the total groundwater microbiome is seepage-borne, likely
312 derived during periodic groundwater recharge or single hydrological extreme events (particularly
313 relevant in karst systems). Recurrent immigration events over thousands of years may have enhanced
314 the persistence of these invaders in the groundwater microbiome, as up to 12.5% of groundwater core
315 ASVs have been identified as being seepage-associated. As these invasive taxa survive and partake in
316 community coalescence, they compete with resident taxa for resources, and closely related species
317 compete for similar resources⁵³. Such competition between species oftentimes results in divergence in
318 niches, including increased host specificity of episymbiotic microorganisms (*e.g.*, Patescibacteria), to
319 reduce competition costs⁵⁴.

320 The effects of hydrological connectivity on microbiome stability extends beyond groundwater,
321 with similar patterns reported in oceans and other aquatic ecosystems^{55,56,57,58,59,60}. Deep ocean
322 microbiomes (890 m depth) are more temporally stable, with mean Bray-Curtis similarities around 64%,
323 compared to 40% in the more hydrologically connected surface layers (5-20 m depth) where currents

324 and nutrient mixing enhance microbial exchange^{55,60}. Somewhat surprisingly, surface/shallow ocean
325 microbiomes are more stable than those of hydrologically connected groundwaters, and show reduced
326 long-term community turnover^{43,55,60}. Compared to groundwaters, marine microbiomes exhibit higher
327 resilience (*i.e.*, lower turnover rates) throughout all depths/layers^{43,55,60}. This suggests that groundwater
328 systems are more vulnerable and less resilient than marine systems⁶¹.

329 The functional potentials of all seven groundwater microbiomes exhibited remarkable temporal
330 stability and resilience despite notable variations in community composition. This aligns with previous
331 groundwater studies that report high functional redundancy in these ecosystems¹¹. Functional
332 redundancy, *i.e.*, multiple microbial taxa performing overlapping ecological roles, ensures ecosystem
333 stability amid species turnover⁶². The observed resiliencies in functional potential suggests that, despite
334 environmental fluctuations, these microbial communities retain the capacity to adapt while maintaining
335 their functional integrity^{8,21}. Significant correlations were observed between microbiome and
336 metabolome changes, pointing to a form of functional redundancy achieved through metabolic diversity.
337 Such a mechanism enables different microbial taxa to produce varied metabolites that fulfill similar
338 ecological functions⁶³. Metabolome compositions exhibited greater temporal consistency across wells
339 than microbiomes. This stands to reason as not all detected metabolites are microorganism-related,
340 which in and of itself highlights the influence of additional controlling factors^{28,30}. The lower temporal
341 stability of groundwater metabolomes compared to corresponding microbiomes' elevated functional
342 potentials corroborates earlier findings which reported substantial metabolome variability at the same
343 site²⁹.

344 Distinct temporal patterns were observed in groundwater metabolome and DOM compositions,
345 largely attributable to differences in analytical methods that preferentially target different compound
346 classes^{14,29}. The high temporal stability of DOM across wells may result from its composition, which
347 consists of lignin-degradation products and polyphenolic leachates from plant matter in topsoil^{36,64,65}
348 (Supplementary Fig. 9). These compounds resist microbial degradation, contributing to their persistence
349 in the environment^{36,64,65}. Despite comprising less than 5% of total DOM, the microorganism-associated
350 DOM fraction likely plays a significant role in subsurface organic matter cycling. Its influence on spatial

351 and temporal variation in microbiome composition and function^{36,66,67} was particularly evident in the
352 shallowest well (*e.g.*, H14, H32). The contrasting dynamics observed between DOM and metabolomes
353 emphasize the value of integrating diverse analytical approaches^{28,32}.

354 This long-term study shows that contrasting hydrogeological conditions render the Hainich CZE
355 an ideal setting to elucidate the various mechanisms driving groundwater microbiome stability and
356 vulnerability in carbonate-rock aquifers. Highly karstified regions, rich with extensive fractures and
357 conduits, are globally crucial for drinking water as they facilitate rapid groundwater flow and abundant
358 water storage (in contrast to less permeable, dense carbonate rocks)⁴. Karst aquifer microbiomes are
359 typically influenced most largely by stochastic processes and are more susceptible to disturbances from
360 surface-derived inputs (*e.g.*, organic contaminants, pathogens)^{22,67,68,69}. As such, improper surface
361 management can rapidly increase their vulnerability and threaten groundwater health. Extensive source
362 control to limit surface contaminant ingress to aquifers is crucial for maintaining the health of these
363 ecosystems²³. Moreover, hydroclimatic extremes, such as heavy precipitation and drought, evidently
364 exacerbate groundwater vulnerability. These extremes facilitate the high ingress of surface-derived
365 organic molecules (*e.g.*, xenobiotic substances) into groundwater by evading microbial processing¹⁴.
366 The high ingress of surface-associated bacteria during recharge and heavy precipitation events raises
367 concerns about pathogen contamination, threatening groundwater quality. These findings underscore
368 the vulnerability of highly hydrologically connected aquifers, such as karst aquifers, under
369 hydroclimatic extremes.

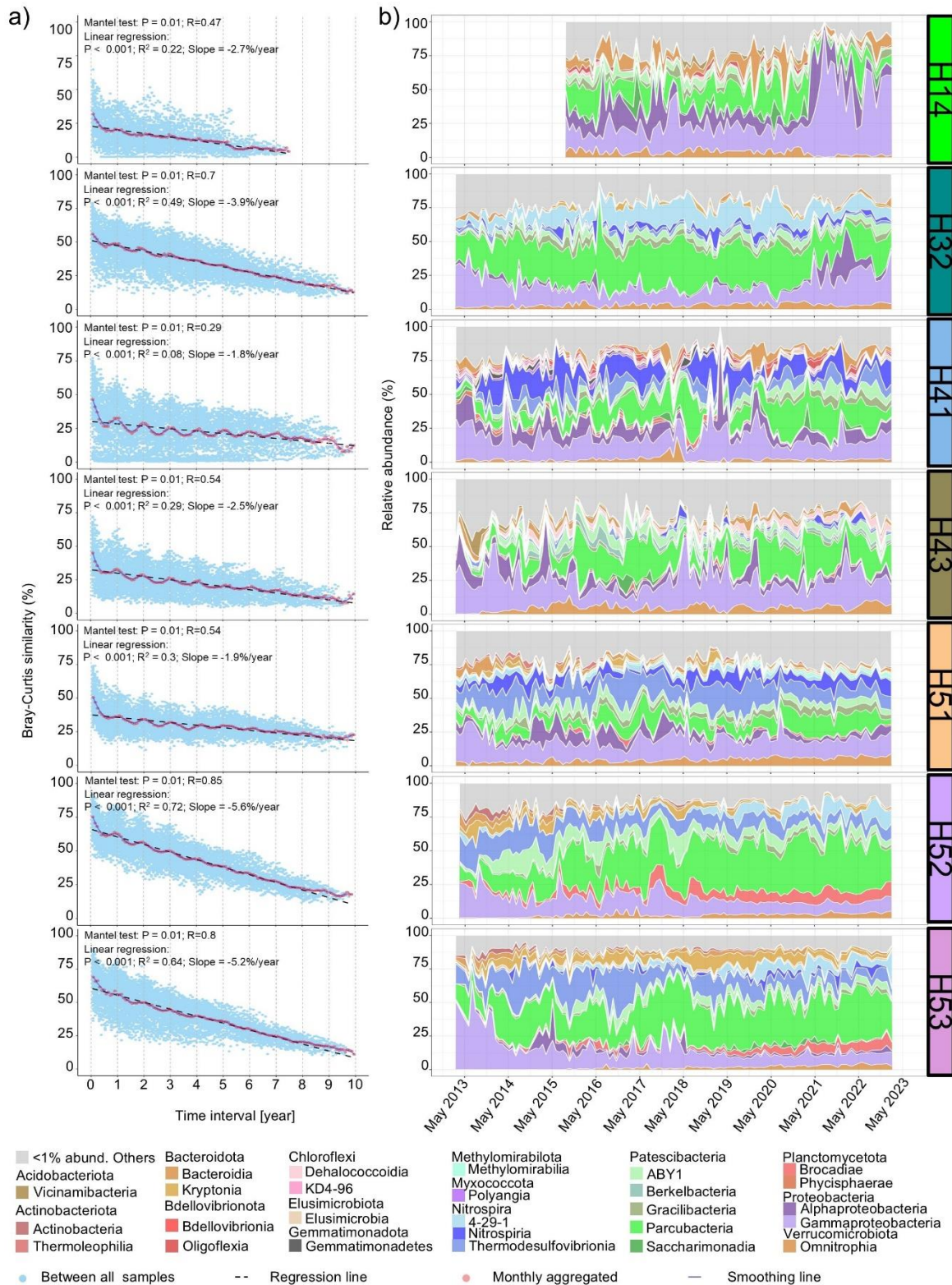
370 In contrast, microbiomes in more hydrologically isolated groundwaters (*e.g.*, dense limestone
371 formations) tend to be governed by deterministic processes and often resist disturbances^{70,71}. However,
372 this study demonstrates that continuous environmental changes result in decreasing temporal stability
373 in microbiomes in hydrologically isolated groundwater over time, even in the absence of contamination.
374 Such changes may be driven by prolonged droughts affecting groundwater recharge. While these
375 systems may exhibit low short-term variability and greater resistance to contamination, their recovery
376 following severe contamination is likely prolonged, as a new ecological balance emerges.

377

378 **Conclusion**

379 Overall, the results of this study show that both hydrologically connected and isolated groundwater
380 ecosystems in carbonate aquifers are increasingly at risk under scenarios with more frequent
381 hydrological extremes that alter the connections between surface water and groundwater. As
382 hydroclimatic extremes intensify, the persistence of core microbial functions becomes increasingly
383 important to maintaining ecosystem stability. The findings discussed here underscore the importance of
384 innovative integrated management strategies (*e.g.*, robust surface control measures, long-term
385 monitoring of aquifer health) to safeguard Europe's groundwater resources from the stresses of
386 contamination and climate change.

387 **Figures**

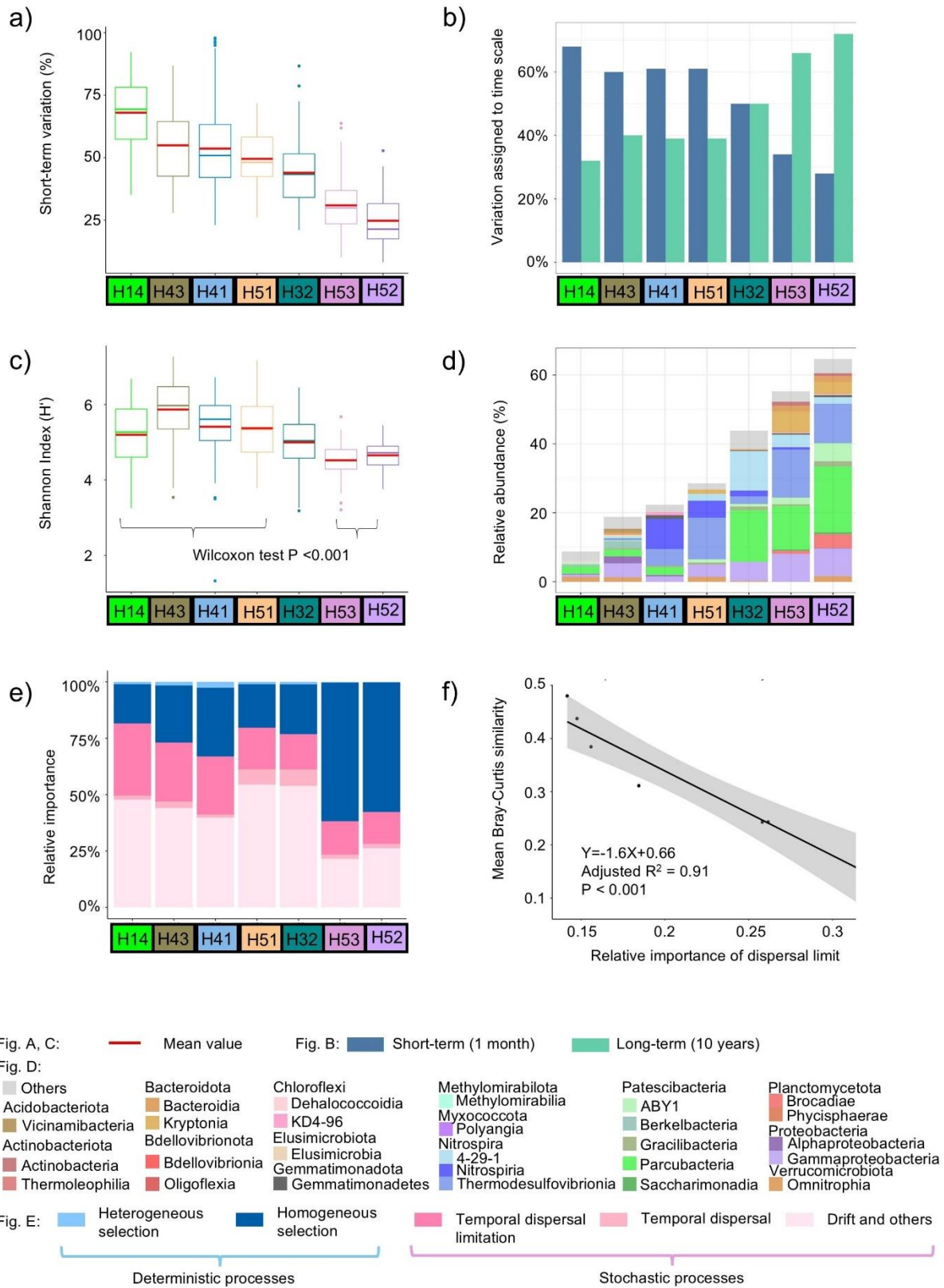


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389 Fig. 1 Sinusoidal and long-term variability in groundwater microbiomes. a) Temporal patterns in
 390 groundwater microbial community similarity: blue points represent the raw Bray-Curtis similarity
 391 between all samples against their sampling time intervals while red points represent the monthly average
 392 Bray-Curtis similarity between all samples. Regression lines and smoothing lines are applied to better
 393 envision temporal patterns. The turnover rate (%/year) is calculated as the slope of the regression lines.
 394 b) Temporal patterns in groundwater microbial community diversity at different wells: colors represent

395 bacterial classes with relative abundances greater than 1%. The first six years of microbiome data at
 396 wells of H41, H43, and H52 were published in Yan *et al*¹³.

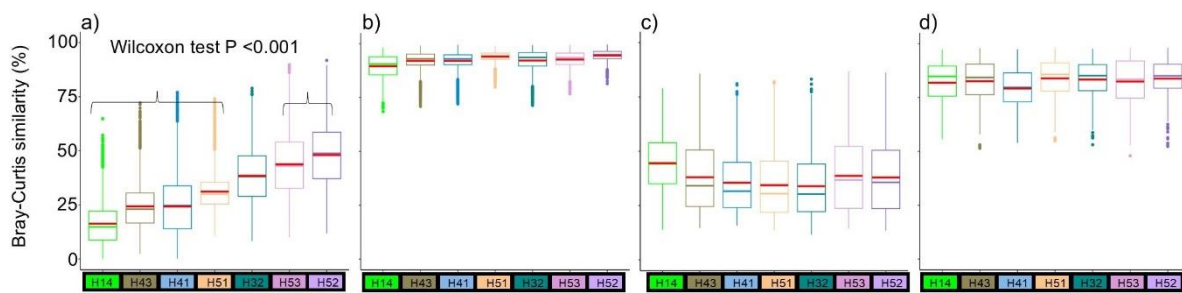
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398

399 **Fig. 2 Distinct temporal patterns of groundwater microbiomes.** a) Short-term variations in
400 groundwater microbiome diversity based on Bray-Curtis distances between samples collected one
401 month apart (short red horizontal lines denote corresponding mean values). b) Variation in groundwater
402 microbiomes assigned to short-term (1 month in blue bar) and long-term (10 years in green bar) time
403 scales. c) Shannon indices of groundwater microbiomes (short red horizontal lines denote
404 corresponding mean values). d) Fraction of core microorganisms (ASVs present $\geq 80\%$ of collected
405 well samples; colors represent different bacterial classes). e) Relative importance of groundwater
406 microbiome assembly processes: green hues represent deterministic processes (heterogeneous selection
407 and homogeneous selection), while red hues represent stochastic processes (dispersal limitation,
408 horizontal dispersal, and drift and others). Since these analyses were conducted on a single site over
409 time, the dispersal here was related to time rather than space. f) Significant linear model between the
410 relative abundance of (temporal) dispersal limitation and temporal stability of groundwater microbiome
411 diversity (using mean Bray-Curtis similarity).

412

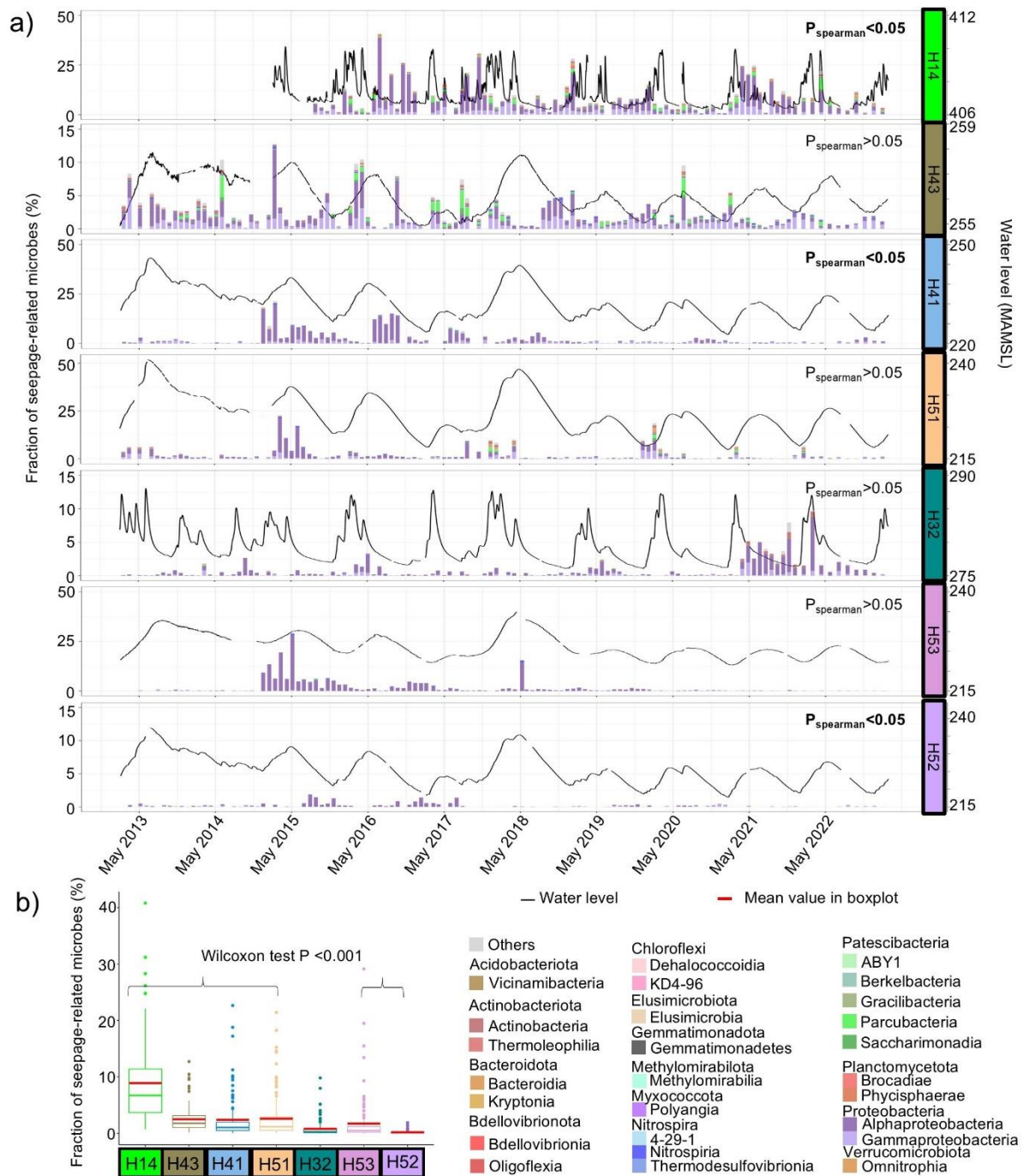


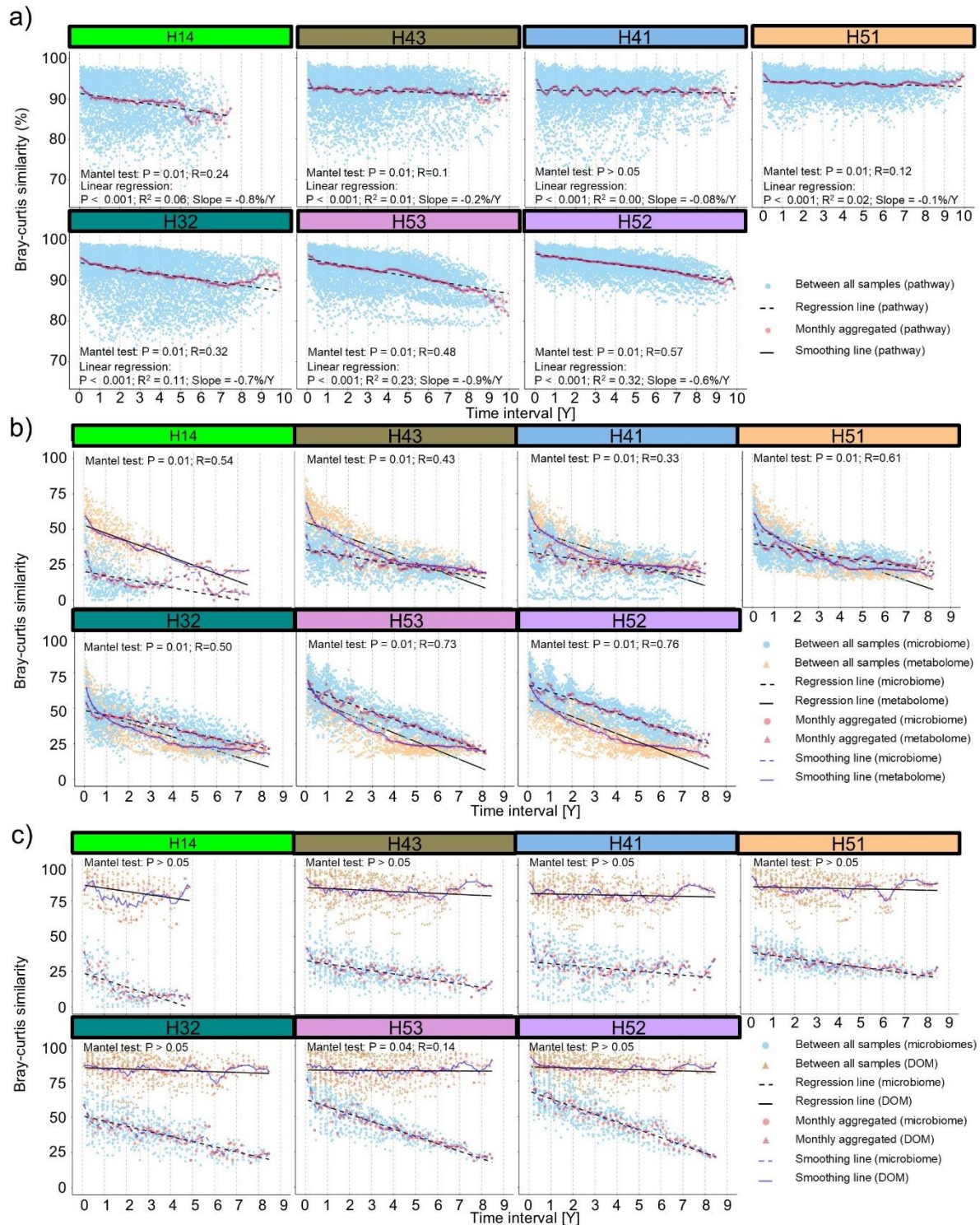
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414 **Fig. 3 Temporal stability of a) groundwater microbiome compositions, b) metabolic potential of**
415 **groundwater microbiomes, c) metabolome compositions, and d) bulk DOM compositions across**
416 **wells per Bray-Curtis similarity.** The metabolic potential of groundwater microbiomes in figure c)
417 was inferred from 16S rRNA gene-sequences. Red horizontal lines denote mean values.

418

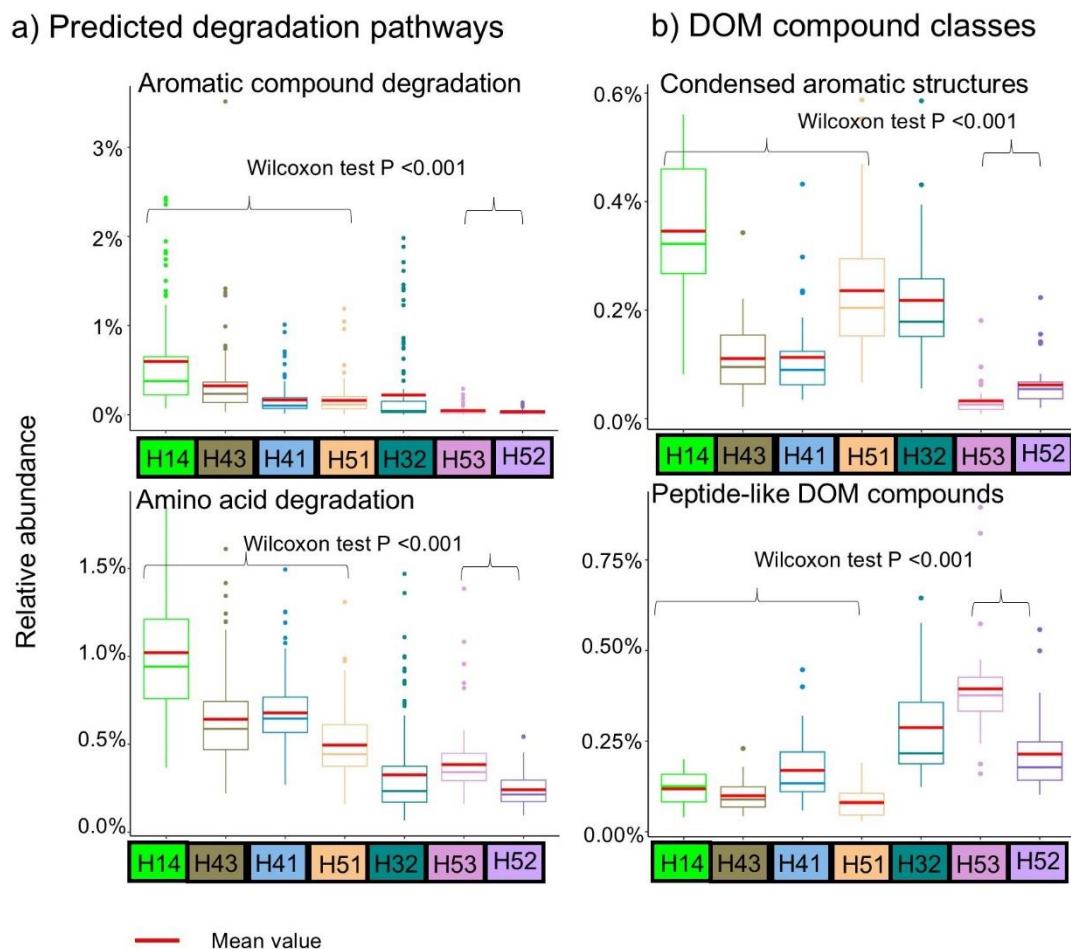
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426

427 **Fig. 5 Temporal patterns of metabolic potential of microbiomes, metabolomes, and DOM**
 428 **compositions.** a) Temporal patterns of metabolic pathways predicted by PICRUSt2 using 16S
 429 rRNA gene-sequences; b) Temporal patterns of groundwater metabolomes alongside a subset
 430 of concurrently sampled groundwater microbiomes; c) Temporal patterns of groundwater DOM
 431 compositions alongside a subset of concurrently sampled groundwater microbiomes. Mantel
 432 tests were conducted to assess correlations a) between changes in predicted pathway and time;
 433 b) between changes in microbiomes and metabolomes; c) between changes in microbiomes and
 434 DOM.



435

436 Fig. 6 **Boxplot of selected a) predicted degradation pathways and b) DOM compound classes.** The
437 predicted degradation pathways were inferred from 16S rRNA gene-sequences, while DOM compound
438 classes were inferred from DI-HRMS derived bulk DOM. Red horizontal lines denote mean values.

439

440 **Methods**

441 **Study sites and sampling**

442 The study site is located in the Hainich CZE in central Germany. Detailed site information and sampling
443 procedures can be found in Kohlhepp *et al.*³⁴, Küsel *et al.*³⁵, and Lehmann and Totsche⁵. Briefly, the
444 CZE features alternating thin-bedded limestone-mudstone in hillslope terrains, which is a common and
445 widely distributed geological setting. The bedrock of the low-mountain hillslope in the eastern Hainich
446 CZE consists primarily of Upper Muschelkalk of marine origin (Germanic Triassic), parts of which
447 harbor abundant groundwater resources. Along this hillslope, our seven monitoring wells are embedded

448 in a transect spanning 5.4 km (Supplementary Fig. 1A), encompassing various relief positions, aquifers,
449 and depths³⁵. Aquifer compartments at Hainich CZE differ significantly in oxygen availability. Wells
450 H14, H32, H41, and H51 are oxic, H43 is suboxic (< 1 mg/L dissolved oxygen), and H52 and H53 are
451 anoxic (< 0.1 mg/L dissolved oxygen). Redox potentials range around 400 mV for wells H14, H32,
452 H41, and H51, and around 200 mV for wells H43, H52, and H53. The wells are situated in areas of
453 extensive land management, including forest, pasture, and cropland. The shallowest well (H14), located
454 on the upper hillslope, is the only well located in the preferential groundwater recharge area.

455 Groundwater samples were collected monthly to monitor standard hydrochemical parameters,
456 microbial communities, and untargeted metabolomics. DOM was measured every three months. From
457 February 2013 to February 2023, we collected 815 groundwater samples from seven groundwater wells
458 for microbial analysis and hydrochemical parameters, 226 groundwater samples for DOM, and 387
459 samples for metabolomic analyses (Supplementary Fig. 12). Once the physico-chemical parameters of
460 pumped groundwater stabilized, groundwater was collected from each well using a submersible pump
461 (MPI, Grundfos) and placed into sterilized bottles. These samples included 5–10 L for microbiological
462 analysis, duplicate 10 L samples for DOM analysis, triplicate 5 L samples for metabolomics, and 100
463 mL for DOC concentration measurements. Groundwaters from which to isolate genomic DNA was
464 processed through 0.2 µm filters (PES or polycarbonate from Supor, Pall Corporation and Merck–
465 Millipore, respectively) via vacuum pumping, and filters were stored at –80°C prior to DNA extraction.
466 Acidified filtered groundwaters (0.7 µm filters, pH = 2 with HCl) were stored at 4°C in the dark until
467 further DOM processing, while raw groundwaters were stored in the dark and chilled until further
468 metabolomics analyses. Other groundwaters were filtered through 0.7 µm filters and stored at 4°C in
469 the dark until further DOC quantification.

470 Groundwater hydrochemical parameters including groundwater levels, temperature, pH,
471 specific electrical conductivity (EC₂₅; reference T: 25°C), dissolved oxygen content, redox potential
472 (ORP), acidity (neutralizing capacity), alkalinity (neutralizing capacity), total inorganic carbon (TIC),
473 and ion concentrations were measured as described by Lehmann and Totsche⁵ and Kohlhepp *et al.*³⁴.
474 Element concentration including Ca, K, Mg, Na, and S were measured with ICP-MS (inductively

475 coupled plasma mass spectrometry; 8900 Triple Quadrupole ICP-MS, Agilent, Germany), while major
476 anions Cl^- was measured by IC (ion chromatography; Dionex IC20, Thermo Fisher Scientific, USA).
477 DOC concentration was quantified as non-purgeable organic carbon on a vario TOC cube (Elementar
478 Analysensysteme, Germany) with a detection limit of 0.5 mg L^{-1} .

479 The sampling details for the 174 seepage samples collected in this study can be found in
480 Hermann *et al.*¹⁶ and Lehmann *et al.*⁷² (Supplementary Fig. 1; Supplementary Tab. 4). Seepage sites
481 were sampled regularly (biweekly) and on an event-basis (weekly). Seepage volumes ranging from 100
482 to 500 mL were filtered through $0.2 \mu\text{m}$ filters (PES; Supor, Pall Corporation) using a vacuum pump,
483 and filters were stored at -80°C prior to DNA extraction.

484 **DNA extraction and amplicon sequencing**

485 Groundwater and seepage genomic DNA were extracted using the DNeasy PowerSoil Pro Kit (Qiagen,
486 Hilden, Germany) per manufacturer's instructions, and extractions were stored at -20°C prior to PCR
487 amplification. PCR amplification of bacterial 16S rRNA gene (V3–V4 region) was performed on an
488 Illumina Miseq platform using v3 chemistry with primers Bakt_0341F and Bakt_0785R⁷³. Most
489 samples (see list in Supplementary Tab. 5) were sequenced in-house following the two-step PCR
490 library preparation procedures described by Krüger *et al.*¹⁷. Amplicon libraries for some samples were
491 generated using the NEBNext Ultra DNA Library Prep Kit for Illumina (New England Biolabs, MA),
492 following methods detailed by Kumar *et al.*⁷⁴ All remaining samples were processed at LGC Genomics
493 (Berlin, Germany), as previously described⁷⁵.

494 **Molecular composition of DOM**

495 A detailed description of the methods used to elucidate the molecular compositions of DOM is given in
496 Schroeter *et al.*¹⁴. Briefly, DOM was extracted from acidified filtered groundwater via solid phase
497 extraction with PPL (styrene–divinylbenzene polymer) Bond Elut cartridges (Agilent Technologies)
498 following the protocols of Dittmar *et al.*⁷⁶. PPL extracts were kept at -80°C prior to DI-HRMS analyses.
499 DI-HRMS analyses were conducted on an Orbitrap Elite mass spectrometer (Thermo Fisher Scientific,
500 USA) with a mass resolution of $555,000 \pm 9,000$ at $m/z = 251$. The electrospray ionization (ESI) was

501 run in negative mode with an ESI needle voltage of 2.65 kV. For each sample, 100 scans of m/z 100-
502 1000 were acquired and averaged. Quality controls and molecular formula assignments were processed
503 using DOMAssignR (<https://github.com/simonschroeter/DOMAssignR>). Spectra were further
504 normalized to sum all peak intensities. We focused on DOM compound classes most likely to serve as
505 substrates for microorganisms, including carbohydrates, condensed aromatic structures, lignins, lipids,
506 peptide-like compounds, and unsaturated hydrocarbons. Identification of DOM clusters was based on
507 their elemental N content and hydrogen/carbon and/or oxygen/carbon elemental ratios (Supplementary
508 Tab. 6)⁷⁷.

509 **Untargeted metabolomics**

510 Sample extraction and analyses are described in detail in Zerfaß *et al.*²⁹. Briefly, 5 L of filtered (GF/C,
511 1.5 μm , VWR) groundwater was subjected to solid phase extraction (SPE) in Strata-X 33 μm polymeric
512 reversed phase cartridges (Phenomenex). Eluates (1:1 methanol:acetonitril) were dried (vacuum,
513 nitrogen stream), and the organic residue was re-dissolved in 100 μl of 1:1 THF:methanol. Extracts (1
514 μl) were then analyzed by LC-HRMS (liquid chromatography-high-resolution mass spectrometry) on
515 a Dionex UltiMate 3000 chromatography system coupled to a Q-Exactive Plus orbitrap mass
516 spectrometer (Thermo Fisher Scientific), m/z -range 100-1,500, alternating acquisition in positive and
517 negative mode. For this study, only positive-mode data was extracted.

518 To assure system suitability in the long-term sampling experiment, the LC-HRMS system was
519 maintained with a weekly MS source cleaning, mass calibration, and consistency-check by injection of
520 a standard (containing Fluorophenylalanine, P-Fluorobenzoic acid, Decanoic acid D-19) for which
521 retention times and peak apex intensities were recorded. All samples were taken in environmental
522 replicates and injected in triplicate analytical replicates, and each set of analytical replicates was
523 processed in randomized sequence. Data processing for peak picking and feature assignment was
524 carried out in XCMS as described in the stated reference. For Bray-Curtis similarity tests (details in
525 succeeding section), replicate means were calculated and peak areas were normalized by the sum of all
526 feature peak areas.

527

528 **Bioinformatics and statistical analyses**

529 Most bioinformatics and/or statistical analyses were conducted in R version 4.2.2⁷⁸ at a significance
530 level of $\alpha = 0.05$. After correcting the orientation of mixed-orientation reads and removing primers with
531 Cutadapt⁷⁹ (v 4.1), R package “dada2”⁸⁰ (v1.26) was employed for quality filtering, denoising, inferring
532 amplicon sequence variants (ASVs), and removing chimeras. Reads were truncated to 265 bp (forward
533 reads) or 235 bp (reverse reads), excluding those with more than two expected errors and those truncated
534 when the quality score was equal to or less than two. ASVs were generated by applying the DADA2
535 core algorithm and combining forward and reverse reads, and those that could not be aligned to the
536 SILVA reference database⁸¹ (v138.1) using Mothur (v1.46.1) were removed. After removing chimeric
537 sequences, taxonomy was assigned to the remaining ASVs based on the SILVA taxonomy reference
538 database v138.1. Further downstream sequence analysis was conducted using the R package “phyloseq”
539 ⁸² (v1.42), and a phylogenetic tree was constructed using FastTree⁸³ (v2.1) after aligning genes with
540 Muscle⁸⁴ (v5) and trimming the alignment with trimAl⁸⁵ (v1.4).

541 Recharge and recession phases (Supplementary Fig. 6 and Tab. 5) were defined by observed
542 changes in groundwater levels. The recession phase was characterized by a sustained decline in
543 groundwater levels for more than five consecutive days, ending in a minimum level or (for well H14) a
544 significant rise to or above the maximum level for that phase. The recharge phase was defined in an
545 opposite manner. We considered one hydrological year (between 1 May and 30 Apr) comprising four
546 seasons: hydrological early summer (May to Jul), late summer (Aug to Oct), early winter (Nov to Jan),
547 and late winter (Feb to Apr). To visualize the heterogeneity of groundwater environments, principal
548 component analysis (PCA) was conducted with 15 hydrochemical parameters, including water level,
549 water temperature, specific electrical conductivity, dissolved oxygen concentration, redox potential,
550 acidity, alkalinity, total inorganic carbon, element and ion concentration (Cl⁻, S, Ca, K, Na, and Mg).
551 Pearson and Spearman correlations from Hmisc (Rpackage v5.1-1) were used to determine significant
552 correlations.

553 We used mean pairwise Bray-Curtis similarity to evaluate the temporal stability of groundwater
554 microbiome composition and function, while short-term variabilities was assessed using Bray-Curtis
555 distances for sample pairs from the same well (sampling intervals of 15 to 44 days). To explore temporal
556 patterns, we first plotted raw pairwise Bray-Curtis similarity against sampling time intervals. To
557 illustrate community similarity decay rates (resilience), we modeled the regression line of the raw
558 pairwise Bray-Curtis similarity. To identify periodic patterns, we divided the pairwise Bray-Curtis
559 similarity into monthly intervals (*e.g.*, a time interval of 1 month: 15-44 days; a time interval of 2
560 months: 45-74 days), and applied a smoothing line based on a moving average filter. The same
561 technique was applied to groundwater DOM and metabolomes using Bray-Curtis similarity,
562 groundwater microbiome compositions at the phylogenetic level using UniFrac similarity, and
563 environmental parameters using Euclidean similarity. The Bray-Curtis similarity (distances) and
564 UniFrac similarity (distance) mentioned in this manuscript were based on relative abundance, while the
565 Euclidean similarity (distance) was calculated after normalizing the aforementioned 15 hydrochemical
566 parameters to exclude the effect(s) of absolute abundance (values) differences. Variation assigned to
567 short-term or long-term variability is represented as the proportion of the respective variability in one-
568 decade variability. We first evaluated the 10-year variability of microbiome compositions using the
569 regression model of raw pairwise Bray-Curtis similarity, at which time long-term variability is then
570 calculated by subtracting the aforementioned short-term variability from the 10-year variability.

571 To elucidate the impact of surface-subsurface connectivity in the temporal stability of
572 groundwater microbial communities, we tracked changes in contribution of seepage-associated
573 microorganisms to the groundwater microbiome via SourceTracker⁸⁶ (Rpackage v1.0.1 under R version
574 4.0.0). During these analyses, we applied all seepage samples as the source and the groundwater samples
575 as the sink. Analyses were performed on rarefied ASV (rarefaction depth of 1,000 reads) abundances
576 using default settings ($\alpha = 0.001$, β_1 and β_2 : 0.01).

577 In this study, we define core species as bacterial ASVs present in > 80% of the groundwater
578 samples collected in every well. Distance-based redundancy analysis (dbRDA) based on Bray-Curtis
579 dissimilarity was carried out to assess and validate whether selected environmental parameters (*e.g.*,

580 hydrological seasons, incidence of seepage-associated microorganisms) significantly impact
581 groundwater microbiome composition (Rpackage vegan; adjusted R^2 was used). The significance of
582 dbRDA tests was reported by permutation tests of “anova.cca”. The metabolic functions of groundwater
583 microbial communities were predicted by PICRUST2 software based on taxonomy annotations from
584 16S rRNA gene sequences³³. We employed inferred community assembly mechanisms using a
585 phylogenetic bin-based null model (iCAMP, Rpackage v1.6.5) to evaluate the contribution of ecological
586 processes on groundwater microbiome assembly⁸⁷. All analyses were performed using recommended
587 default settings with 48 bins and confidence for null model significant tests.

588 **Data and code availability**

589 Raw amplicon sequencing data reads for all studied samples have been deposited in the European
590 Nucleotide Archive (details in the Supplementary Material). Raw DOM data from DI-HRMS were
591 deposited under <https://doi.org/10.17617/3.2TZM6C>, while raw metabolome data from LC-HRMS
592 were deposited via the Metabolights repository⁸⁸ under MTBLS3450, MTBLS8433, and MTBLS11375.
593 Groundwater hydrochemical parameters are provided as Supplementary Material.

594 Raw amplicon sequencing data and codes will be released upon publication of the manuscript.

595

596 **References**

- 597 1. Gleeson, T., Wada, Y., Bierkens, M. F. P. & Van Beek, L. P. H. Water balance of global aquifers
598 revealed by groundwater footprint. *Nature* **488**, 197–200 (2012).
- 599 2. Rodell, M. *et al.* Emerging trends in global freshwater availability. *Nature* **557**, 651–659 (2018).
- 600 3. Cuthbert, M. O. *et al.* Observed controls on resilience of groundwater to climate variability in sub-
601 Saharan Africa. *Nature* **572**, 230–234 (2019).
- 602 4. Ford, D. & Williams, P. Karst Hydrogeology and Geomorphology. *John Wiley Sons* (2007).
- 603 5. Lehmann, R. & Totsche, K. U. Multi-directional flow dynamics shape groundwater quality in
604 sloping bedrock strata. *J. Hydrol.* **580**, 124291 (2020).
- 605 6. Retter, A., Karwautz, C. & Griebler, C. Groundwater Microbial Communities in Times of Climate
606 Change. *Curr. Issues Mol. Biol.* 509–538 (2021).
- 607 7. Griebler, C. & Lueders, T. Microbial biodiversity in groundwater ecosystems. *Freshw. Biol.* **54**,
608 649–677 (2009).
- 609 8. Griffiths, B. S. & Philippot, L. Insights into the resistance and resilience of the soil microbial
610 community. *FEMS Microbiol. Rev.* **37**, 112–129 (2013).
- 611 9. Philippot, L., Griffiths, B. S. & Langenheder, S. Microbial Community Resilience across
612 Ecosystems and Multiple Disturbances. *Microbiol. Mol. Biol. Rev.* **85**, e00026-20 (2021).
- 613 10. Lin, X. *et al.* Spatial and temporal dynamics of the microbial community in the Hanford unconfined
614 aquifer. *ISME J.* **6**, 1665–1676 (2012).
- 615 11. Zelaya, A. J. *et al.* High spatiotemporal variability of bacterial diversity over short time scales with
616 unique hydrochemical associations within a shallow aquifer. *Water Res.* **164**, 114917 (2019).
- 617 12. Zhou, Y., Kellermann, C. & Griebler, C. Spatio-temporal patterns of microbial communities in a
618 hydrologically dynamic pristine aquifer. *FEMS Microbiol. Ecol.* **81**, 230–242 (2012).
- 619 13. Yan, L. *et al.* Groundwater bacterial communities evolve over time in response to recharge. *Water*
620 *Res.* **201**, 117290 (2021).
- 621 14. Schroeter, S. A. *et al.* Hydroclimatic extremes threaten groundwater quality and stability. *Nat.*
622 *Commun. Accepted, to be published.*

- 623 15. Herrmann, M. *et al.* Predominance of Cand. Patescibacteria in groundwater is caused by their
624 preferential mobilization from soils and flourishing under oligotrophic conditions. *Front. Microbiol.*
625 **10**, 1407 (2019).
- 626 16. Herrmann, M., Lehmann, K., Totsche, K. U. & Küsel, K. Seepage-mediated export of bacteria from
627 soil is taxon-specific and driven by seasonal infiltration regimes. *Soil Biol. Biochem.* **187**, 109192
628 (2023).
- 629 17. Krüger, M. *et al.* Drought and rewetting events enhance nitrate leaching and seepage-mediated
630 translocation of microbes from beech forest soils. *Soil Biol. Biochem.* **154**, 108153 (2021).
- 631 18. Chaudhari, N. M., Pérez-Carrascal, O. M., Overholt, W. A., Totsche, K. U. & Küsel, K. Genome
632 streamlining in Parcubacteria transitioning from soil to groundwater. *Environ. Microbiome* **19**, 41
633 (2024).
- 634 19. Zhang, Y., Huang, C., Zhang, W., Chen, J. & Wang, L. The concept, approach, and future research
635 of hydrological connectivity and its assessment at multiscales. *Environ. Sci. Pollut. Res.* **28**, 52724–
636 52743 (2021).
- 637 20. Zhong, S. *et al.* Ecological differentiation and assembly processes of abundant and rare bacterial
638 subcommunities in karst groundwater. *Front. Microbiol.* **14**, 1111383 (2023).
- 639 21. Shade, A. *et al.* Fundamentals of Microbial Community Resistance and Resilience. *Front.*
640 *Microbiol.* **3**, (2012).
- 641 22. Chik, A. H. S. *et al.* Evaluation of groundwater bacterial community composition to inform
642 waterborne pathogen vulnerability assessments. *Sci. Total Environ.* **743**, 140472 (2020).
- 643 23. Campanale, C., Losacco, D., Triozzi, M., Massarelli, C. & Uricchio, V. F. An Overall Perspective
644 for the Study of Emerging Contaminants in Karst Aquifers. *Resources* **11**, 105 (2022).
- 645 24. Hose, G. C. *et al.* Assessing groundwater ecosystem health, status, and services. *Groundwater*
646 *Ecology and Evolution*, 501–524 (2023).
- 647 25. Fuhrman, J. A., Cram, J. A. & Needham, D. M. Marine microbial community dynamics and their
648 ecological interpretation. *Nat. Rev. Microbiol.* **13**, 133–146 (2015).
- 649 26. Aguilar, P. & Sommaruga, R. The balance between deterministic and stochastic processes in
650 structuring lake bacterioplankton community over time. *Mol. Ecol.* **29**, 3117–3130 (2020).

- 651 27. Riekeberg, E. & Powers, R. New frontiers in metabolomics: from measurement to insight.
652 *F1000Research* **6**, 1148 (2017).
- 653 28. Pemberton, J. A. *et al.* Untargeted characterisation of dissolved organic matter contributions to
654 rivers from anthropogenic point sources using direct-infusion and high-performance liquid
655 chromatography/Orbitrap mass spectrometry. *Rapid Commun. Mass Spectrom.* **34**, (2020).
- 656 29. Zerfaß, C. *et al.* Groundwater metabolome responds to recharge in fractured sedimentary strata.
657 *Water Res.* **223**, 118998 (2022).
- 658 30. Zerfaß, C., Lehmann, R., Ueberschaar, N., Totsche, K. U. & Pohnert, G. Current-use and legacy
659 contaminants evidence dissolved organic matter transfer and dynamics across a fractured-rock
660 groundwater recharge area. Preprint at <https://doi.org/10.1101/2024.06.03.596186> (2024).
- 661 31. Grasset, C., Groeneveld, M., Tranvik, L. J., Robertson, L. P. & Hawkes, J. A. Hydrophilic Species
662 Are the Most Biodegradable Components of Freshwater Dissolved Organic Matter. *Environ. Sci.*
663 *Technol.* **57**, 13463–13472 (2023).
- 664 32. Catalá, T. S., Shorte, S. & Dittmar, T. Marine dissolved organic matter: a vast and unexplored
665 molecular space. *Appl. Microbiol. Biotechnol.* **105**, 7225–7239 (2021).
- 666 33. Douglas, G. M. *et al.* PICRUSt2 for prediction of metagenome functions. *Nat. Biotechnol.* **38**, 669–
667 673 (2020).
- 668 34. Kohlhepp, B. *et al.* Aquifer configuration and geostructural links control the groundwater quality
669 in thin-bedded carbonate–siliciclastic alternations of the Hainich CZE, central Germany. *Hydrol.*
670 *Earth Syst. Sci.* **21**, 6091–6116 (2017).
- 671 35. Küsel, K. *et al.* How Deep Can Surface Signals Be Traced in the Critical Zone? Merging
672 Biodiversity with Biogeochemistry Research in a Central German Muschelkalk Landscape. *Front.*
673 *Earth Sci.* **4**, (2016).
- 674 36. Benk, S. A. *et al.* Fueling Diversity in the Subsurface: Composition and Age of Dissolved Organic
675 Matter in the Critical Zone. *Front. Earth Sci.* **7**, 296 (2019).
- 676 37. Nowak, M. E. *et al.* Carbon Isotopes of Dissolved Inorganic Carbon Reflect Utilization of Different
677 Carbon Sources by Microbial Communities in Two Limestone Aquifer Assemblages. *Hydrology*
678 *and Earth System Sciences* **21**, 4283–4300 (2017).

- 679 38. Schwab, V. F. *et al.* ¹⁴ C-free carbon is a major contributor to cellular biomass in geochemically
680 distinct groundwater of shallow sedimentary bedrock aquifers. *Water Resour. Res.* **55**, 2104–2121
681 (2019).
- 682 39. Geesink, P., Taubert, M., Jehmlich, N., Von Bergen, M. & Küsel, K. Bacterial necromass is rapidly
683 metabolized by heterotrophic bacteria and supports multiple trophic levels of the groundwater
684 microbiome. *Microbiol. Spectr.* **10**, e00437-22 (2022).
- 685 40. Mead, R. N. *et al.* Insights into dissolved organic matter complexity in rainwater from continental
686 and coastal storms by ultrahigh resolution Fourier transform ion cyclotron resonance mass
687 spectrometry. *Atmospheric Chem. Phys.* **13**, 4829–4838 (2013).
- 688 41. Farnleitner, A. H. *et al.* Bacterial dynamics in spring water of alpine karst aquifers indicates the
689 presence of stable autochthonous microbial endokarst communities. *Environ. Microbiol.* **7**, 1248–
690 1259 (2005).
- 691 42. Villeneuve, K., Violette, M. & Lazar, C. S. From recharge, to groundwater, to discharge Areas in
692 aquifer systems in Quebec (Canada): shaping of microbial diversity and community structure by
693 environmental factors. *Genes* **14**, 1 (2022).
- 694 43. Chow, C.-E. T. *et al.* Temporal variability and coherence of euphotic zone bacterial communities
695 over a decade in the Southern California Bight. *ISME J.* **7**, 2259–2273 (2013).
- 696 44. Cram, J. A. *et al.* Seasonal and interannual variability of the marine bacterioplankton community
697 throughout the water column over ten years. *ISME J.* **9**, 563–580 (2015).
- 698 45. Sharma, A. *et al.* Iron coatings on carbonate rocks shape the attached bacterial aquifer community.
699 *Sci. Total Environ.* **917**, 170384 (2024).
- 700 46. Lehman, R. M., Colwell, F. S. & Bala, G. A. Attached and unattached microbial communities in a
701 simulated basalt aquifer under fracture- and porous-flow conditions. *Appl. Environ. Microbiol.* **67**,
702 2799–2809 (2001).
- 703 47. Fuhrman, J. A. *et al.* Annually reoccurring bacterial communities are predictable from ocean
704 conditions. *Proc. Natl. Acad. Sci.* **103**, 13104–13109 (2006).
- 705 48. Gilbert, J. A. *et al.* Defining seasonal marine microbial community dynamics. *ISME J.* **6**, 298–308
706 (2012).

- 707 49. Yan, L. *et al.* Environmental selection shapes the formation of near-surface groundwater
708 microbiomes. *Water Res.* **170**, 115341 (2020).
- 709 50. Ning, D. *et al.* Environmental stress mediates groundwater microbial community assembly. *Nat.*
710 *Microbiol.* **9**, 490-501 (2024).
- 711 51. Merino, N. *et al.* Subsurface microbial communities as a tool for characterizing regional-scale
712 groundwater flow. *Sci. Total Environ.* **842**, 156768 (2022).
- 713 52. Roxburgh, S. H., Shea, K. & Wilson, J. B. The intermediate disturbance hypothesis: Patch dynamics
714 and mechanisms of species coexistence. *Ecology* **85**, 359–371 (2004).
- 715 53. Liu, X. & Salles, J. F. Lose-lose consequences of bacterial community-driven invasions in soil.
716 *Microbiome* **12**, 57 (2024).
- 717 54. Reif, J., Reifová, R., Skoracka, A. & Kuczyński, L. Competition-driven niche segregation on a
718 landscape scale: Evidence for escaping from syntopy towards allotopy in two coexisting sibling
719 passerine species. *J. Anim. Ecol.* **87**, 774–789 (2018).
- 720 55. Yeh, Y.-C. & Fuhrman, J. A. Contrasting diversity patterns of prokaryotes and protists over time
721 and depth at the San-Pedro Ocean Time series. *ISME Commun.* **2**, 36 (2022).
- 722 56. Langenheder, S., Berga, M., Östman, Ö. & Székely, A. J. Temporal variation of β -diversity and
723 assembly mechanisms in a bacterial metacommunity. *ISME J.* **6**, 1107–1114 (2012).
- 724 57. Portillo, M. C., Anderson, S. P. & Fierer, N. Temporal variability in the diversity and composition
725 of stream bacterioplankton communities: Temporal variability in stream bacterioplankton
726 communities. *Environ. Microbiol.* **14**, 2417–2428 (2012).
- 727 58. Yan, Q. *et al.* Nearly a decade-long repeatable seasonal diversity patterns of bacterioplankton
728 communities in the eutrophic Lake Donghu (Wuhan, China). *Mol. Ecol.* **26**, 3839–3850 (2017).
- 729 59. Zhang, L. *et al.* Untangling Microbiota Diversity and Assembly Patterns in the World’s Largest
730 Water Diversion Canal. *Water Res.* **204**, 117617 (2021).
- 731 60. Ferrera, I. *et al.* Seasonal and interannual variability of the free-living and particle-associated
732 bacteria of a coastal microbiome. *Environ. Microbiol. Rep.* **16**, e13299 (2024).
- 733 61. Aslam, R. A., Shrestha, S. & Pandey, V. P. Groundwater vulnerability to climate change: A review
734 of the assessment methodology. *Sci. Total Environ.* **612**, 853–875 (2018).

- 735 62. Allison, S. D. & Martiny, J. B. H. Resistance, resilience, and redundancy in microbial communities.
736 *Proc. Natl. Acad. Sci.* **105**, 11512–11519 (2008).
- 737 63. Shaffer, J. P. *et al.* Standardized multi-omics of Earth’s microbiomes reveals microbial and
738 metabolite diversity. *Nat. Microbiol.* **7**, 2128–2150 (2022).
- 739 64. Hu, A. *et al.* Microbial and Environmental Processes Shape the Link between Organic Matter
740 Functional Traits and Composition. *Environ. Sci. Technol.* **56**, 10504–10516 (2022).
- 741 65. Sumranwanich, T. *et al.* Evaluating lignin degradation under limited oxygen conditions by bacterial
742 isolates from forest soil. *Sci. Rep.* **14**, 13350 (2024).
- 743 66. Osterholz, H. *et al.* Terrigenous dissolved organic matter persists in the energy-limited deep
744 groundwaters of the Fennoscandian Shield. *Nat. Commun.* **13**, 4837 (2022).
- 745 67. Shen, Y., Chapelle, F. H., Strom, E. W. & Benner, R. Origins and bioavailability of dissolved
746 organic matter in groundwater. *Biogeochemistry* **122**, 61–78 (2015).
- 747 68. Lapworth, D. J., Baran, N., Stuart, M. E. & Ward, R. S. Emerging organic contaminants in
748 groundwater: A review of sources, fate and occurrence. *Environ. Pollut.* **163**, 287–303 (2012).
- 749 69. Musgrove, M., Jurgens, B. C. & Opsahl, S. P. Karst groundwater vulnerability determined by
750 modeled age and residence time tracers. *Geophys. Res. Lett.* **50**, e2023GL102853 (2023).
- 751 70. Mehrshad, M. *et al.* Energy efficiency and biological interactions define the core microbiome of
752 deep oligotrophic groundwater. *Nat. Commun.* **12**, 4253 (2021).
- 753 71. Soares, A. *et al.* A global perspective on bacterial diversity in the terrestrial deep subsurface.
754 *Microbiology* **169**, (2023).
- 755 72. Lehmann, K., Lehmann, R. & Totsche, K. U. Event-driven dynamics of the total mobile inventory
756 in undisturbed soil account for significant fluxes of particulate organic carbon. *Sci. Total Environ.*
757 **756**, 143774 (2021).
- 758 73. Klindworth, A. *et al.* Evaluation of general 16S ribosomal RNA gene PCR primers for classical and
759 next-generation sequencing-based diversity studies. *Nucleic Acids Res.* **41**, e1–e1 (2013).
- 760 74. Kumar, S. *et al.* Thiosulfate- and hydrogen-driven autotrophic denitrification by a microbial
761 consortium enriched from groundwater of an oligotrophic limestone aquifer. *FEMS Microbiol. Ecol.*
762 **94**, (2018).

- 763 75. Rughöft, S. *et al.* Community Composition and Abundance of Bacterial, Archaeal and Nitrifying
764 Populations in Savanna Soils on Contrasting Bedrock Material in Kruger National Park, South
765 Africa. *Front. Microbiol.* **7**, (2016).
- 766 76. Dittmar, T., Koch, B., Hertkorn, N. & Kattner, G. A simple and efficient method for the solid-phase
767 extraction of dissolved organic matter (SPE-DOM) from seawater. *Limnol. Oceanogr. Methods* **6**,
768 230–235 (2008).
- 769 77. Laszakovits, J. R. & MacKay, A. A. Data-Based Chemical Class Regions for Van Krevelen
770 Diagrams. *J. Am. Soc. Mass Spectrom.* **33**, 198–202 (2022).
- 771 78. R Core Team. R: A language and environment for statistical computing. (2022).
- 772 79. Martin, M. Cutadapt removes adapter sequences from high-throughput sequencing reads. *EMBnet.*
773 *journal* **17**, 10–12 (2011).
- 774 80. Callahan, B. J. *et al.* DADA2: High-resolution sample inference from Illumina amplicon data. *Nat.*
775 *Methods* **13**, 581–583 (2016).
- 776 81. Quast, C. *et al.* The SILVA ribosomal RNA gene database project: improved data processing and
777 web-based tools. *Nucleic Acids Res.* **41**, D590–D596 (2013).
- 778 82. McMurdie, P. J. & Holmes, S. phyloseq: An R package for reproducible interactive analysis and
779 graphics of microbiome census data. *PLoS ONE* **8**, e61217 (2013).
- 780 83. Price, M. N., Dehal, P. S. & Arkin, A. P. FastTree: Computing Large Minimum Evolution Trees
781 with Profiles instead of a Distance Matrix. *Mol. Biol. Evol.* **26**, 1641–1650 (2009).
- 782 84. Edgar, R. C. Muscle5: High-accuracy alignment ensembles enable unbiased assessments of
783 sequence homology and phylogeny. *Nat. Commun.* **13**, 6968 (2022).
- 784 85. Capella-Gutiérrez, S., Silla-Martínez, J. M. & Gabaldón, T. trimAl: a tool for automated alignment
785 trimming in large-scale phylogenetic analyses. *Bioinformatics* **25**, 1972–1973 (2009).
- 786 86. Knights, D. *et al.* Bayesian community-wide culture-independent microbial source tracking. *Nat.*
787 *Methods* **8**, 761–763 (2011).
- 788 87. Ning, D. *et al.* A quantitative framework reveals ecological drivers of grassland microbial
789 community assembly in response to warming. *Nat. Commun.* **11**, 4717 (2020).

790 88. Yurekten, O. *et al.* MetaboLights: open data repository for metabolomics. *Nucleic Acids Res.* **52**,
791 D640–D646 (2024).

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801 **Author contributions**

802 HW and KK conceptualized the manuscript. KK, MH, and HW provided microbiome data. SAS, AI,
803 GG, and SET provided DOC and bulk DOM data. CZ and GP provided metabolomics data. RL, KL,
804 and KUT managed field installations and provided hydrochemical parameter analyses. HW processed
805 data, performed bioinformatic analyses and wrote the manuscript. All authors discussed the results and
806 implications and commented on the manuscript at all stages.

807 **Competing interests**

808 The authors declare no competing interests.

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